

Estimating surface rain rate from satellite passive microwave observations

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Outline

- **How does passive MW works?**
 - Passive MW in a nutshell

- **Inverse algorithm**
 - basic concepts

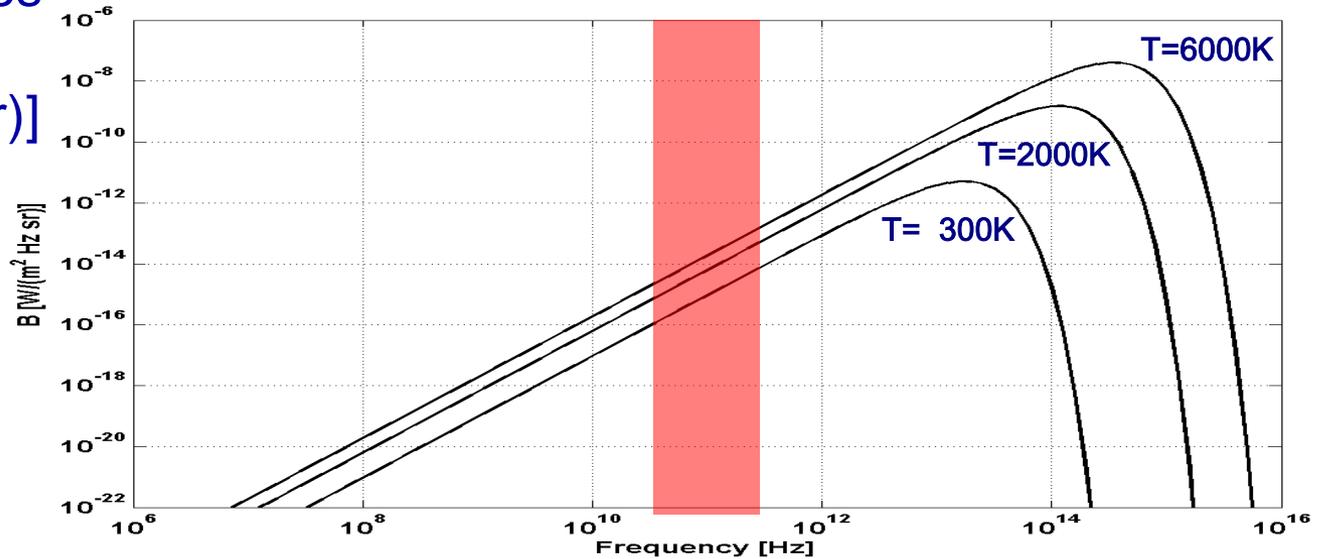
- **Examples**
 - Near-real-time
 - Validation



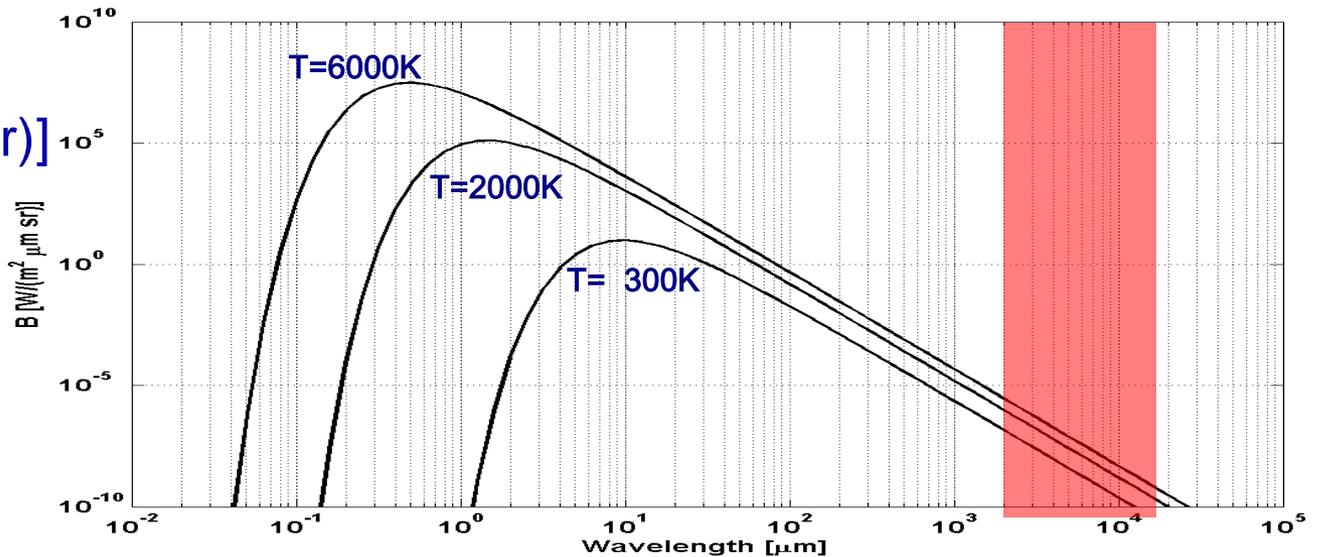
Emission of MW radiation

Planck's curves

$B[\text{W}/(\text{m}^2 \text{ Hz sr})]$
 vs. $F[\text{Hz}]$



$B[\text{W}/(\text{m}^2 \mu\text{m sr})]$
 vs. $\lambda[\mu\text{m}]$



Emission of MW radiation

PLANCK FUNCTION APPROXIMATIONS

high ν (IR) $h\nu \gg KT \Rightarrow \frac{h\nu}{KT} \gg 1 \Rightarrow \frac{1}{e^{\frac{h\nu}{KT}} - 1} \approx \frac{1}{e^{\frac{h\nu}{KT}}} = e^{-\frac{h\nu}{KT}} \Rightarrow B_\nu \approx \frac{2h\nu^3}{c^2} e^{-\frac{h\nu}{KT}}$

low ν (MW) $h\nu \ll KT \Rightarrow \frac{h\nu}{KT} \ll 1 \Rightarrow \frac{1}{e^{\frac{h\nu}{KT}} - 1} \approx \frac{1}{\frac{h\nu}{KT}} = \frac{KT}{h\nu} \Rightarrow B_\nu \approx 2KT \frac{\nu^2}{c^2} = \frac{2KT}{\lambda^2}$

$$B_\nu = \frac{2h\nu^3}{c^2} \frac{1}{e^{\frac{h\nu}{KT}} - 1} \quad \text{Planck function}$$

$$B_\nu \approx \frac{2h\nu^3}{c^2} e^{-\frac{h\nu}{KT}} \quad \text{Wien approximation (IR)}$$

$$B_\nu \approx 2KT \frac{\nu^2}{c^2} \quad \text{Reyleigh-Jeans approximation (MW)}$$



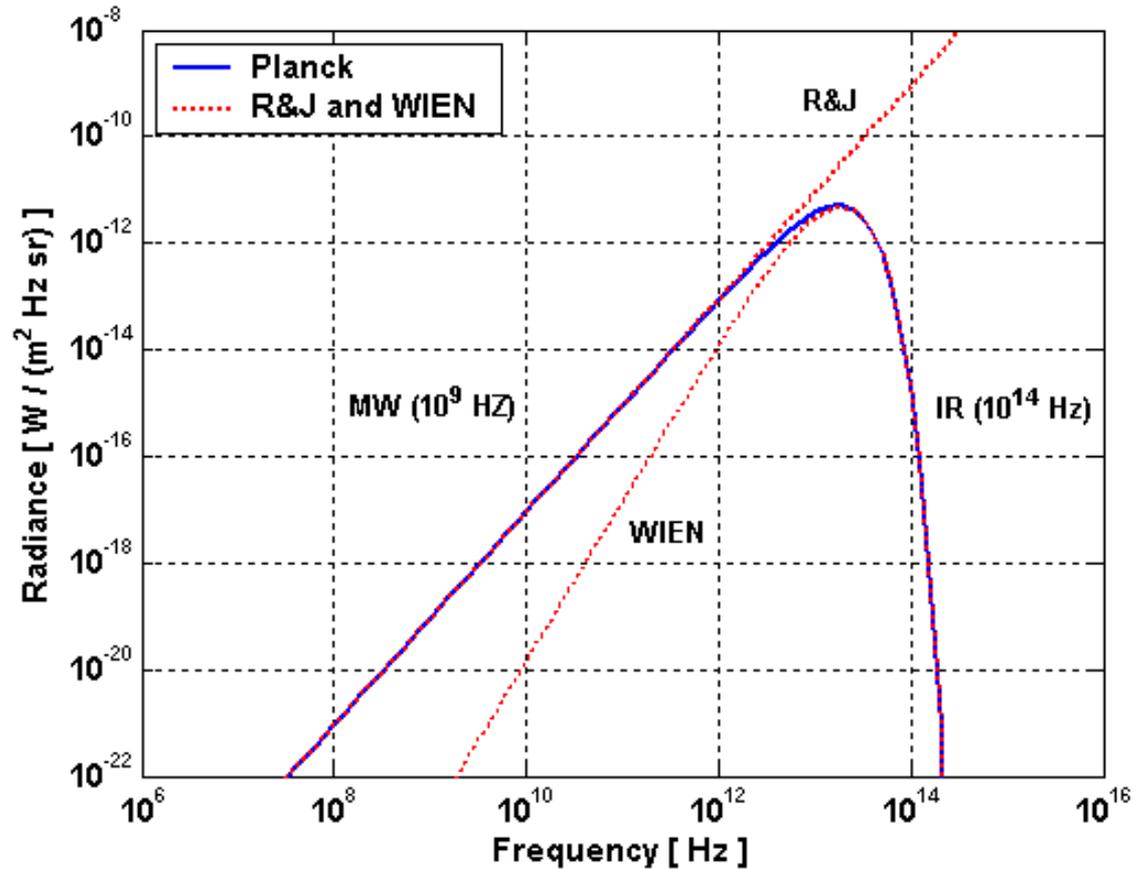
Emission of MW radiation

PLANCK FUNCTION APPROXIMATIONS

$$B_\nu = \frac{2h\nu^3}{c^2} \frac{1}{e^{\frac{h\nu}{KT}} - 1}$$

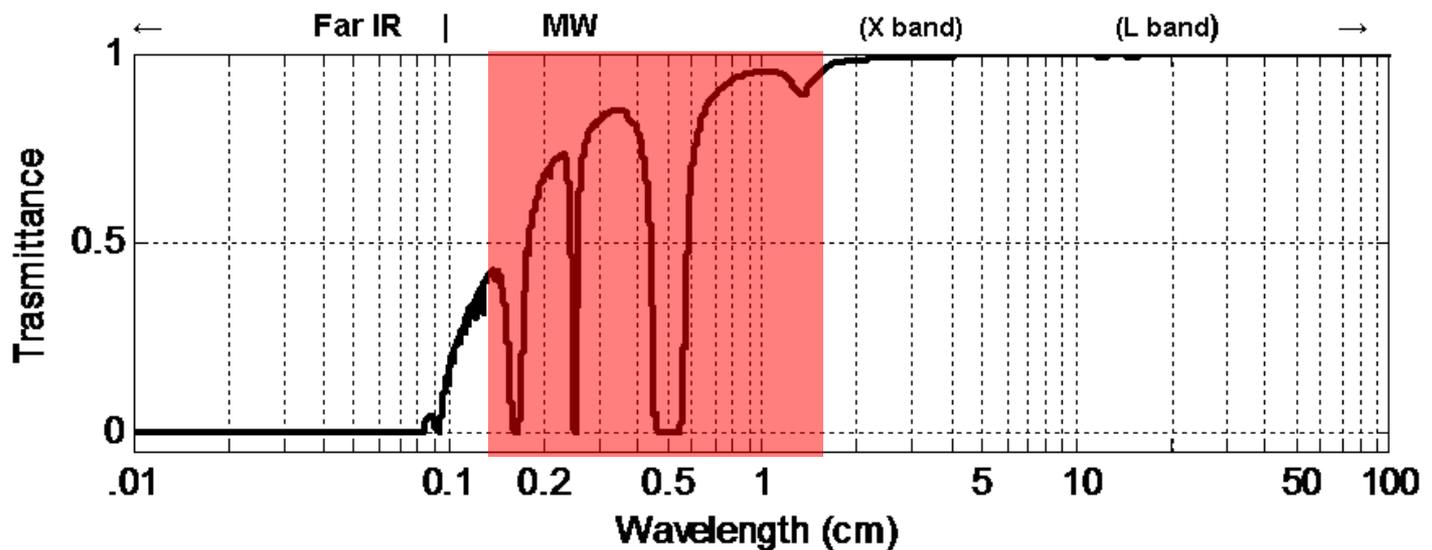
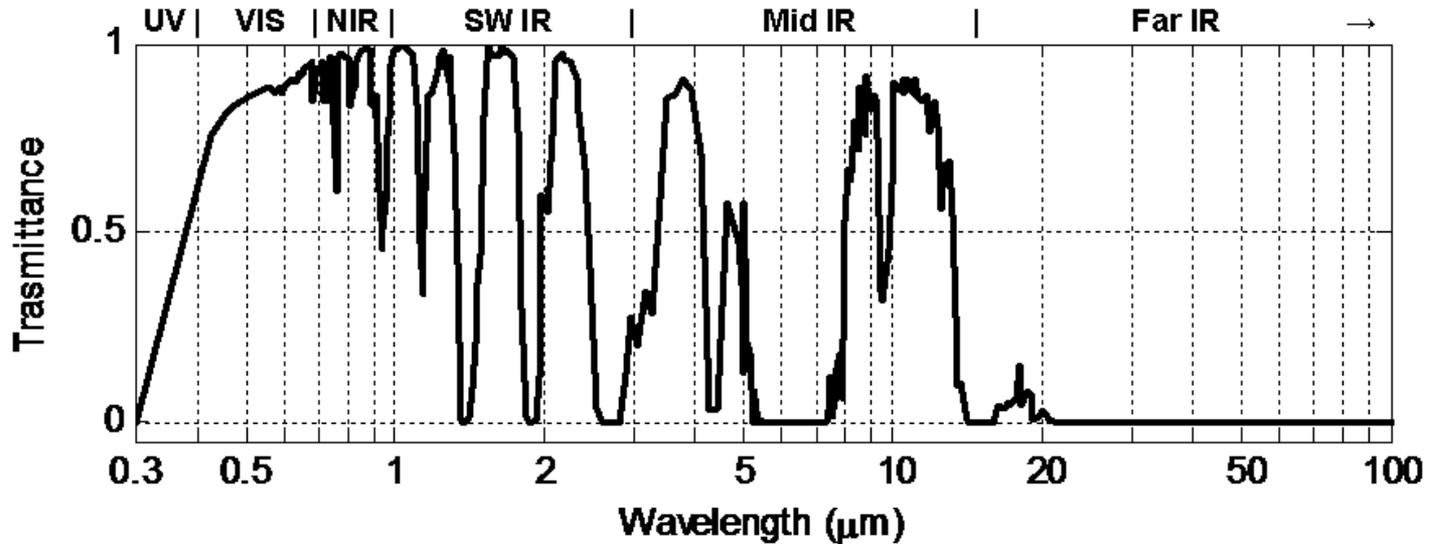
$$B_\nu \approx \frac{2h\nu^3}{c^2} e^{-\frac{h\nu}{KT}}$$

$$B_\nu \approx 2KT \frac{\nu^2}{c^2}$$



Transmission of MW radiation

- Atmospheric Transmittance (due to absorption only)



Scattering of MW radiation

- EM wave scattering happens when atmospheric particles deviate incident radiation from its original direction of propagation.
- How much scattering actually happens depends on radiation λ (or ν), molecule/particle characteristics (abundance, size, ...).

□ Atmospheric particle	typical size	
○ Molecule:	$10^{-4} - 10^{-3}$	μm
○ Aerosol:	$0.1 - 10$	μm
○ Hydrometeors:	$10^{-3} - 10^{+5}$	μm
○ Haze:	$10^{-3} - 1$	μm
○ Fog:	$1 - 100$	μm
○ Cloud:	$1 - 10^3$	μm
○ Ice crystals:	$1 - 5 \cdot 10^3$	μm
○ Rain:	$0.01 - 1$	cm
○ Snow:	$1 - 5$	cm

Increasing size



Scattering of MW radiation

Radiation scattering depends on the ratio particle size over wavelength

Dimension parameter : $x=2\pi r/\lambda$

If	$r \ll \lambda$	$x \rightarrow 0$	negligible scattering
If	$r \sim \lambda$	$x \ll 1$	Rayleigh limit (simplified solution)
If	$r \sim \lambda$	$x \sim 1$	Mie conditions (rigorous solution)
If	$r \gg \lambda$	$x \rightarrow \infty$	geometric optics limit (non-selective)

Increasing x



Rayleigh scattering: small particles ($x \ll 1$); molecules, fine dust,
Ex: blue sky (visible)

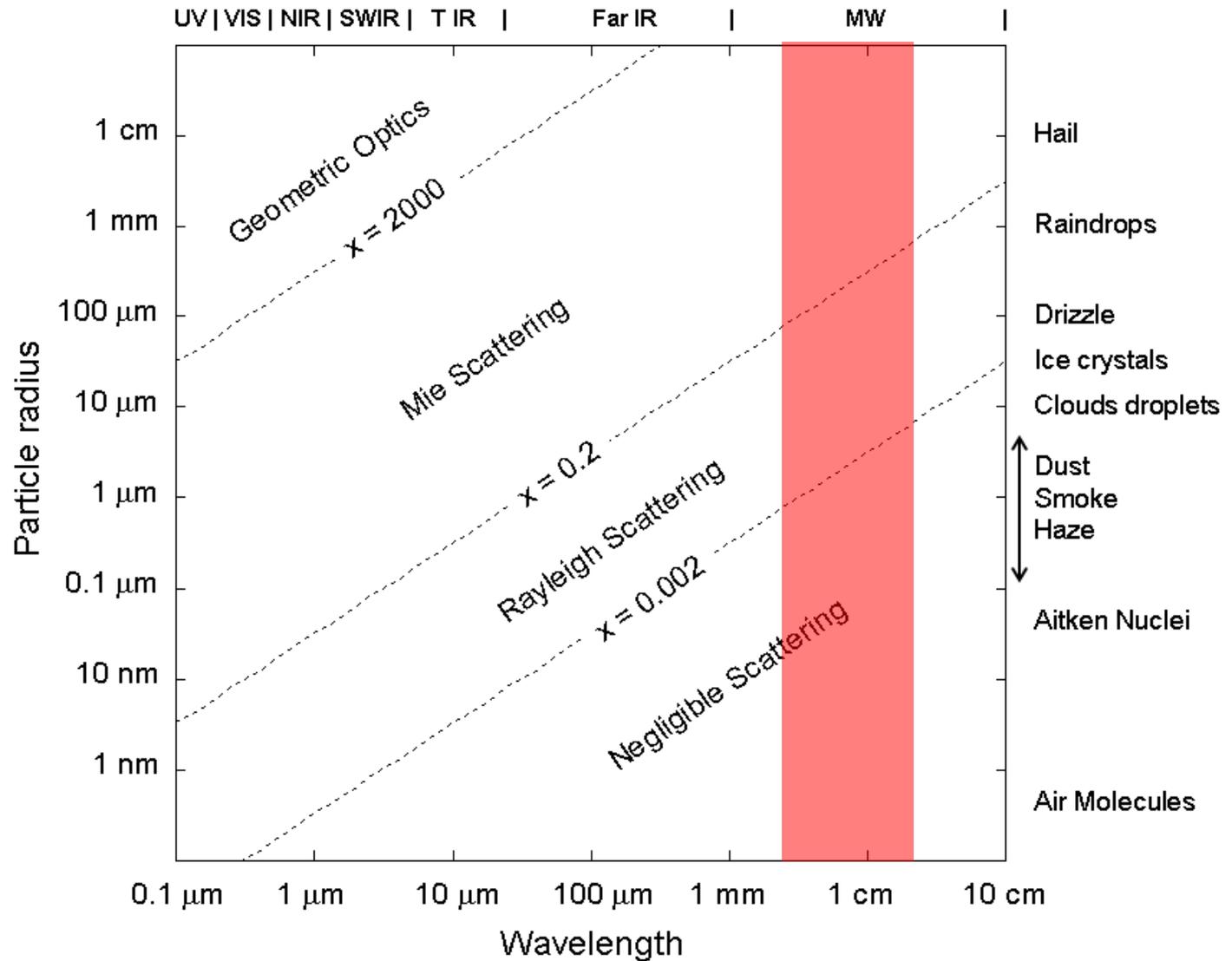
Mie scattering: medium particles ($x \sim 1$); dust, pollen, hydrometeors
Ex: rain (mw)

Non-selective scattering: large particles ($x \gg 1$); hydrometeors,
Ex: white clouds (visible)



Scattering of MW radiation

□ Scattering regimes



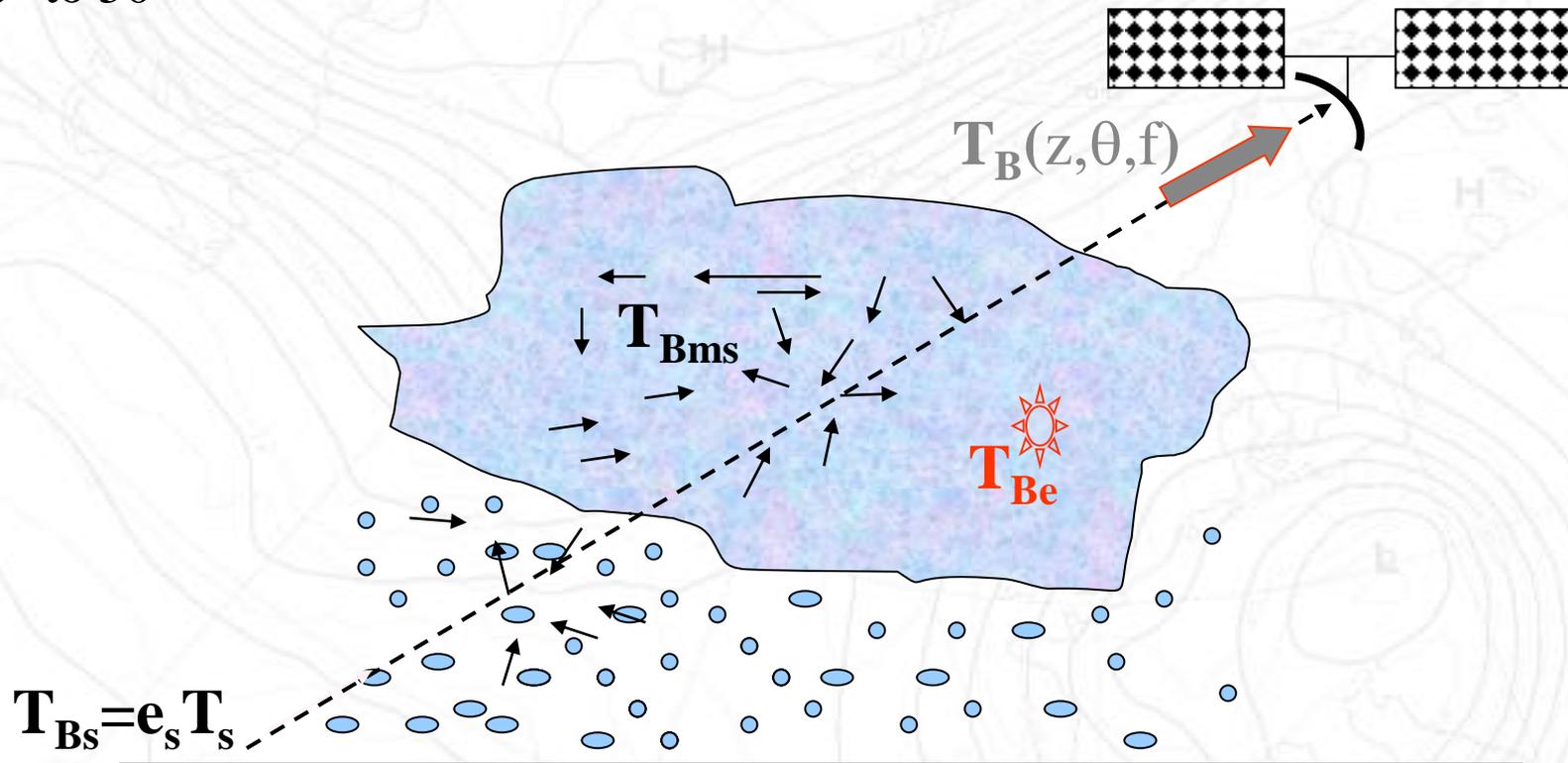
FORWARD PROBLEM

Physical basis

T_B : brightness temperature

$f=10-160$ GHz

$\theta=0^\circ$ to 50°



e_s : surface emissivity
 T_s : surface temperature

T_{Be} : emission T_B
 T_{Bms} : multiple scattering T_B

FORWARD PROBLEM

Radiative transfer equation

- T_B in a plane-parallel medium: non-scattering case

$$\frac{dT_B(\tau, \Omega)}{d\tau} = -\underbrace{T_B(\tau, \Omega)}_{\text{Extinction}} + \underbrace{T(\tau)}_{\text{Emission}}$$

⇒ **Ordinary differential equation: linearization of F**

⇒ **Inverse problem as a *Fredholm integral equation* (e.g., temperature retrieval)**

- T_B in a plane-parallel medium: scattering case

$$\frac{dT_B(\tau, \Omega)}{d\tau} = -\underbrace{T_B(\tau, \Omega)}_{\text{Extinction}} + \underbrace{\frac{w}{4\pi} \int_{4\pi} p(\Omega, \Omega') T_B(\tau, \Omega) d\Omega'}_{\text{Multiple scattering}} + \underbrace{(1-w)T(\tau)}_{\text{Emission}}$$

⇒ **Integro-differential equation: strongly non-linear F (e.g., rainfall retrieval)**



Precipitation in MW — Theory/Basis

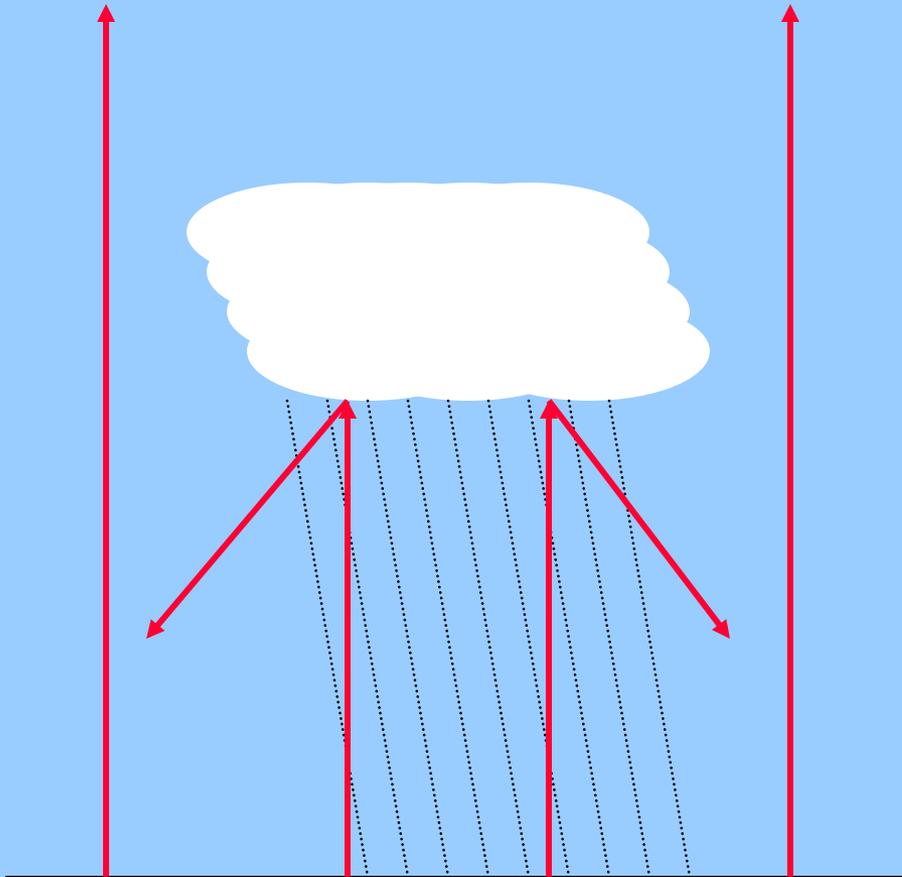
- **Scattering signal**
 - ice in clouds scatters terrestrial radiation (cold areas over warm bckg)
 - Rainfall rates are related to the magnitude of the resulting brightness temperature depression.
 - **Strength**: can be applied to high-frequency channels: works over both land and ocean
 - **Weakness**: poor at detecting precipitation clouds with little or no ice (e.g. warm orographic clouds in the tropics)

- **Emission signal**
 - water in clouds emits radiation, (warm areas over cold bckg, e.g. ocean)
 - Rainfall rates are related to the magnitude of the resulting brightness temperature difference
 - **Strength**: sensitive to clouds with little or no ice
 - **Weakness**: must know terrestrial radiances without cloud beforehand; generally applicable over oceans but not land



Lower T_b
above cloud

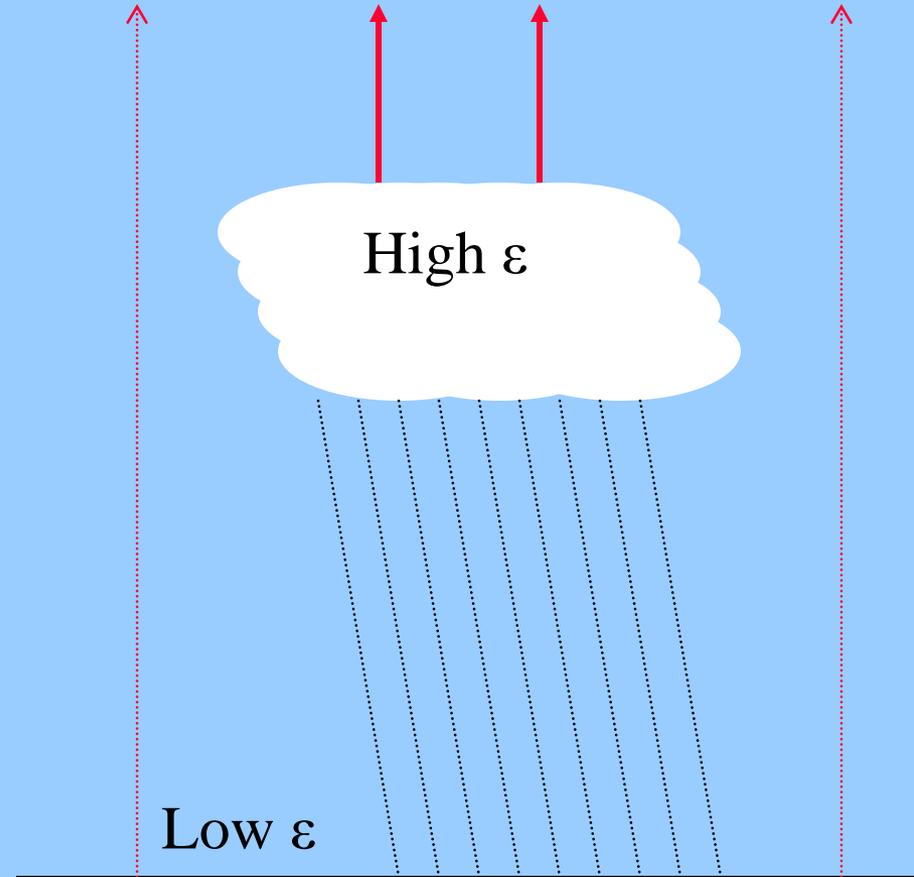
Higher T_b above
clear air



Scattering

Higher T_b
above cloud

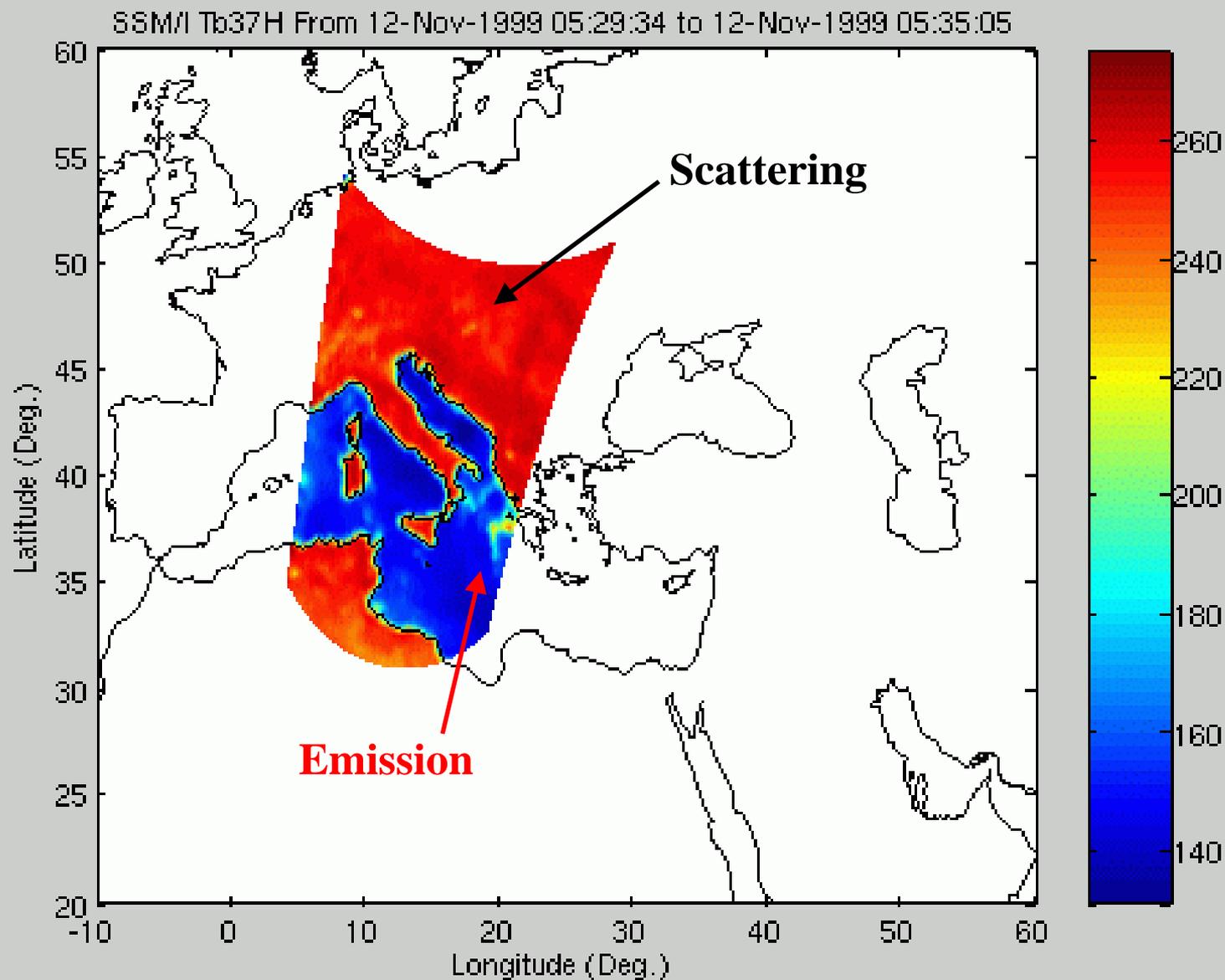
Lower T_b above
clear air



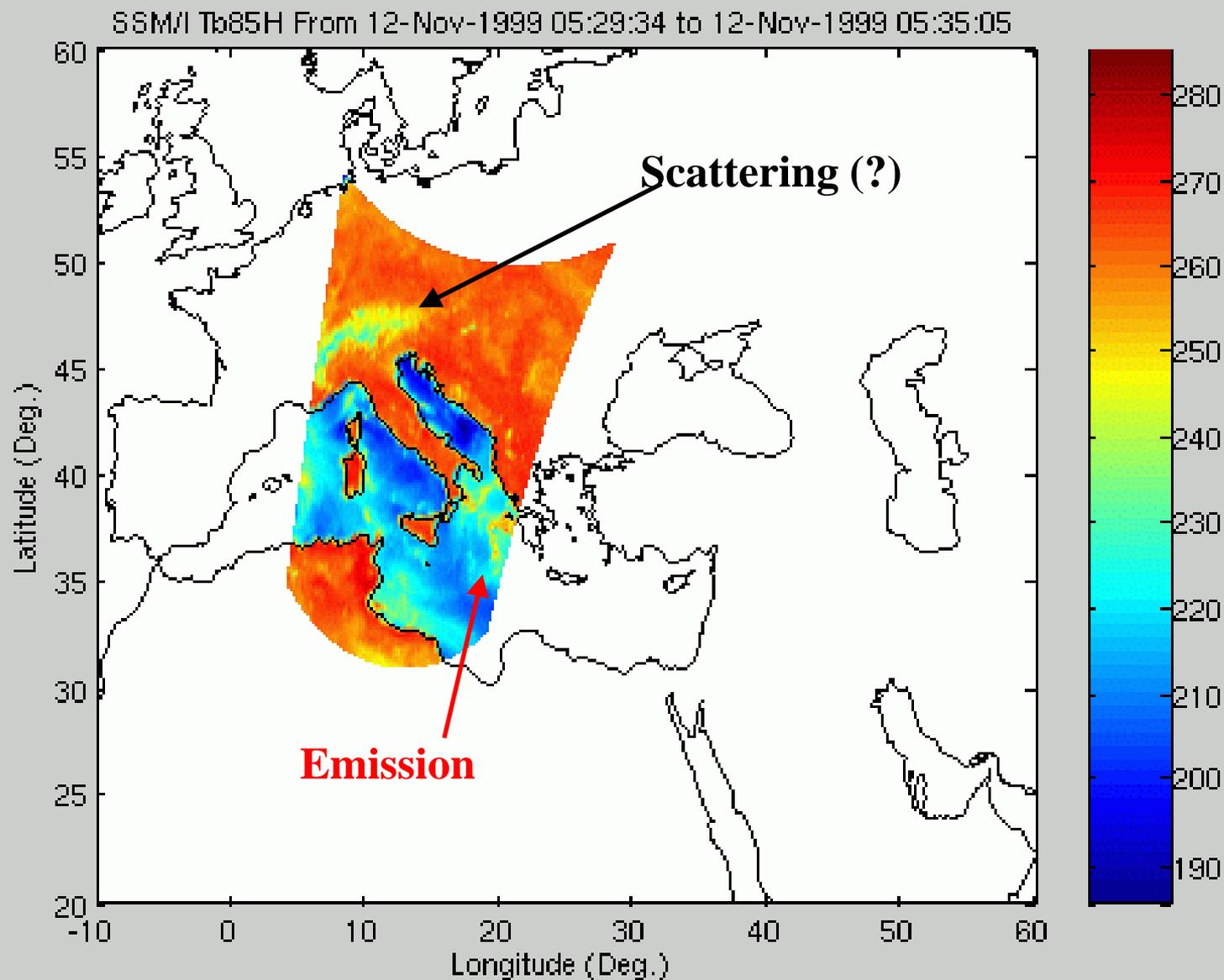
Emission (over ocean)



Tb at 37 GHz H

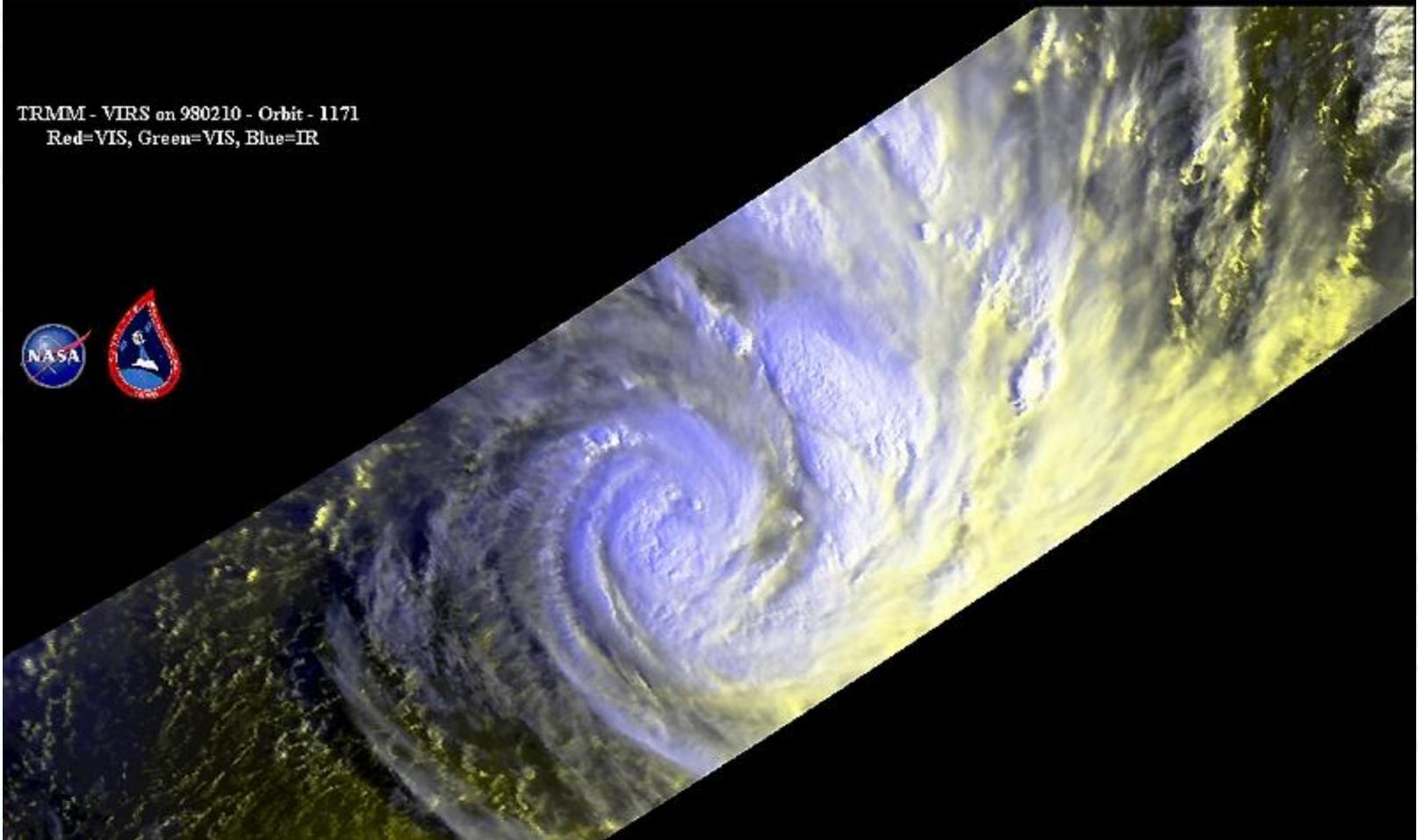


Tb at 85 GHz H

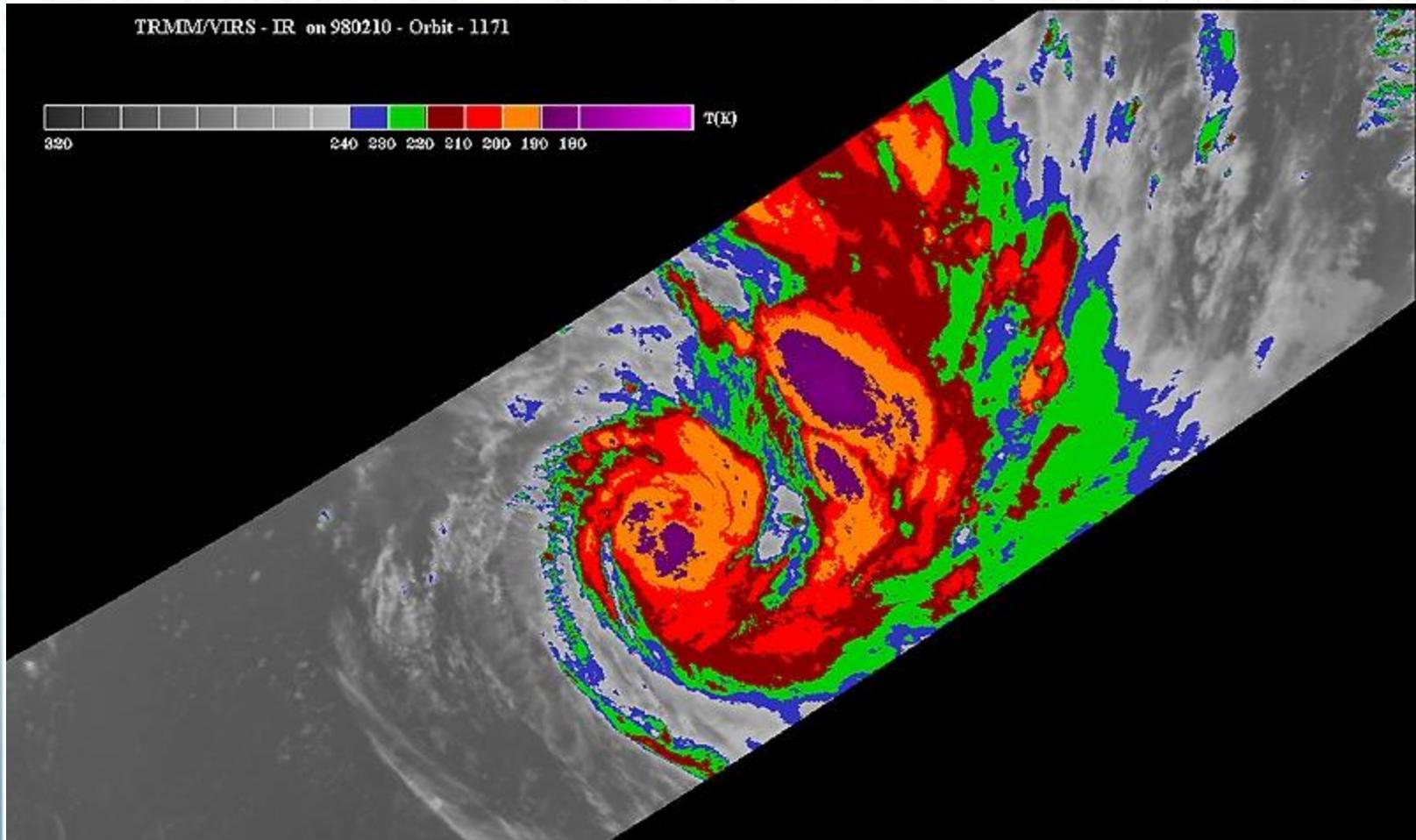


TRMM VIRS (RGB)

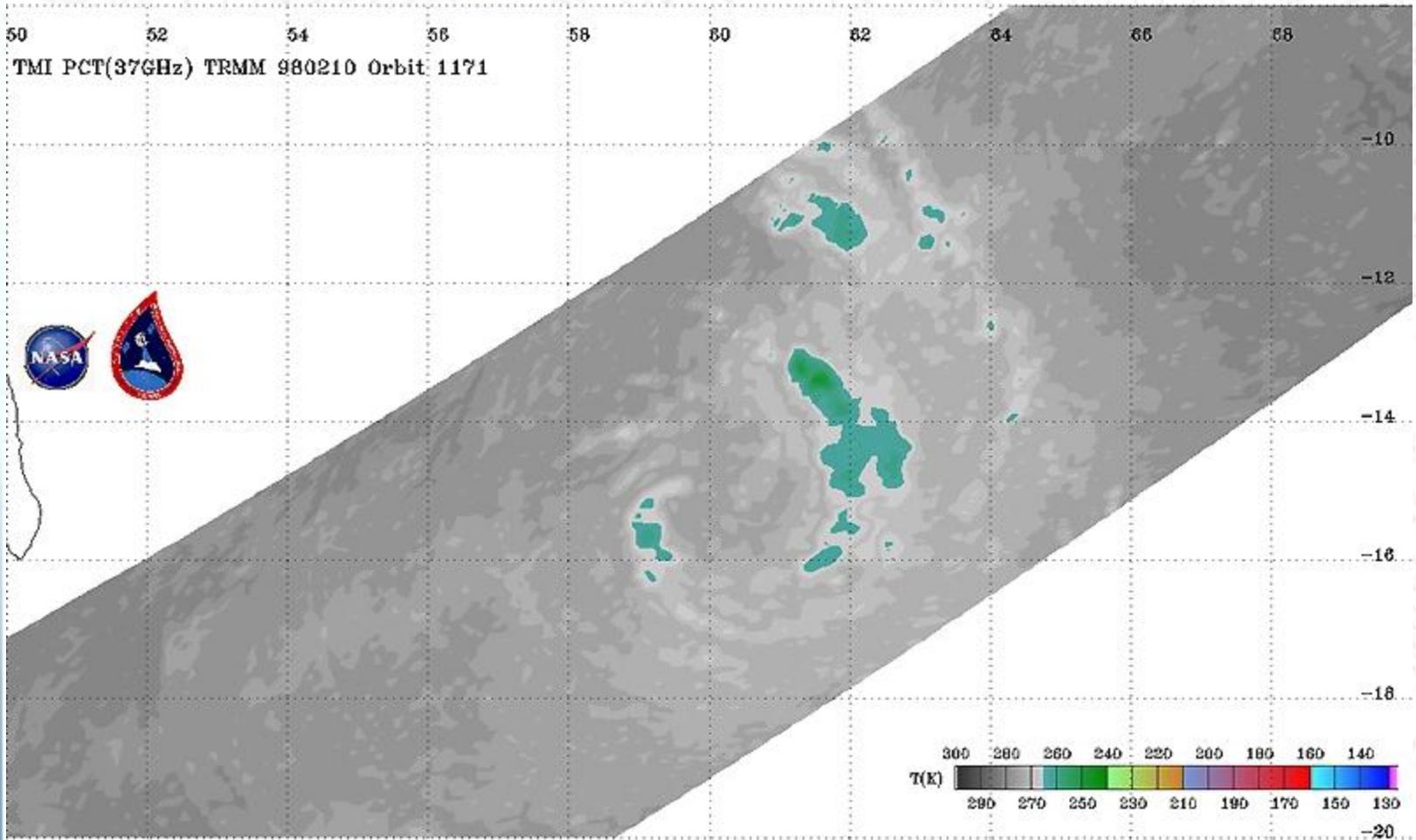
TRMM - VIRS on 980210 - Orbit - 1171
Red=VIS, Green=VIS, Blue=IR



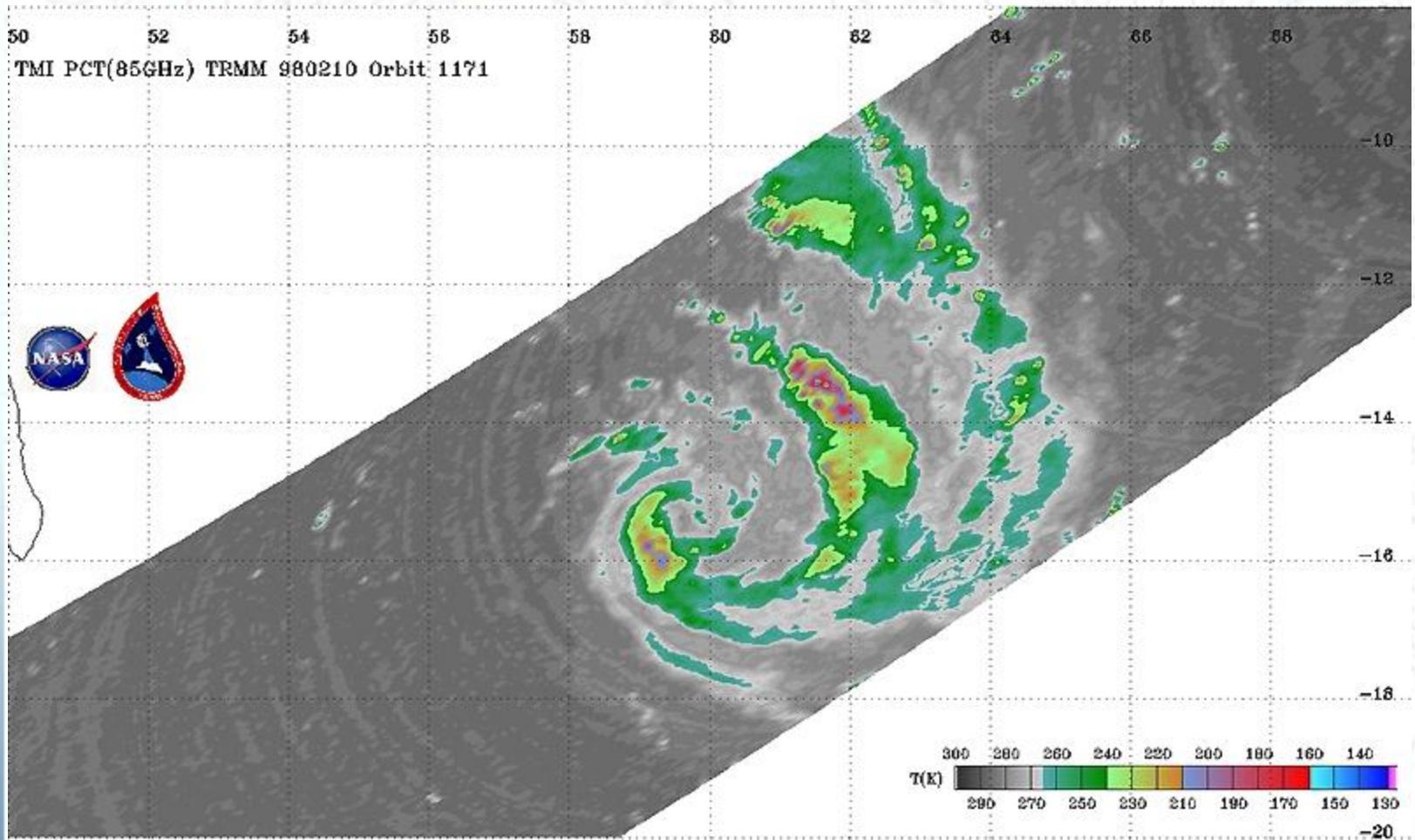
TRMM VIRS (IR)



TRMM MICROWAVE IMAGER (37GHz)



TRMM MICROWAVE IMAGER (85GHz)

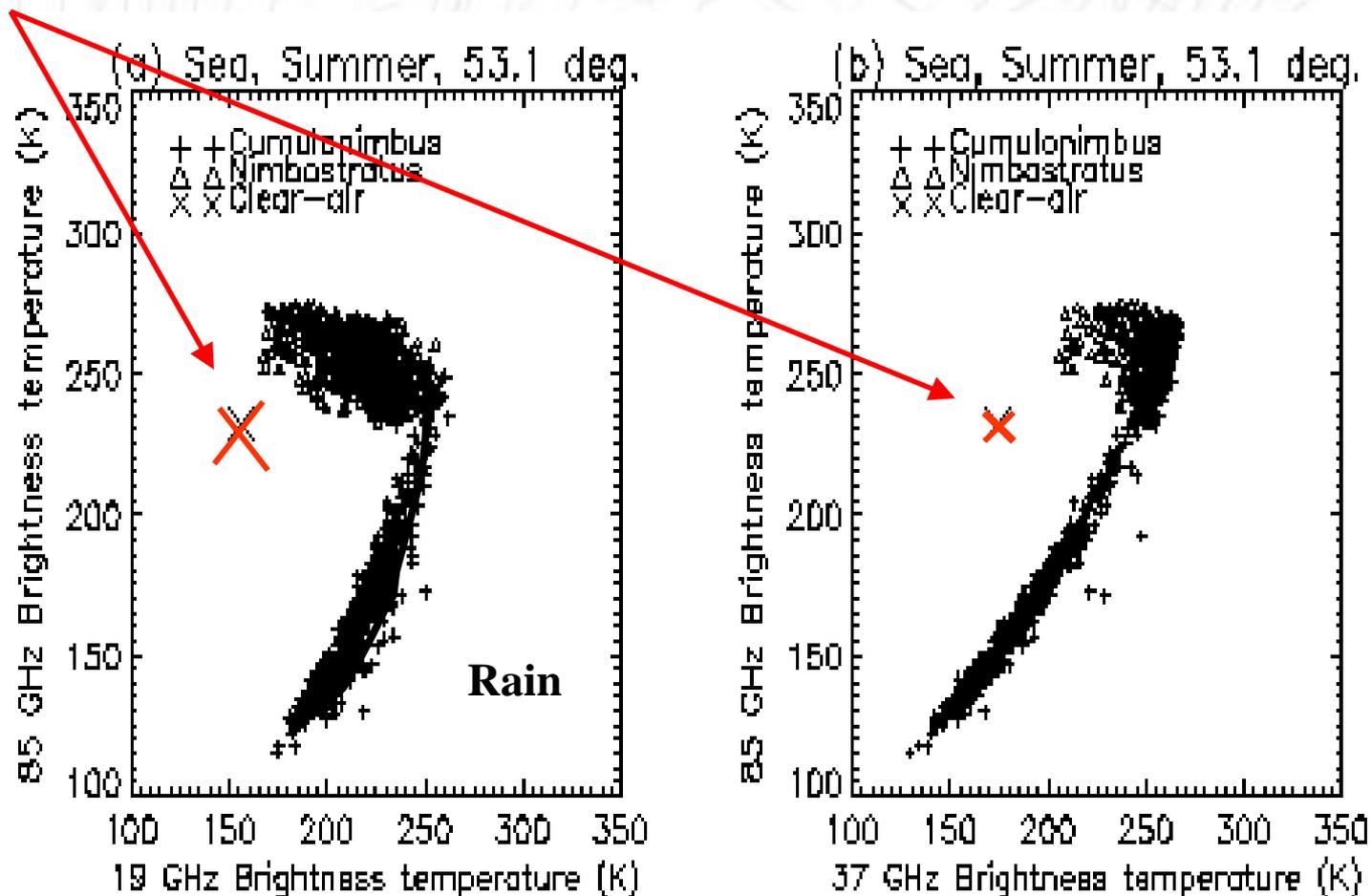


FORWARD PROBLEM

Radiative transfer models

Clear air

T_B simulations over ocean

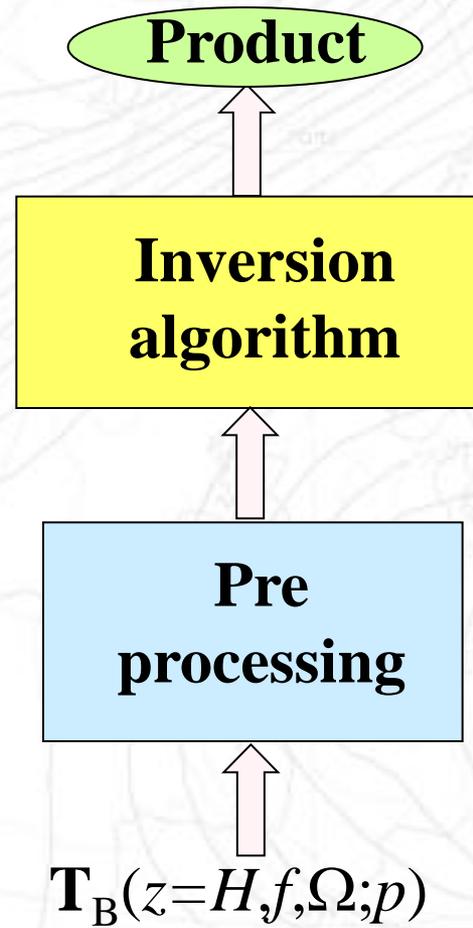
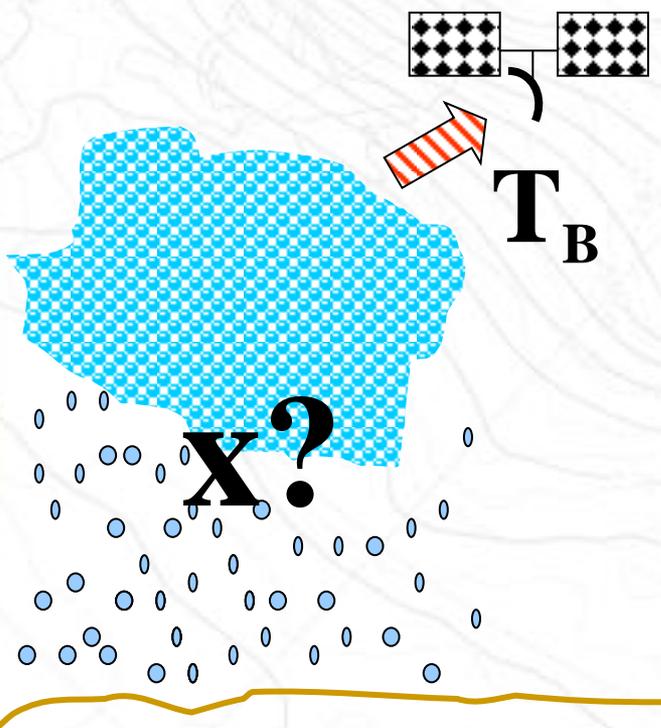


T_B [e.g., 8 T_B 's H/V at 10, 19, 37, 85 GHz]



INVERSE PROBLEM

Rainfall retrieval



- Hydrometeor profile
 - Latent heat
 - Rain rate
-
- Training data set
 - A priori information
 - Ancillary data
-
- Geo-location
 - Calibration
 - Filtering
 - Deconvolution



Instruments and Algorithms

Satellite passive microwave observations

- Rainfall observations from satellite are performed operationally by several agencies using data from MW radiometers as SSM/I, AMSU-A/B, MHS.
- Here we use the surface rain rate estimated operationally at IMAA-CNR in collaboration with CETEMPS
 - **Algorithm:** PEMW (Precipitation Estimation from MicroWave observations) developed at IMAA-CNR
 - **Instrument:** MHS and AMSU-B
 - **Satellites:** NOAA N16-18-19; EUMETSAT MetOp A
 - **Orbit:** polar low-earth-orbit (LEO)

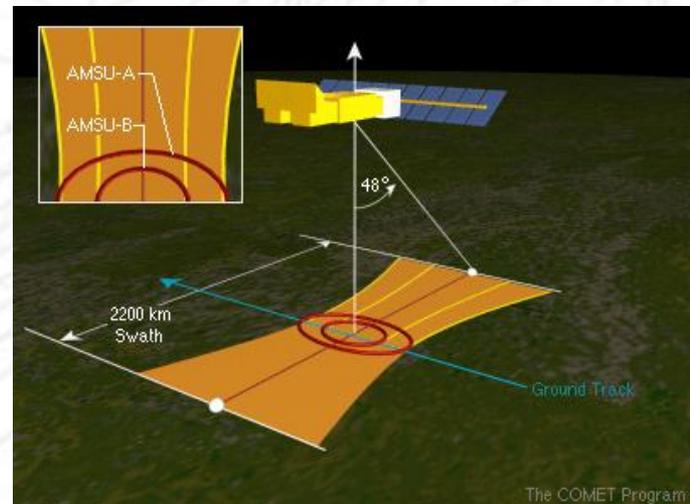


Instruments and Algorithms

Satellite passive microwave observations

AMSU-B (Advanced Microwave Sounding Unit B) MHS (Microwave Humidity Sounder)

- ❑ Cross-track linear scanning MW radiometers
- ❑ 5 channels
 - 2 window
 - 3 opaque
- ❑ Antenna beam width: 1.1°
- ❑ Spatial resolution: 16-17 km at nadir
- ❑ 90 consecutive FOV
- ❑ Coverage: Twice daily full global coverage



AMSU-B

# chan	f (GHz)	# passband	bandwidth per passband (MHz)	POL
1	89.0+-0.9	2	1000	V
2	150.0+-0.9	2	1000	V
3	183.31+-1.00	2	500	V
4	183.31+-3.00	2	1000	V
5	183.31+-7.00	2	2200	V

MHS

f (GHz)	# passband	bandwidth per passband (MHz)	POL
89.0	1	2800	V
150.0	1	2800	V
183.31+-1.00	2	500	H
183.31+-3.00	2	1000	H
190.31	1	2200	V



Instruments and Algorithms

Satellite passive microwave observations

Precipitation Estimation from AMSU-B / MHS

- Features:
 - Increased spatial resolution wrt lower freq instruments
 - Increased sensitivity to light rain wrt lower freq instruments
 - Good temporal coverage: some 7-10 daily overpasses with AMSU-B / MHS currently flying on 4 platforms



Instruments and Algorithms

PEMW Algorithm

- PEMW automatically selects the most appropriate atmospheric scenario from a pool of 81
- Each scenario is associated with coefficients which fit a model between the observations and rain rate
- Scenarios (and associated coefficients) were identified based on a data set of simulations and observations

Retrieval procedure

Di Tomaso et al., J. Geoph. Res., 2009

INPUT (MHS or AMSU-B)

$(TB_1; TB_2; TB_3; TB_4; TB_5)$

TB Differences

$$\Delta_1 = TB_1 - TB_2$$

$$\Delta_2 = TB_3 - TB_5$$

$$\Delta_3 = TB_4 - TB_5$$

Model

$$f(rr, \Delta_i | \theta_{i_1}, \dots, \theta_{i_j}, \dots, \theta_{i_c})$$

$$i = 1, 2, 3$$

$$\Theta_k = \{\theta_{1_1}, \dots, \theta_{1_c}, \dots, \theta_{3_1}, \dots, \theta_{3_c}\}$$

$$k = 1, \dots, s$$

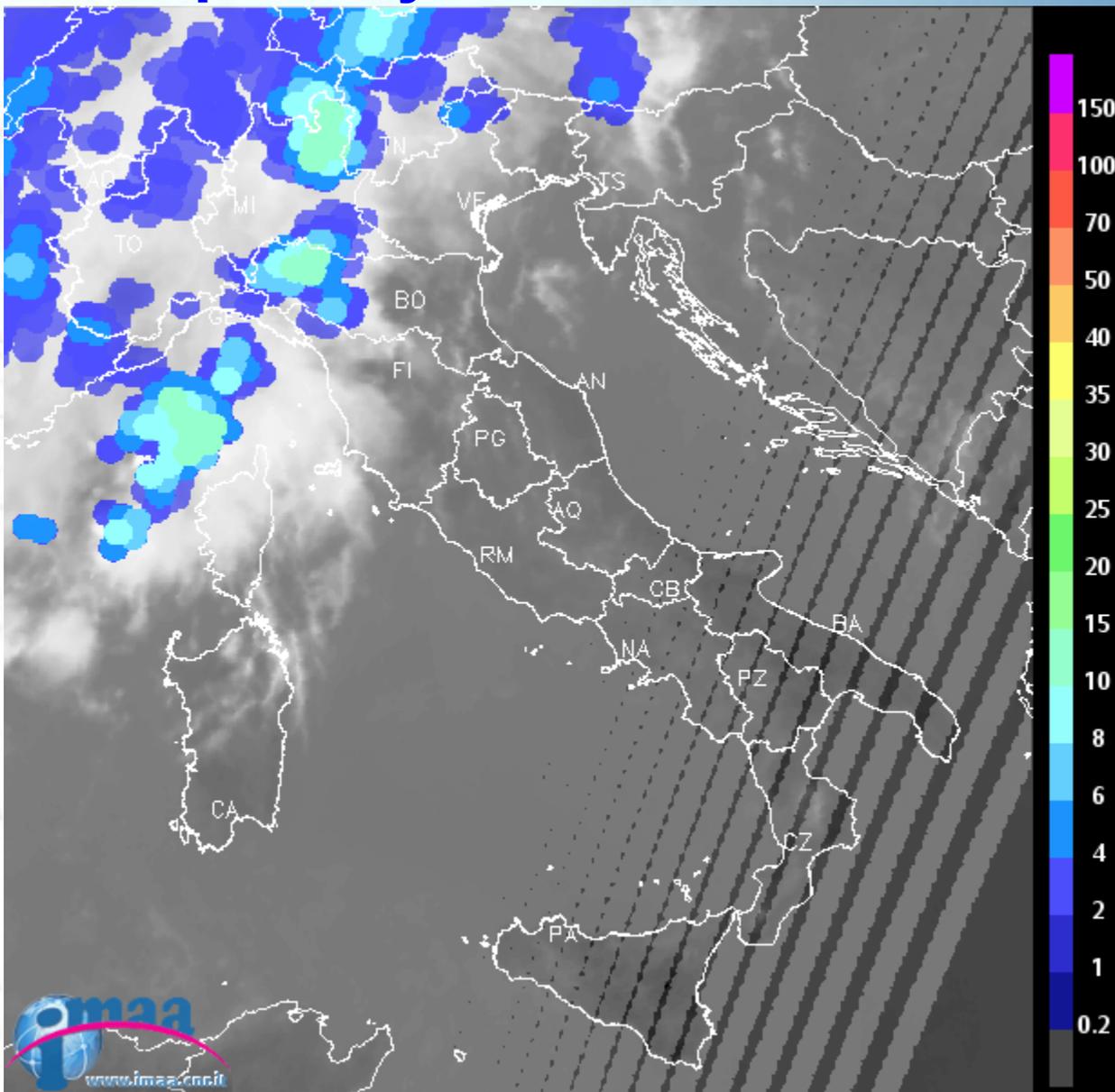
OUTPUT

Surface rain rate (mm/h)

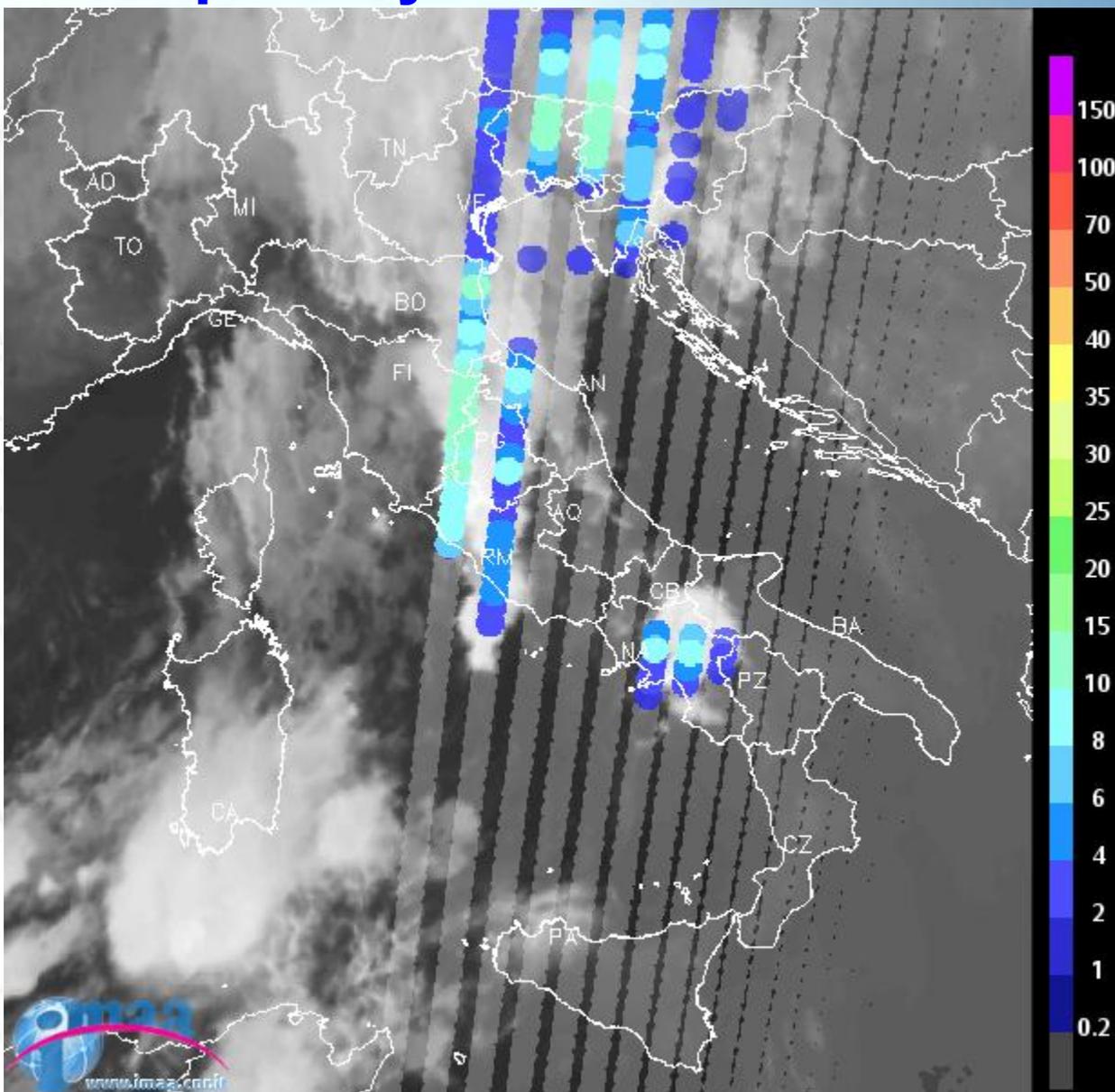
Scenario selection



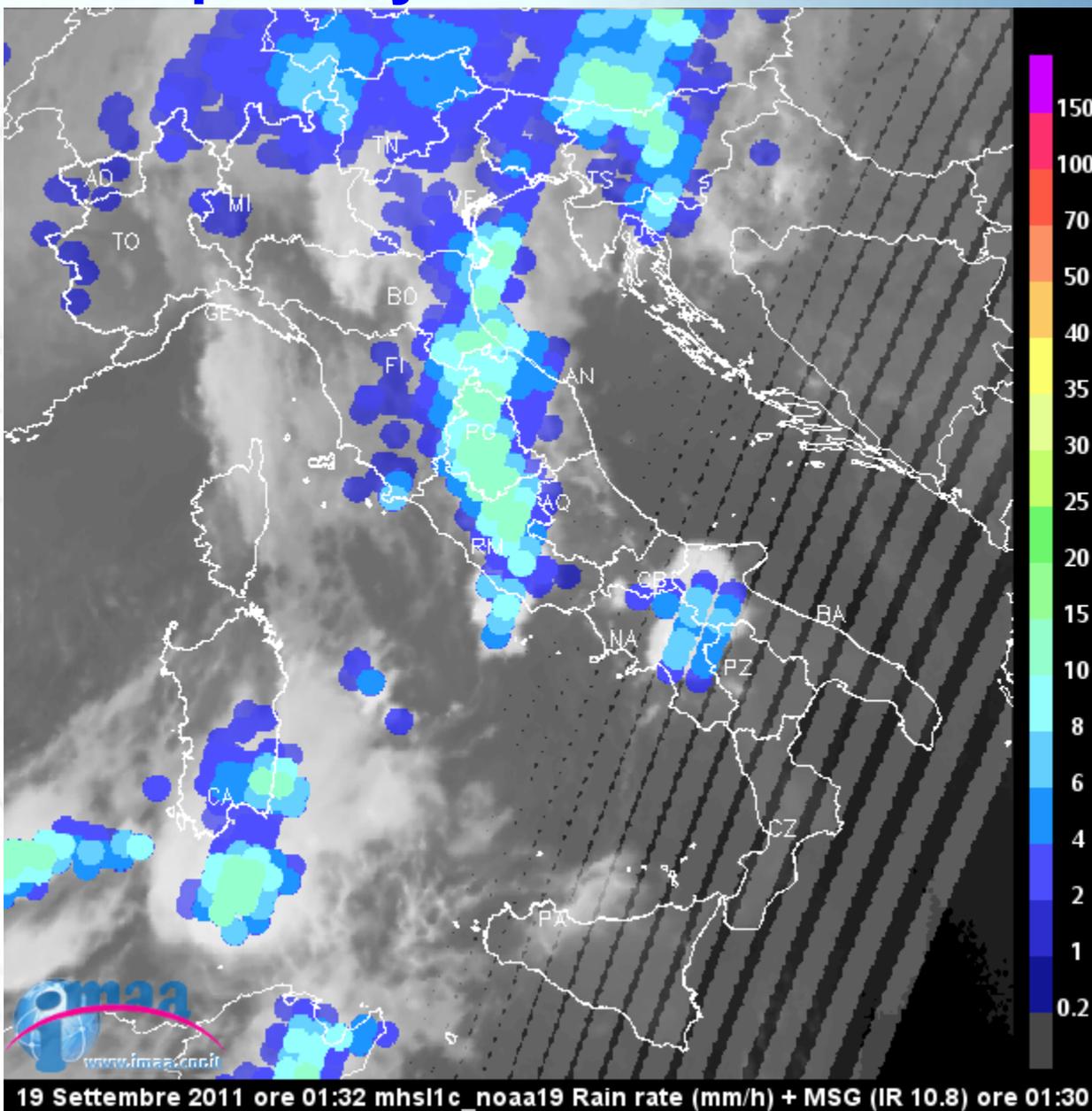
What's up today? 2011/09/18 11:27



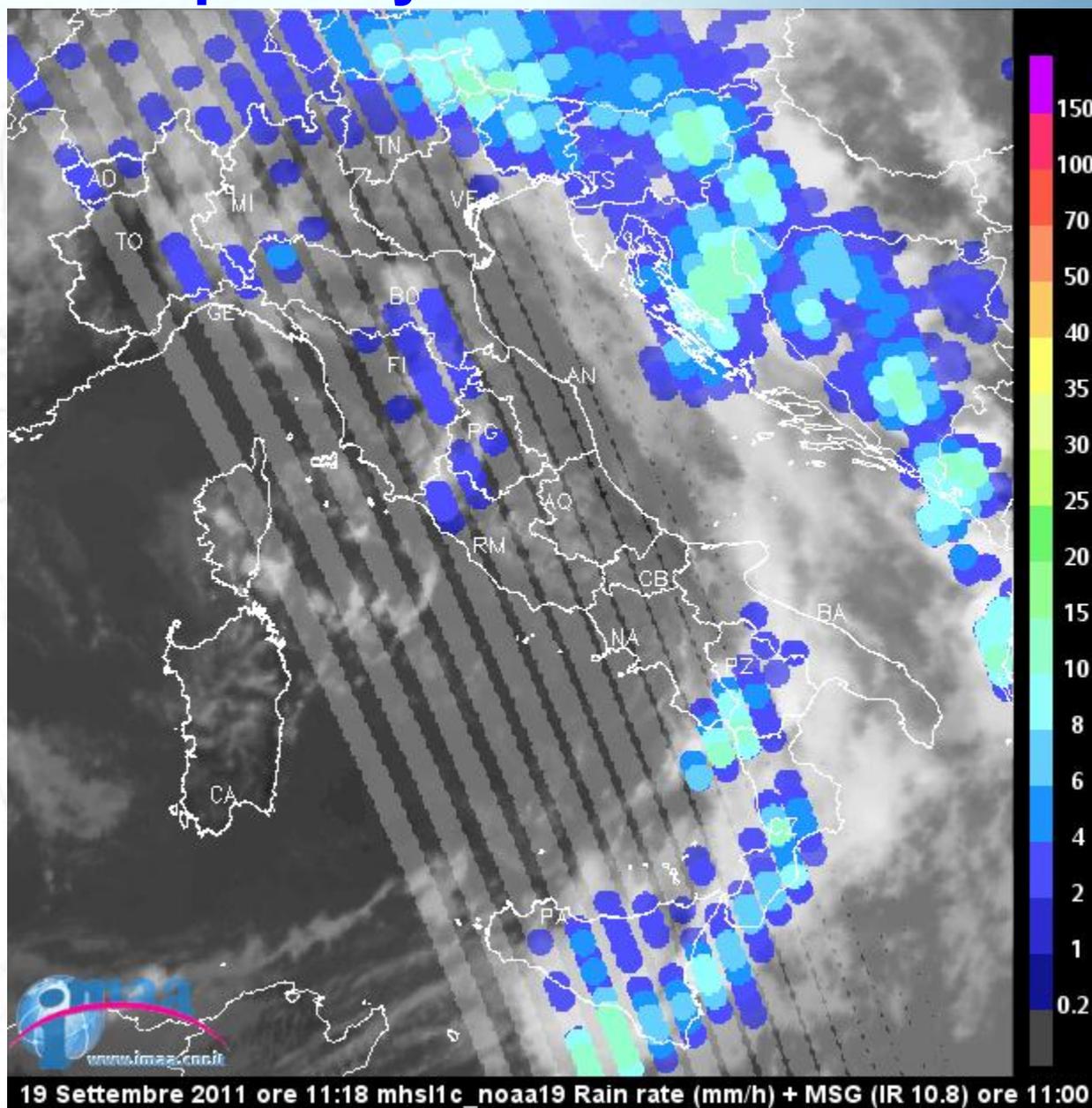
What's up today? 2011/09/19 03:07



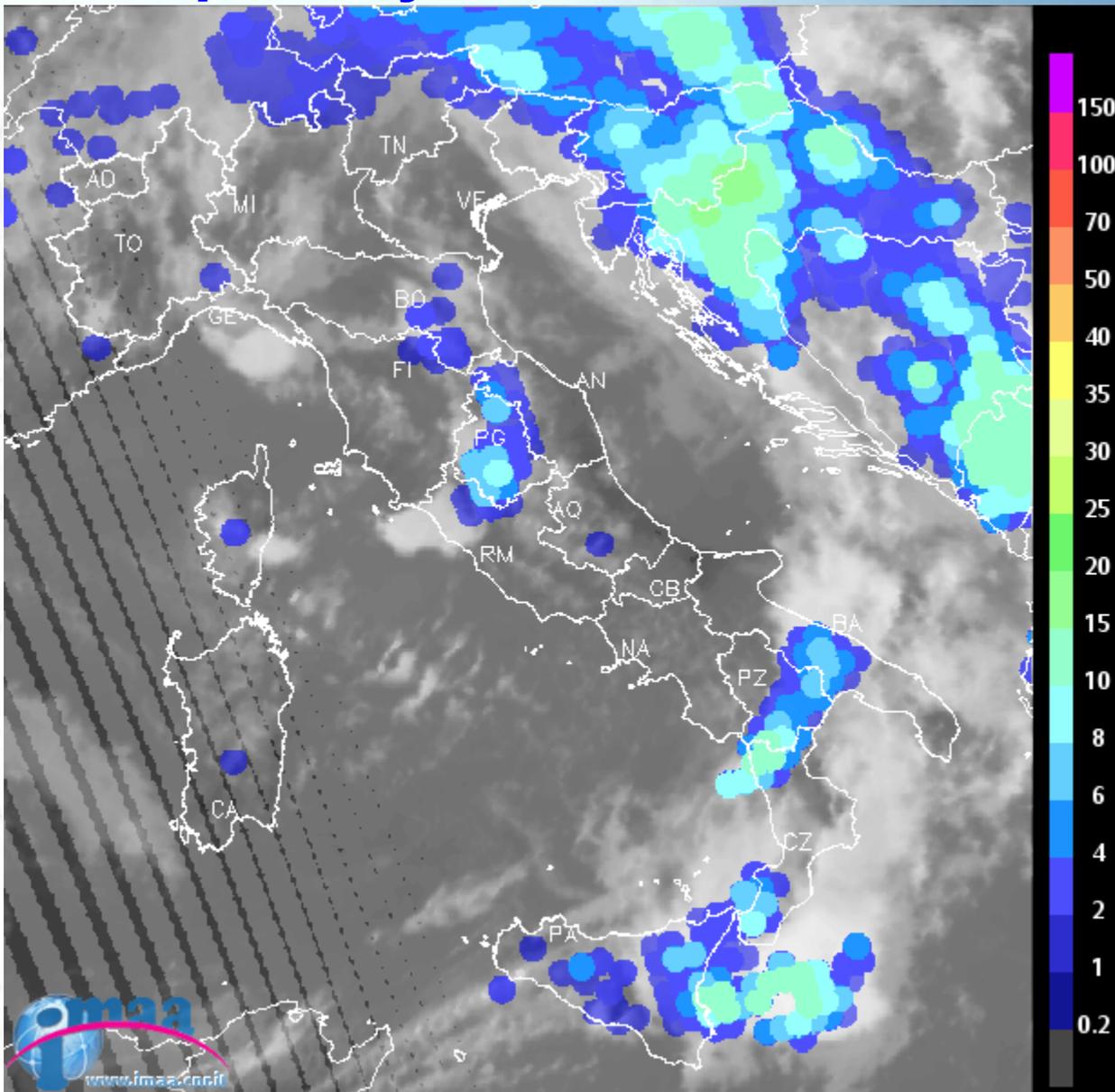
What's up today? 2011/09/19 03:32



What's up today? 2011/09/19 13:18



What's up today? 2011/09/19 14:34



Validation of PEMW with RNC

- How can we validate our product?!??
 - Very difficult – no reference “truth” available
 - Raingauge – point measurement
 - Radar – volume measurement, but radial wrt to radar location

- Previously PEMW has been validated with selected case studies (using raingauges and France-UK radar network)



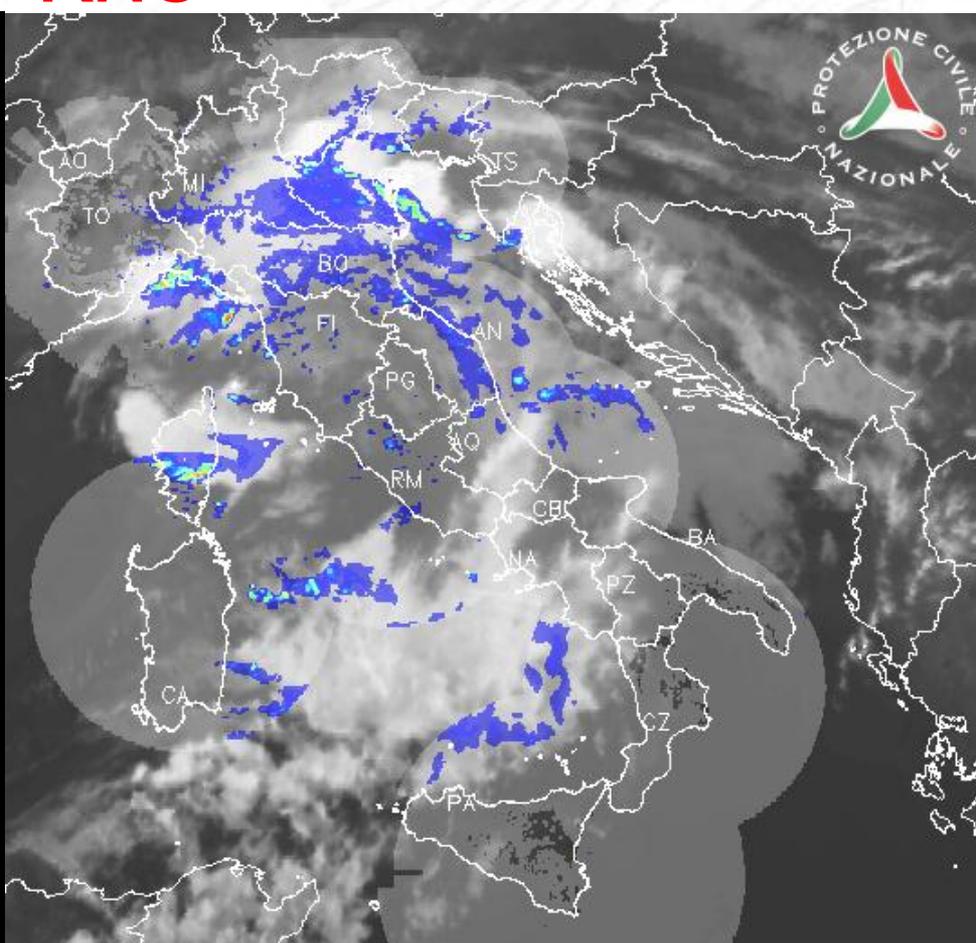
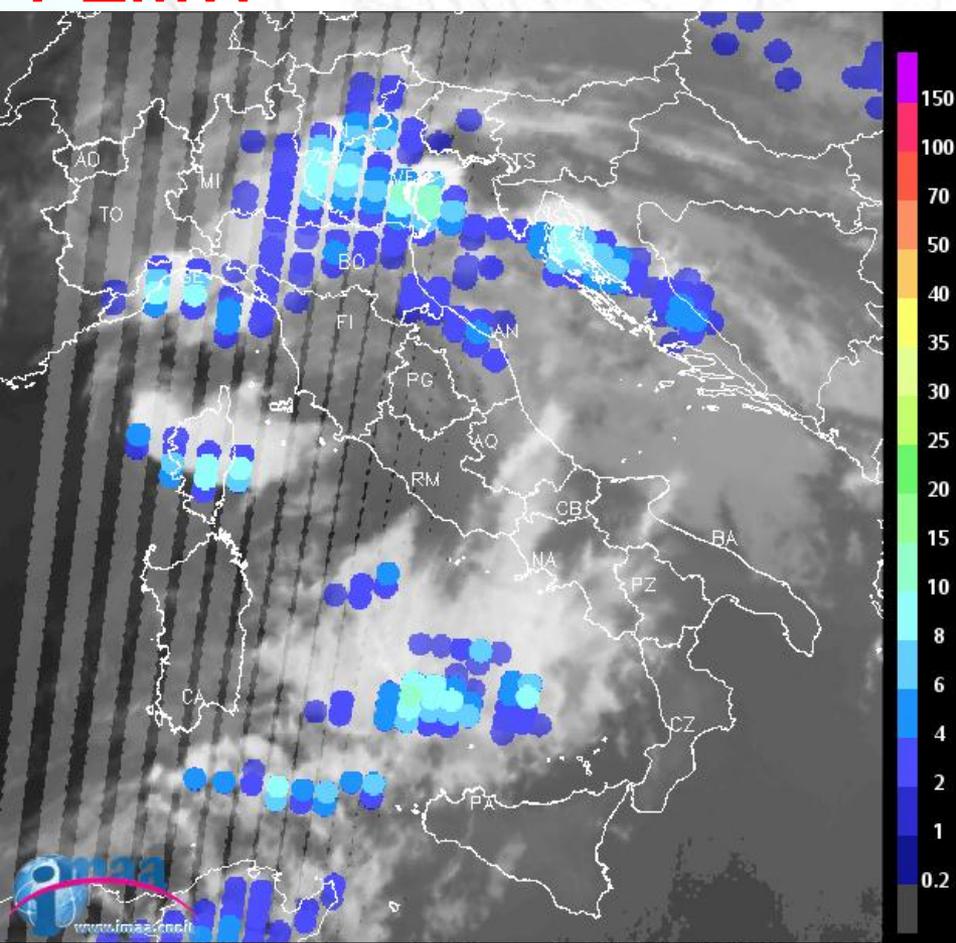
Case Studies

2011/07/05 01:22 UTC

Note: Storm in North East and southern Tyrrhenian

PEMW

RNC



Validation of PEMW with RNC

- Radar Network Composite allows a systematic validation of operational PEMW over Italy
 - Temporal colocation: RNC data within ± 7.5 min of sat overpass
 - Spatial colocation: RNC data convoluted within PEMW FOVs

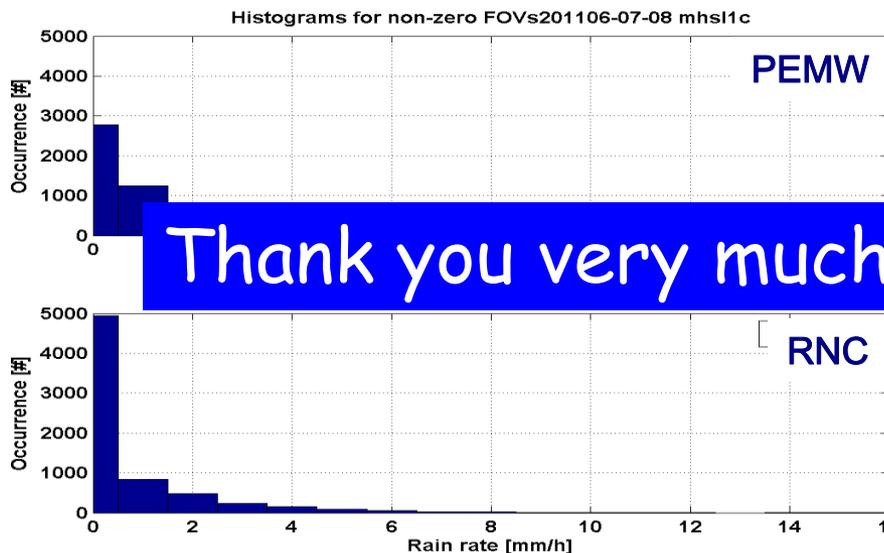
- Quantitative scores are computed to investigate the consistence between these two source of rainfall information
 - Dichotomous assessment
 - Contingency table, Accuracy, bias score, POD, FAR, HSS, PODN, CSI, HK, ISE, FOM, FOH, POFD, DFR, FOCN, TSS,...
 - Continuous assessment
 - Bias, rmse, FSE, FMR, FVR, NEB,...



Validation of PEMW with RNC

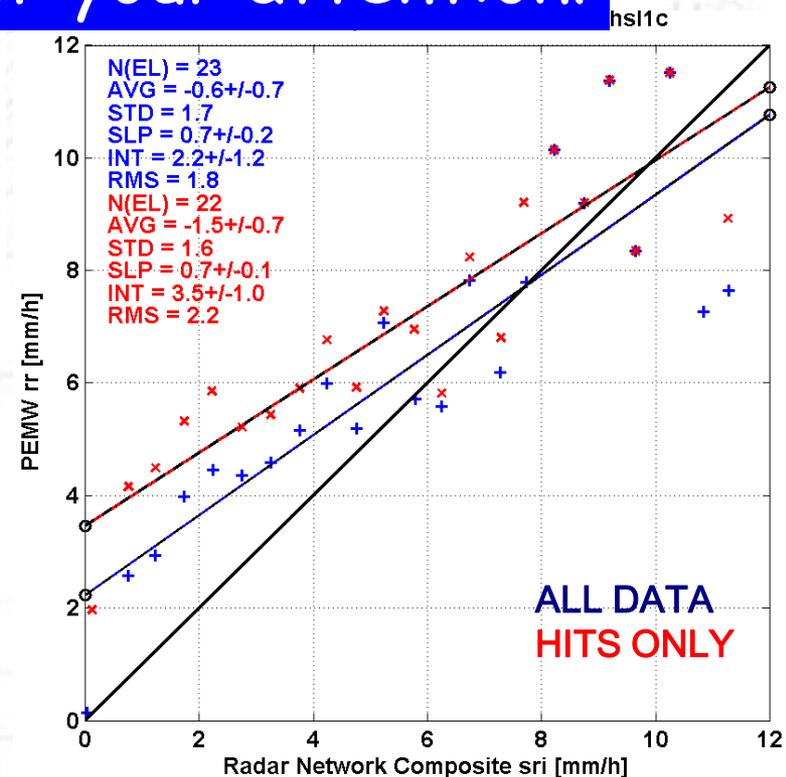
Seasonal product

3-month colocated PEMW and RNC data set
 Example: Summer 2011 (J-J-A)



	RNC (Y)	RNC (N)	
PEMW (Y)	H = 1797 H = 1348	F = 4024 F = 1738	5821 3086
PEMW (N)	M = 5100 M = 619	CN = 353173 CN = 360389	358273 361008
	6897	357197	364094 364094

Thank you very much for your attention!



	All FOVs	Neg. ≤ 0.5 mm/h	RNC $\geq 1 / 5$ mm/h
Accuracy	0.97	0.99	-
Bias	0.84	1.57	-
POD	0.26	0.68	0.73 / 0.91
FAR	0.69	0.56	-
HSS	0.27	0.53	-



Publications

- Ricciardelli E., F. Romano, D. Cimini, F. S. Marzano, V. Cuomo, A statistical approach for rainfall confidence estimation using MSG-SEVIRI observations, EUMETSAT Meteorological Satellite Conference, Cordoba 20-24 September 2010
- Di Tomaso E., F. Romano, and V. Cuomo, Rainfall estimation from satellite passive microwave observations in the range 89 GHz to 190 GHz, J. of Geophysical Res., VOL. 114, 2009.
- Ricciardelli E., F. Romano, V. Cuomo, Physical and statistical approaches for cloud identification using MSG-SEVIRI data, Remote Sensing of Environment, 112 (2741-2760), 2008.
- Romano F., D. Cimini, R. Rizzi, V. Cuomo, Multilayered cloud parameters retrievals from combined infrared and microwave satellite observations, Journal of Geophysical Research, 112, D08210, doi:10.1029/2006JD007745, 2007.
- Marzano, F. S., D. Cimini, and F. J. Turk, Multivariate Probability Matching for Microwave Infrared Combined Rainfall Algorithm (MICRA), Measuring Precipitation from Space, Levizzani V., P. Bauer, and F. J. Turk Editors, Springer, 2007.
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- Marzano, F. S., M. Palmacci, D. Cimini, G. Giuliani and J. F. Turk: Multivariate Statistical Integration of Satellite Infrared and Microwave Radiometric Measurements for Rainfall Retrieval at the Geostationary Scale, IEEE Transactions on Geoscience and Remote Sensing, Vol. 42, n. 5, pp. 1018-1032, 2004.

