

A climatology of thermodynamic vs. dynamic Arctic wintertime sea ice thickness effects during the CryoSat-2 era

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CMIP6 Models (1990-2009 monthly mean)



Keen et al., 2021

Sea ice thickness growth (simplified) "governing" equation

$$\frac{\partial H}{\partial t} = f(t, H, \mathbf{x}) \xrightarrow{\text{Dynamics}} -\nabla \cdot (\mathbf{u}H)$$

H is ice thickness, t is time, **x** is location, **u** is drift velocity

Thermodynamics: ice thickness change through phase change

Dynamics: ice thickness change through relative motion of ice parcels (compressible advection)

Sea ice thickness growth (simplified) "governing" equation

$$\frac{\partial H}{\partial t} = f(t, H, \mathbf{x}) \stackrel{\text{Dynamics}}{-\nabla \cdot (\mathbf{u}H)}$$

- Affected by differing mechanisms in a changing climate
- Large scale observations of each process and their comparison to climate models are needed but lacking

Science Question:

What are relative effects of dynamic and thermodynamic sea ice thickness processes in the Arctic?

Stefan's Law: simple sea ice thermodynamics



 ρ =ice density, L=latent heat, H=ice depth, k=ice conductivity, T_f=ice freezing point, T_{si}= snow--ice temperature, F_w = ocean heat flux to ice



Allow instantaneous relationship!

Neglect:

- Horizontal conduction
- Thermal inertia
- Internal heat sources

Stefan's Law Integrated Conducted Energy (SLICE)

Passive microwave retrieved snow ice interface temp. (Kilic et al., 2019)



Stefan's Law solution

$$H = \sqrt{H_0^2 + \delta t} \frac{2\kappa_{\text{eff}}}{\rho_{\text{i}}L} (T_{\text{f}} - T_{\text{si}}) - \delta t \frac{F_{\text{w}}}{\rho_{\text{i}}L}$$

Attributes

- Retrieves instantaneous thermodynamic growth
- Daily, basin-wide coverage
- $F_w = 2 w/m^2$

Caveats

- Requires initial condition (H₀) to retrieve absolute thickness
- No melt or sunlight! Growth season only
- >95% ice concentration only

Use CryoSat-2 and SLICE to estimate weekly dynamics

$$\frac{\partial H}{\partial t} = f(t, H, \mathbf{x}) \xrightarrow{\text{Dynamics}} \nabla \cdot (\mathbf{u}H)$$

$$\frac{\partial [CryoSat-2]}{\partial t} = [SLICE] + [residual]$$

CryoSat-2: Alfred Wegener Institute (AWI) CS2SMOS

- weekly CS2/SMOS combination sea ice thickness product
- Covers 2010-2021
- Ricker et al. (2017)

Decompose dynamics into advection and deformation

$$\frac{\partial H}{\partial t} = f(t, H, \mathbf{x}) - (\nabla H) \cdot \mathbf{u} - H(\nabla \cdot \mathbf{u})$$
Advection Deformation
$$\frac{\partial [CryoSat-2]}{\partial t} = [SLICE] - [(\nabla CryoSat - 2) \cdot \mathbf{mot.vec.}] - [residual]$$
Deformation
(Lagrangian dynamics)

Motion vectors

- Approximate advection term with CS2 thickness field and motion vector
- Use satellite-based motion vector product (Tschudi et al., 2019)

2010-2021 wintertime mean



- Thermodynamic growth highest where ice is thinnest
- Dynamics increases thickness north of Canadian Archipelago (CAA) and eastern Siberia

2010-2021 wintertime mean



- BG: advects thick ice towards Alaskan coast where divergence/lead formation reduces thickness, ice is deposited/ridges north of eastern Siberia
- TD: advects thinner ice towards CAA, ridging along drift increases thickness

Dynamics relative to thermodynamics



Some regions are dominated by dynamics, with dynamics more than doubling thermodynamics

Dynamics are -30% of thermodynamics in basin-wide mean



Close agreement with Keen et al., 2021 and Ricker et al., 2021

Conclusions

- Use SLICE and CryoSat-2 to estimate thermo. and dyn. effects
- Dyn. are -30% of thermo. in basin-wide mean, agrees with literature
- Basin-wide, 25km Eulerian, sub-seasonal temporal resolution, and long-term observations expands on literature
- BG: negative deformation/leading in westward leg and positive deformation/ridging eastern Siberian coast and eastward leg
- TD: negative advection coupled with positive deformation/ridging
- Dynamics and deformation effects account for 2x thermodynamic growth in some regions

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SLICE compares well with buoys



Uncertainty

