

# Studies of all-sky radiance assimilation at NCEP

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# Outline

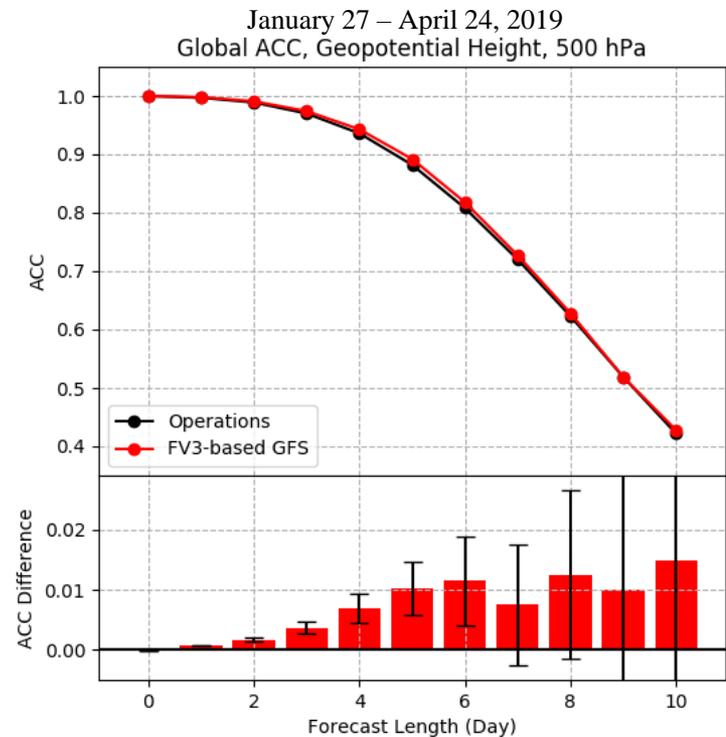
- Status of all-sky radiance assimilation in the operational FV3GFS
- Incorporation of convective clouds in the all-sky radiance assimilation
- All-sky radiance assimilation over land
- Ongoing work and future plan

# FV3GFS hybrid 4DEnVar system became operational in June 2019

$$J(\mathbf{x}'_c, \mathbf{a}) = b_c \frac{1}{2} (\mathbf{x}'_c)^T \mathbf{B}_c^{-1} (\mathbf{x}'_c) + b_e \frac{1}{2} \mathbf{a}^T \mathbf{L}^{-1} \mathbf{a} + \frac{1}{2} \sum_{k=1}^K (\mathbf{H}_k \mathbf{x}'_{(t)k} - \mathbf{y}'_k)^T \mathbf{R}_k^{-1} (\mathbf{H}_k \mathbf{x}'_{(t)k} - \mathbf{y}'_k)$$

$$\mathbf{z} = \mathbf{B}^{-1} \mathbf{x}'_c \quad \mathbf{v} = \mathbf{L}^{-1} \mathbf{a}$$

- FV3 dynamic core, cubed-sphere grid, non-hydrostatic option
- Initial prototyping with (mostly) GFS physics (new: GFDL microphysics)
- Stochastic physics: SPPT + SHUM only
- C768 (~13km) L64 (55km top)
- Data assimilation: adaptation of original operational GSI hybrid 4DEnVar scheme with 80 ensemble members
- Ensemble and increment resolution have been increased from ~39km to ~25 km



(Kleist and Thomas, 2019)

# Status of radiance data assimilation in the FV3GFS

## Microwave:

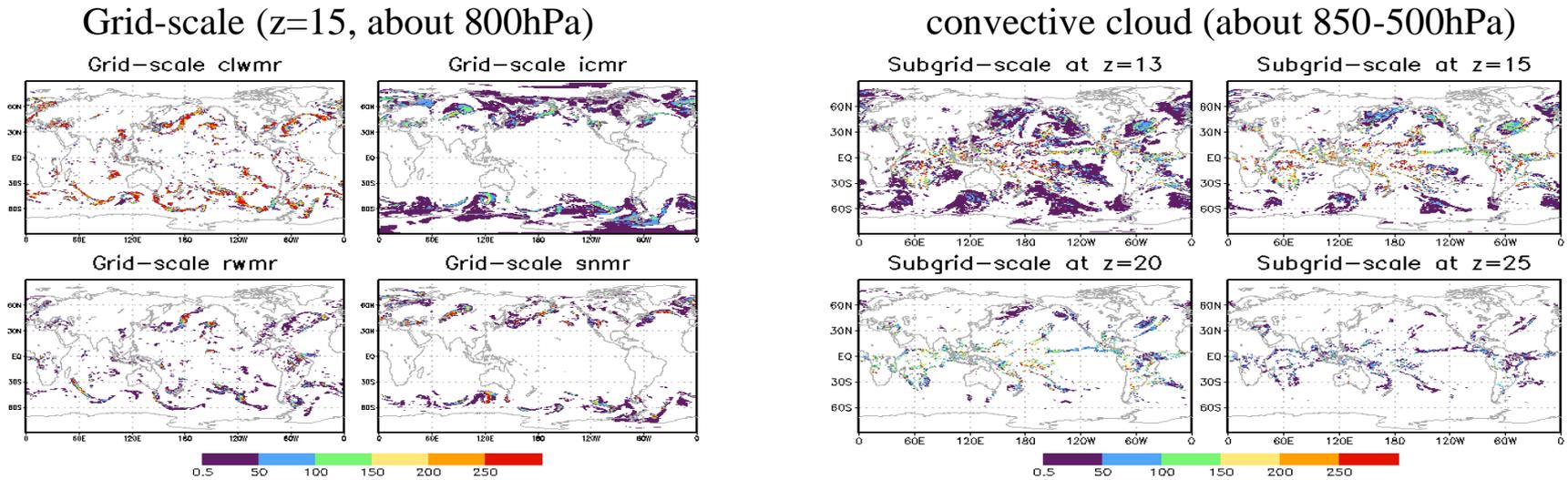
- **AMSU-A**: NOAA-15, 18, 19, MetOp-A, MetOp-B, Aqua
- **ATMS**: NPP, NOAA-20
- MHS: NOAA-18, 19, MetOp-A, MetOp-B
- SSMIS: DSMP-F17
- SAPHIRE: Megha-Tropique

## Infrared:

- AIRS: Aqua
- GOES-15 Sounder
- IASI: MetOp-A, MetOp-B
- CrIS: NPP, NOAA-20
- SEVIRI: MeteoSat-11
- AVHRR: MetOp-A, NOAA-18

- ❑ Over ocean, both clear-sky and cloudy radiances from AMSU-A and ATMS over ocean FOVs are assimilated in the **all-sky** approach (Zhu et al. 2016; Zhu et al. 2019), only clear-sky radiances are assimilated from other sensors
- ❑ Over land, only clear-sky radiances are assimilated for all the sensors

In the cloudy radiance assimilation in the GFS, only grid-scale clouds have been used in the radiance simulation calculation. The lack of clouds in radiance simulation in tropics is clear.

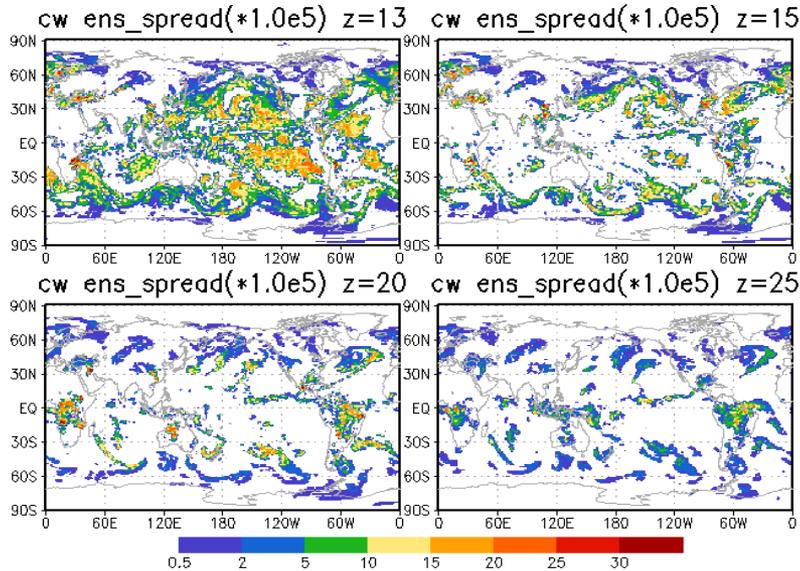


**Convective cloud was added as an optional model output field later 2018.**

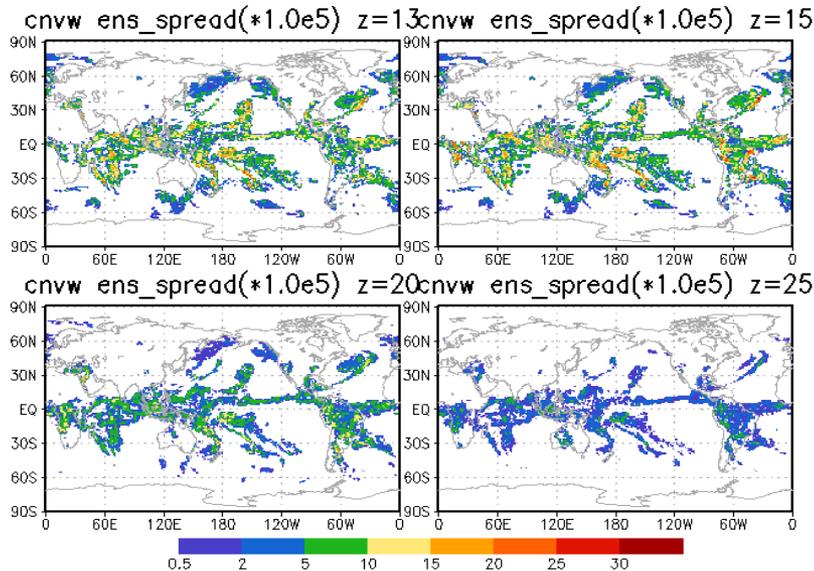
To preserve forecast model water budget, convective cloud increments should not be fed back to the forecast model. Considered approaches to incorporating convective clouds:

- Combine subgrid & grid scale clouds in the GSI, remove subgrid-scale clouds from cloud analyses before passing them back to model, or do not feed back cloud to model;
- Treat convective clouds separately as additional control variable(s).

# Ensemble spread: grid-scale

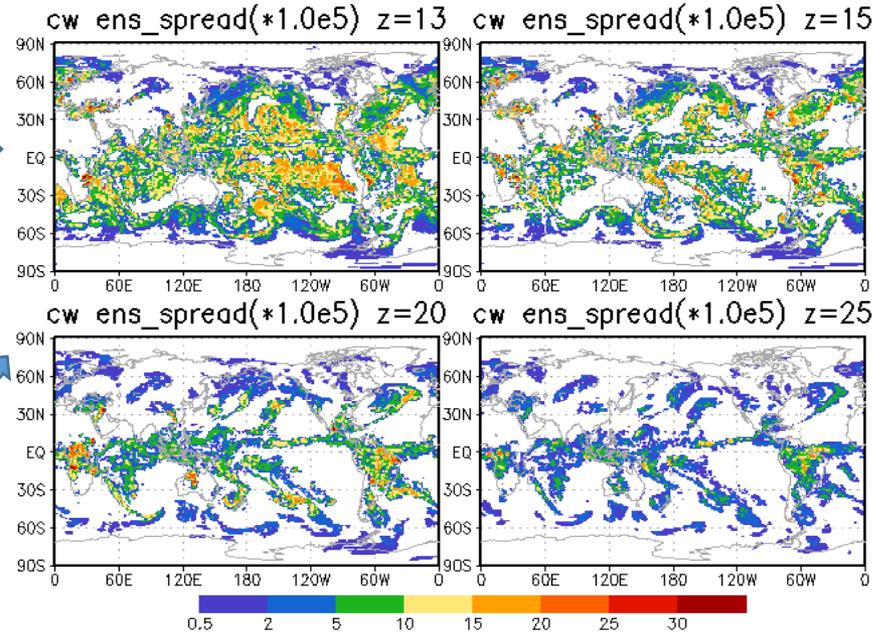


subgrid-scale



Different characteristics were observed for grid-scale and convective clouds

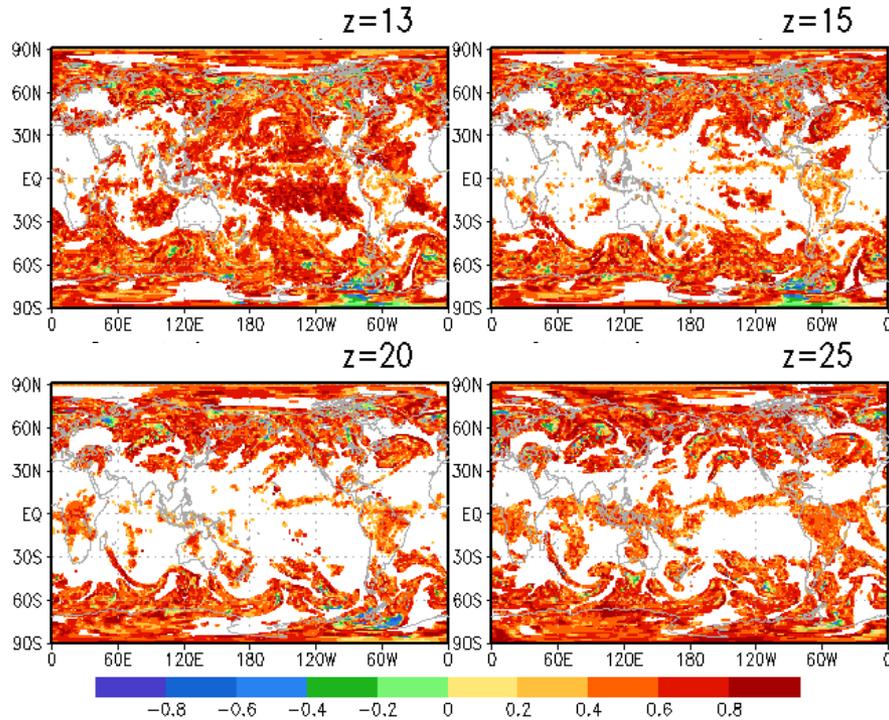
## Combined ensemble spread



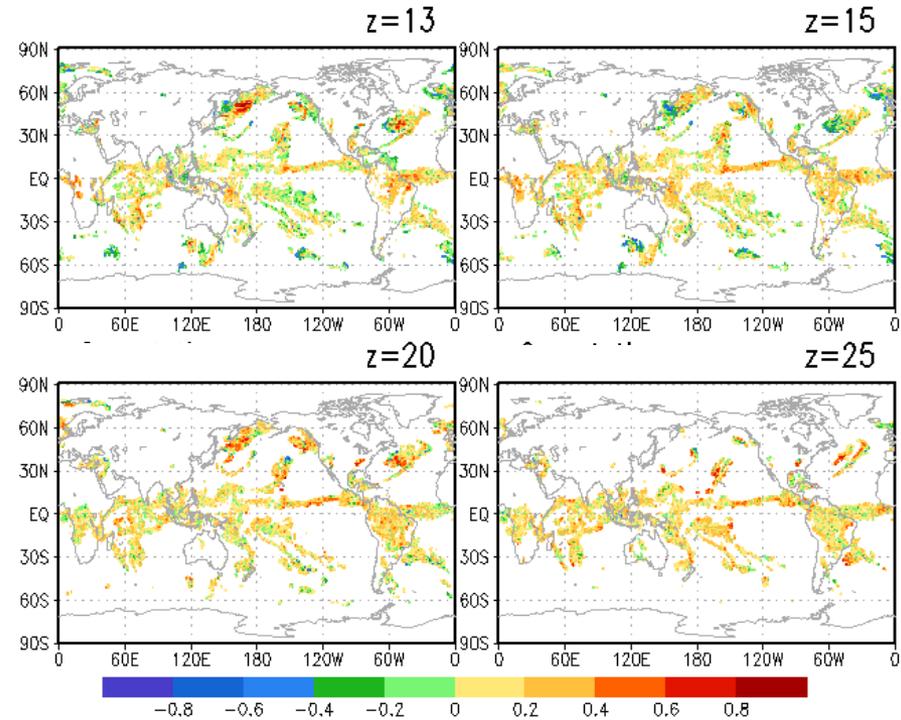
- In the tropics, the contribution from convective clouds is dominant (z14-25)
- The contribution from convective clouds corresponds clearly with the ITCZ and SPCZ

# Point-wise correlation between cloud water and RH

grid-scale (about 850-500hPa)



convective (about 850-500hPa)

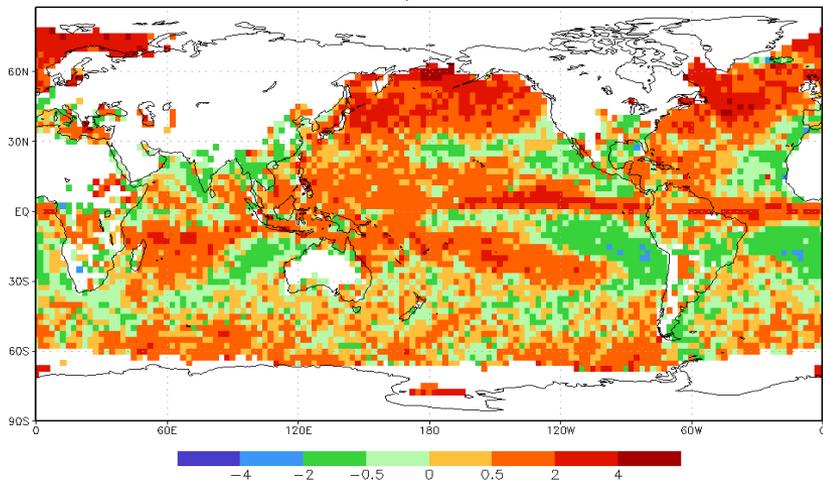


- **High** point-wise correlation between grids-scale clouds and RH
- **Low** correlation between convective clouds and RH
- Because of its large ensemble spread and low correlation with T and RH, we didn't take the approach of using convective clouds as a separate cloud control variable as there are no direct cloud observations to constrain the field

# Impact of incorporating convective clouds

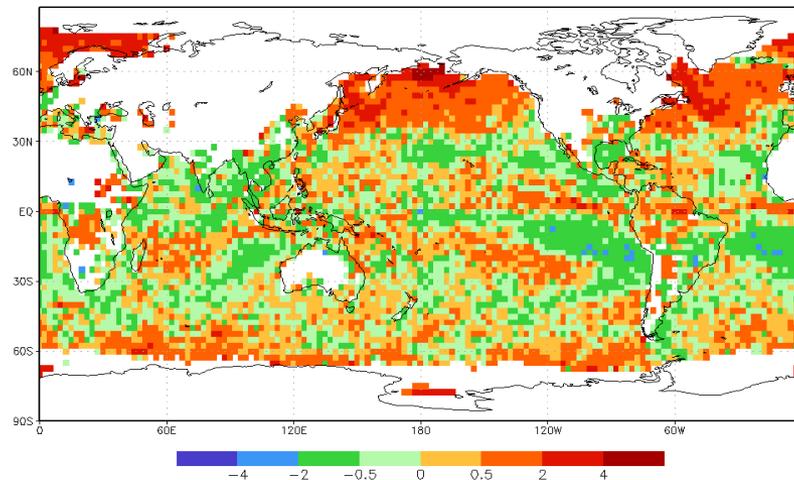
## w/o convective clouds

ATMS NPP Ch2 w/o convective clouds

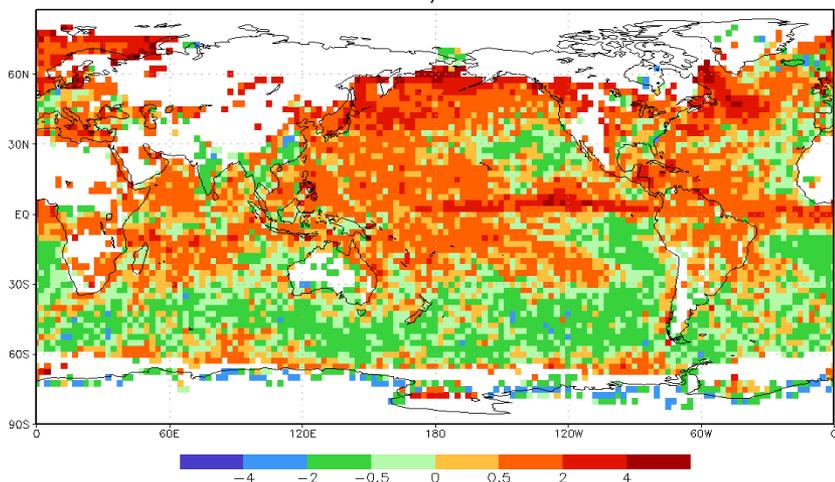


## with convective clouds

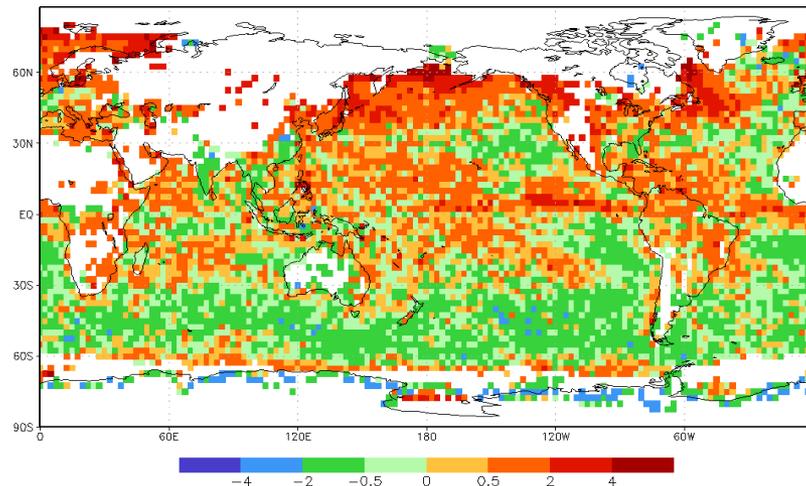
ATMS NPP Ch2 with convective clouds



AMSU-A N19 Ch1 w/o convective clouds



AMSU-A N19 Ch1 with convective clouds



- Improve OmF with the much needed clouds in the tropics

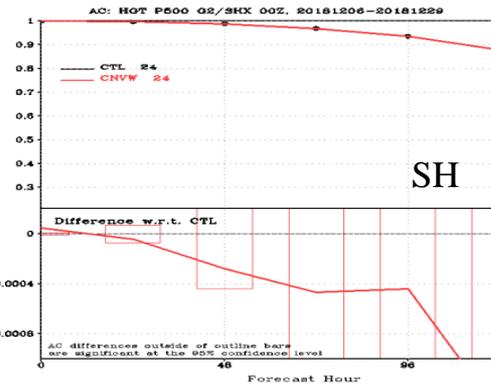
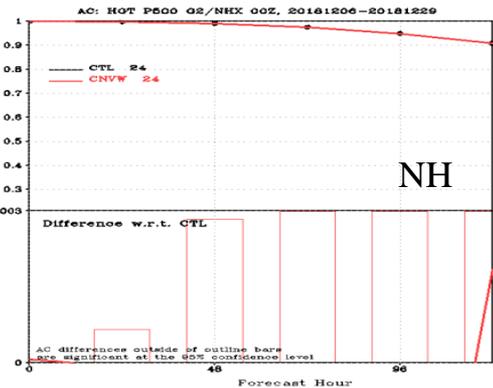
### 3D EnVar experiments were performed

- Improve OmF in the tropics;
- Provide realistic cloud ensemble spread in tropics;
- Improve model forecast spinup with reduced RMSE of wind & temperature;
- Prepare for using features of CRTM 2.3.0 (e.g. cloud fraction acts as reducing cloud for these MW data)
- However, forecast anomaly correlation is slightly degraded in Southern Hemisphere

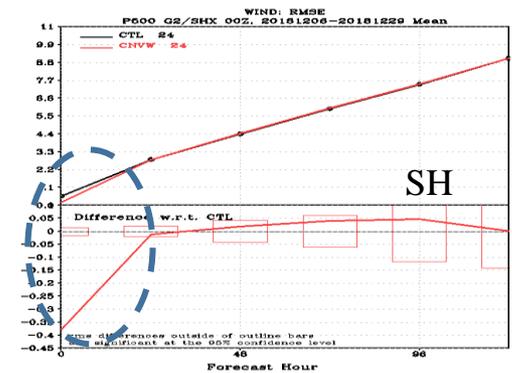
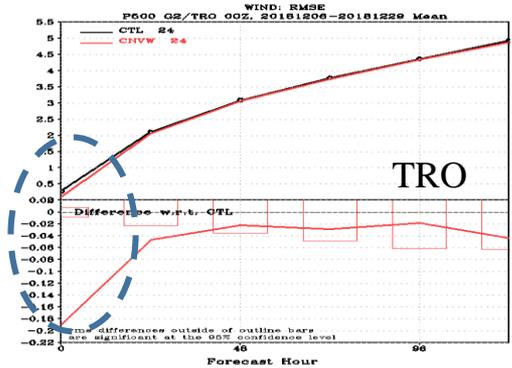
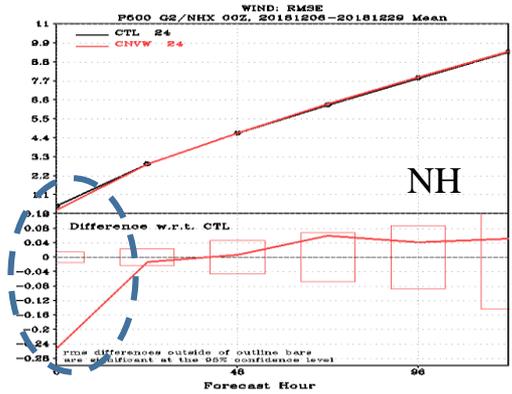
### Work plan:

- Investigate the impact of applying smoothing to convective cloud ensemble as smoothing has been applied to other fields

### Geopotential height anomaly correlation at 500hPa



### Wind RMSE at 500 hPa

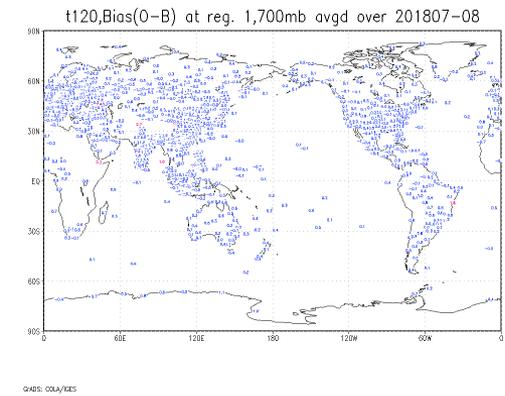
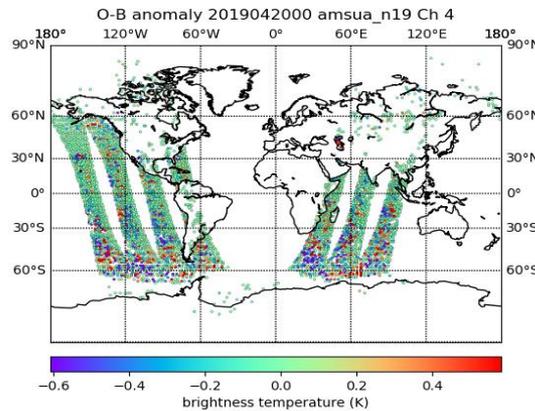
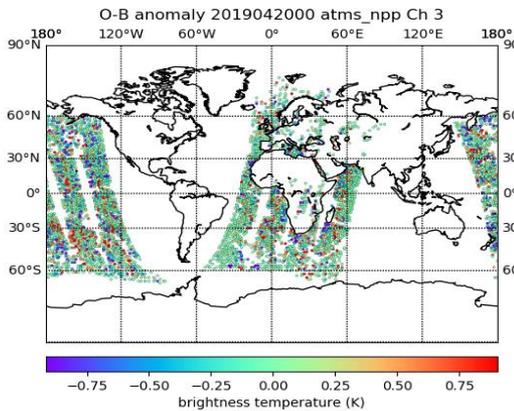


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# Challenges in assimilating surface-sensitive radiances over land

- ❑ Far fewer radiances are used over land than over ocean, only clear-sky radiances are used
- ❑ Radiance bias correction (Derber and Wu 1998; Zhu et al. 2014) is anchored on conventional data, bias correction estimate is currently dominated by radiances over ocean
- ❑ Challenges in assimilating surface-sensitive radiances over land
  - Uncertainties of NESDIS microwave land physical emissivity model in the CRTM
    - Radiance simulation is sensitive to emissivity accuracy
    - Emissivity sensitivity is used in bias correction & quality control, and it is required if emissivity is a control variable
  - Land surface model component in the forecast model: Uncertainties of land surface state properties, e.g. land surface skin temperature (LST) and soil moisture
  - Problematic cloud detection scheme over land where obs are compared with the equivalent clear-sky TBs



Radiosonde observation

# Instantaneous emissivity retrieval

(Prigent et al 2006; Karbou et al 2005; Baordo and Geer 2016; etc)

For a scattering-free atmosphere, assuming a flat and specular surface, observed brightness temperature BT can be expressed as:

$$BT_{obs} = \varepsilon T_s \Gamma + BT_{up} + \Gamma (1 - \varepsilon) BT_{down}$$

- $BT_{up}$  atmospheric upwelling radiation
- $BT_{down}$  atmospheric downwelling radiation
- $\Gamma$  atmospheric surface-to-space transmittance
- $T_s$  surface skin temperature (effective radiating temperature of the surface at the relevant frequency)

Surface emissivity can be calculated as:

$$\varepsilon = \frac{BT_{obs} - BT_{up} - BT_{down} \Gamma}{(T_s - BT_{down}) \Gamma}$$

Or with effective cloud fraction C

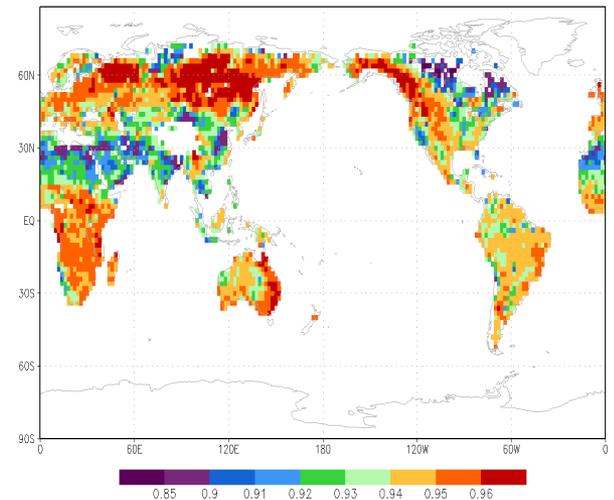
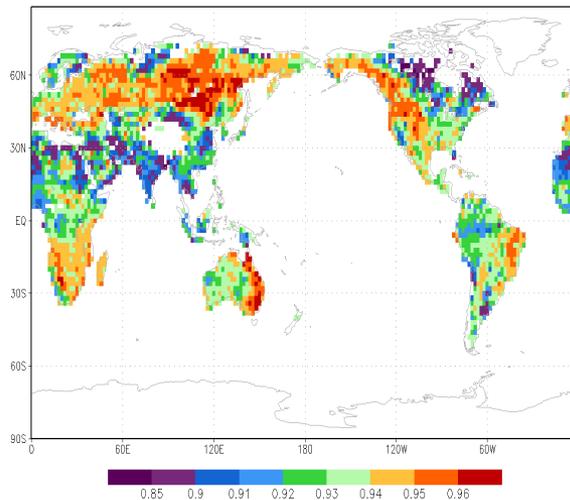
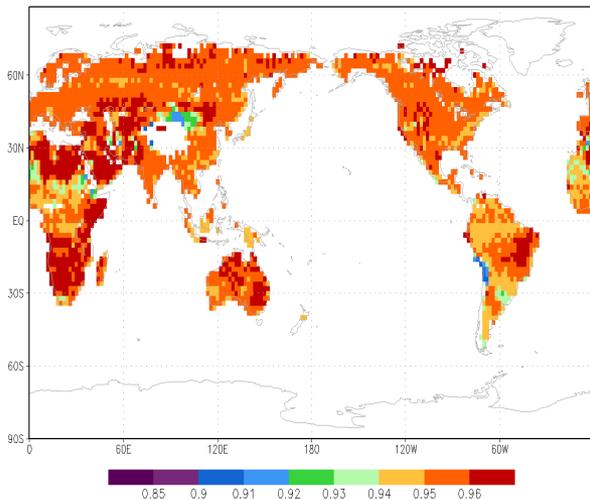
$$\varepsilon = \frac{BT_{obs} - (1-C) (BT_{up}^{clr} + BT_{down}^{clr} \Gamma^{clr}) - C (BT_{up}^{cld} + BT_{down}^{cld} \Gamma^{cld})}{(1-C) (T_s - BT_{down}^{clr}) \Gamma^{clr} + C (T_s - BT_{down}^{cld}) \Gamma^{cld}}$$

# Comparison of a-warm-month averaged emissivity

CRTM emissivity

instantaneous emissivity from GFS

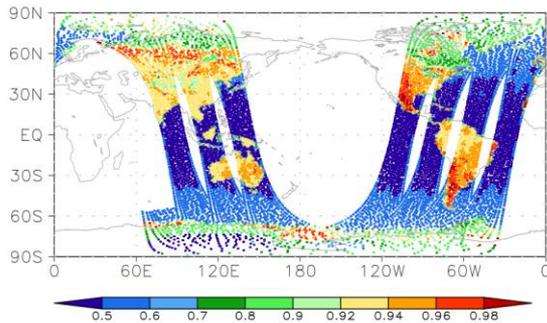
climatology (TELSEM2)



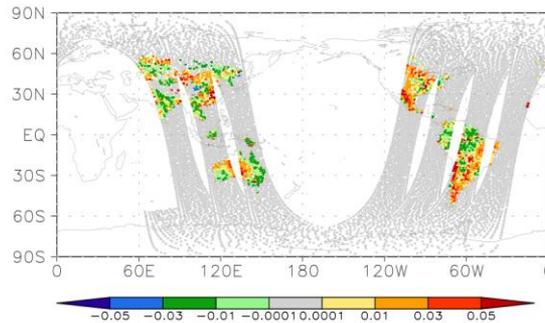
- The patterns of the instantaneous emissivity from the GFS data assimilation system are similar to the climatology from TELSEM2.
- Large differences are observed in some areas between the CRTM emissivity and the climatology.
- Warm and cold months of instantaneous emissivities from the GFS have been provided to CSEM developers for the machine learning study.

# Issues with the emissivity sensitivity in the CRTM

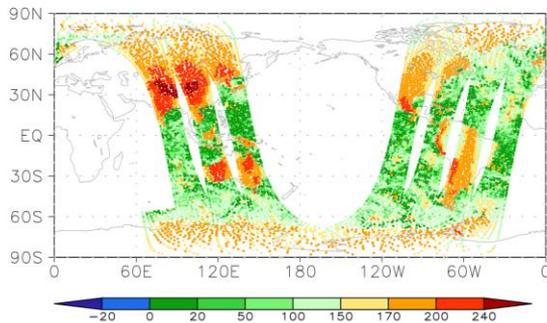
CRTM emissivity



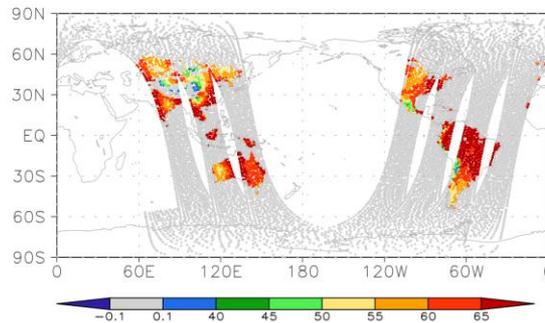
$e_{\text{CRTM}} - e_{\text{(analytical+TELSEM)}}$



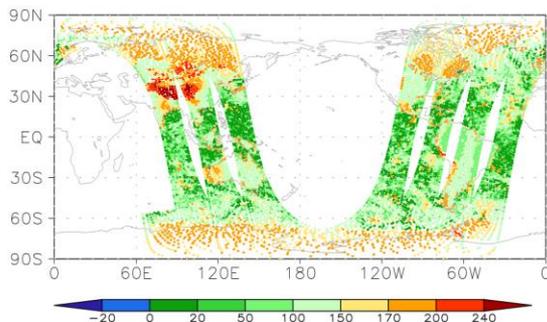
CRTM emissivity sensitivity ( $e_k$ )



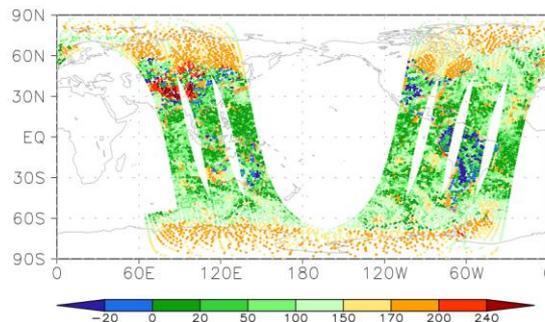
Diff  $e_k$  with user\_emissivity (clr-sky)



$e_k$  with user\_emissivity (clr-sky)



$e_k$  with user emissivity (cld-sky)



When user supplied emissivity is used in the CRTM, emissivity sensitivity calculated is problematic.

- seems too low in clear sky case
- negative emissivity sensitivity in cloudy case, which is not physically meaningful.

The calculation is based on

- the simple Kirchhoff 's law over specular surface.
- the assumption that emissivity is an independent variable, which actually depends on surface temperature strongly.

# Ongoing work and future plan

- **Intermediate goal:** using instantaneous emissivity retrieval from the GSI combined with TELSEM atlas, assimilate all-sky microwave radiances over land
  - Help Community Surface Emissivity Model (CSEM) developers to improve CSEM
  
- **Long-term goal:** with the improved CSEM, soil moisture and LST analyses using radiances from low-frequency (e.g. L-band) microwave satellite sensors, such as AMSR2, SMOS, GMI
  - Near-surface temperature and humidity observations (currently not assimilated in FV3GFS) are strongly influenced by soil moisture in appropriate conditions, and will also help to constrain radiance assimilation
  - Radiance observations involve variables from more than one components: both atmosphere and land data assimilations. Coupled data assimilation (Kleist 2019, personal communication)