

# MODIS and VIIRS Data Environmental Applications: Part 2

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Hawaii Direct Broadcast Polar Orbiter Workshop  
University of Hawaii Manoa  
21 August 2013



# Aviation Applications

Turbulence, Clouds



# Atmospheric Turbulence

## What is Turbulence?



This smoke pattern shows turbulence as rapid, abrupt and chaotic changes in the speed and direction of air flow.

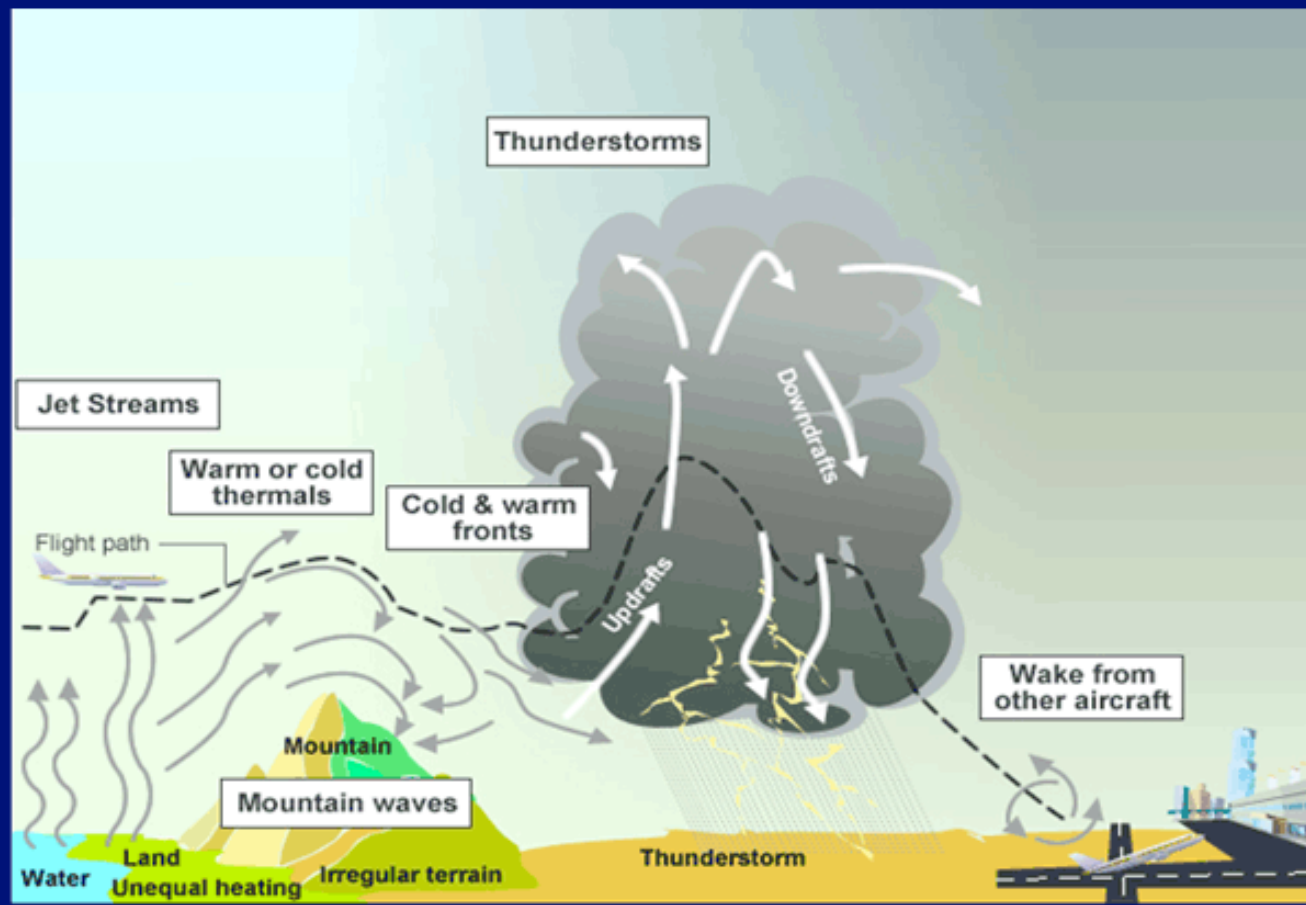


Colored smoke is used to show "wake turbulence" generated by an aircraft upon take-off or landing.



# Atmospheric Turbulence

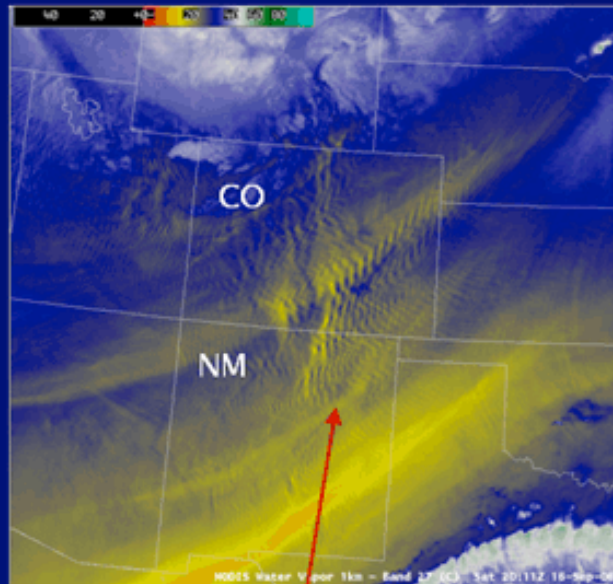
## Causes of Turbulence



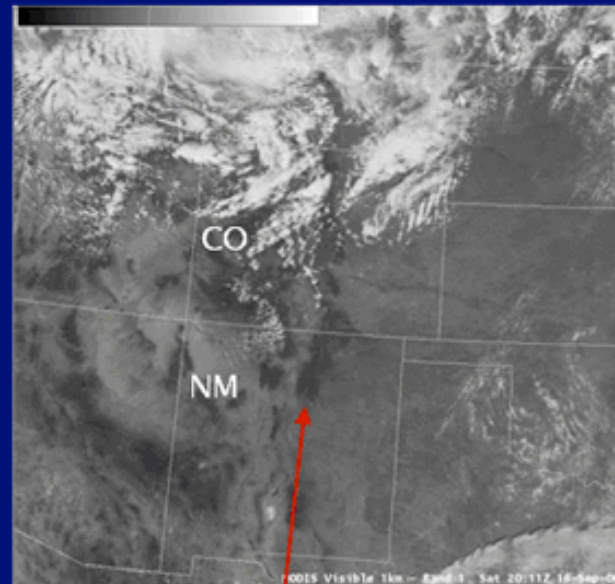


# Why is $6.7\text{ }\mu\text{m}$ Important for the Detection of Turbulence?

## Water Vapor Channel & Visible Channel



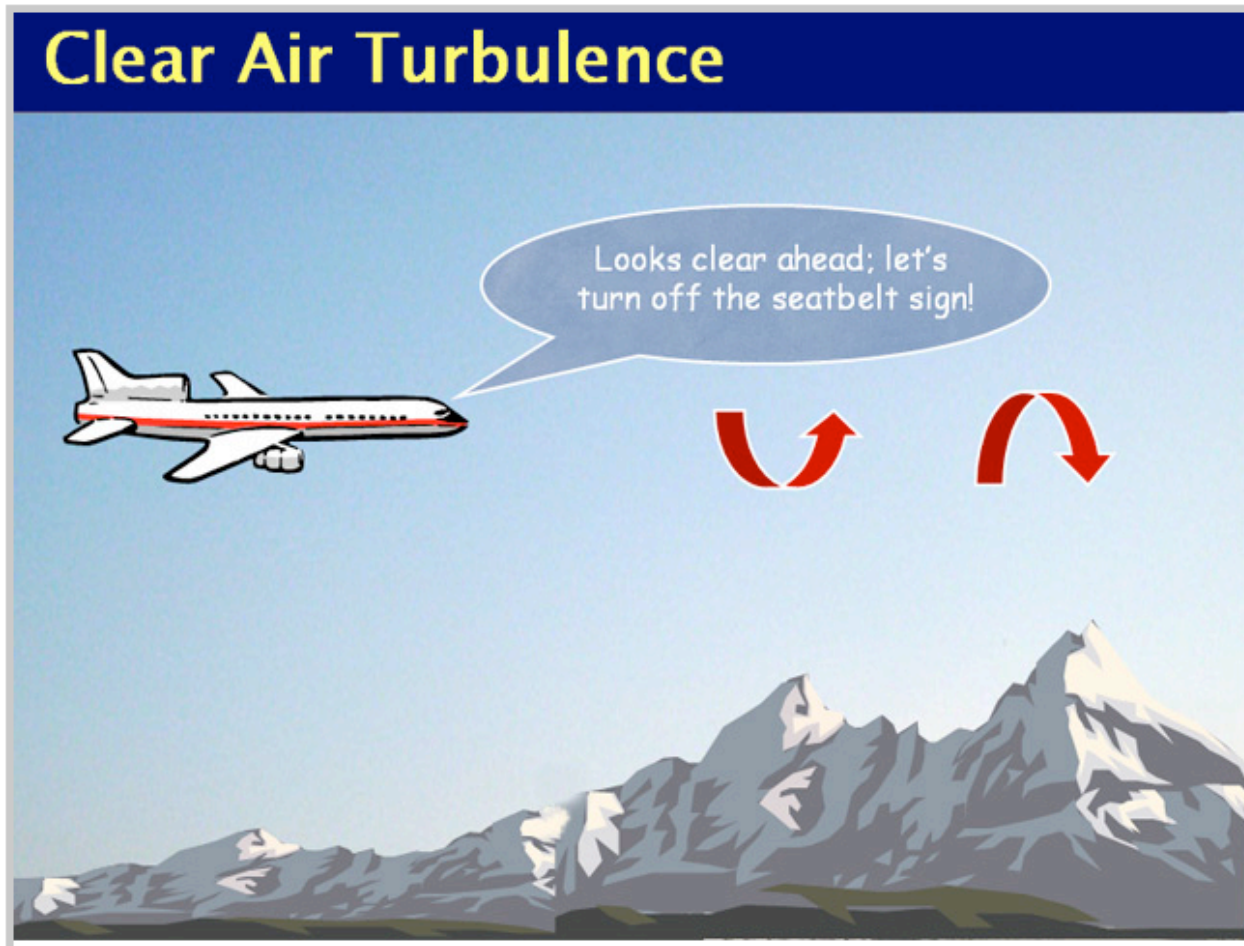
Mountain waves over  
southeastern CO and northeastern NM



Same area is almost cloud-free in the  
visible channel

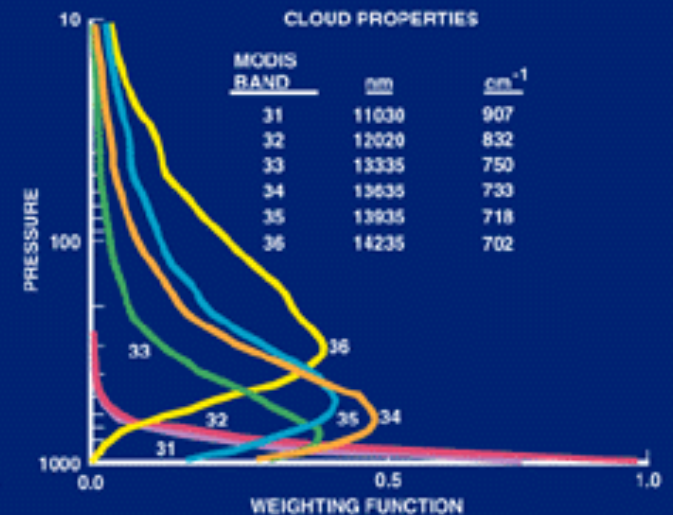
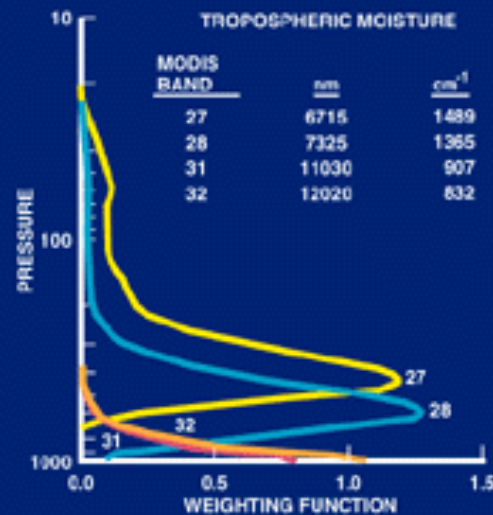
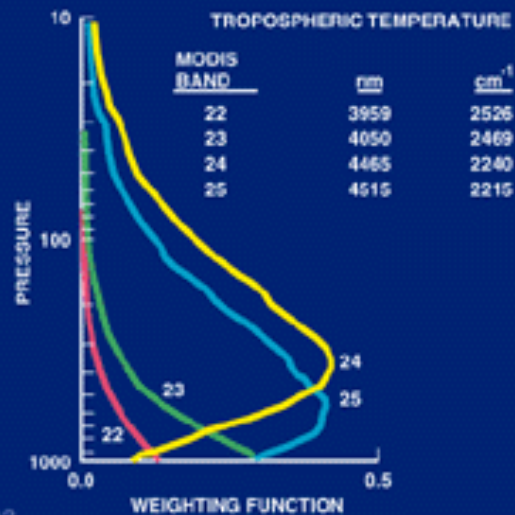
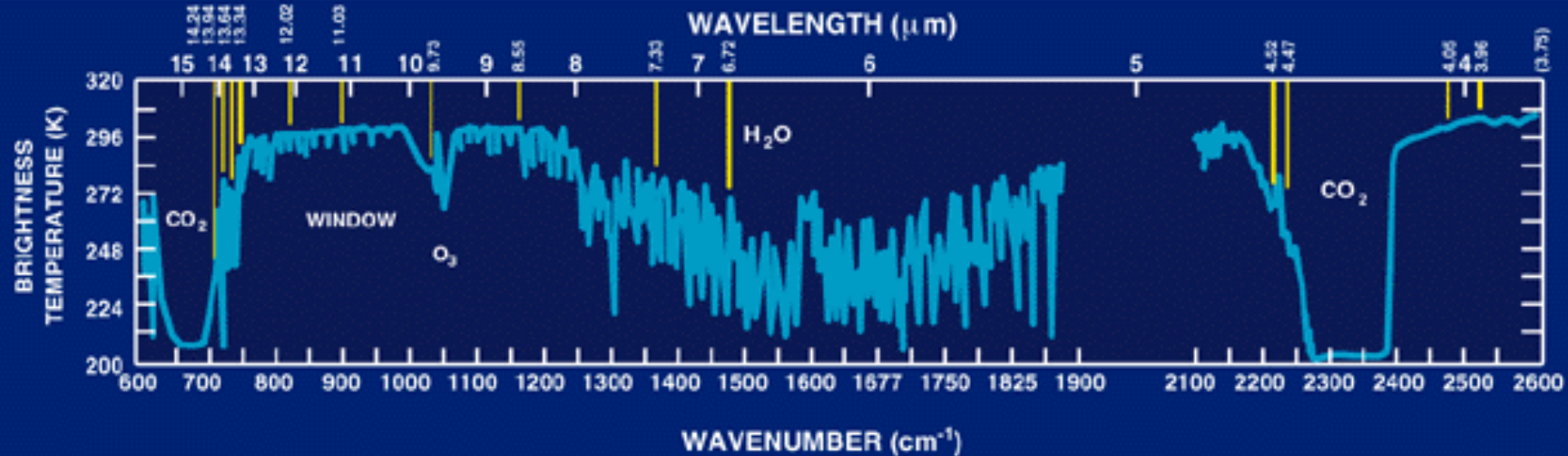


# Why is This Important?



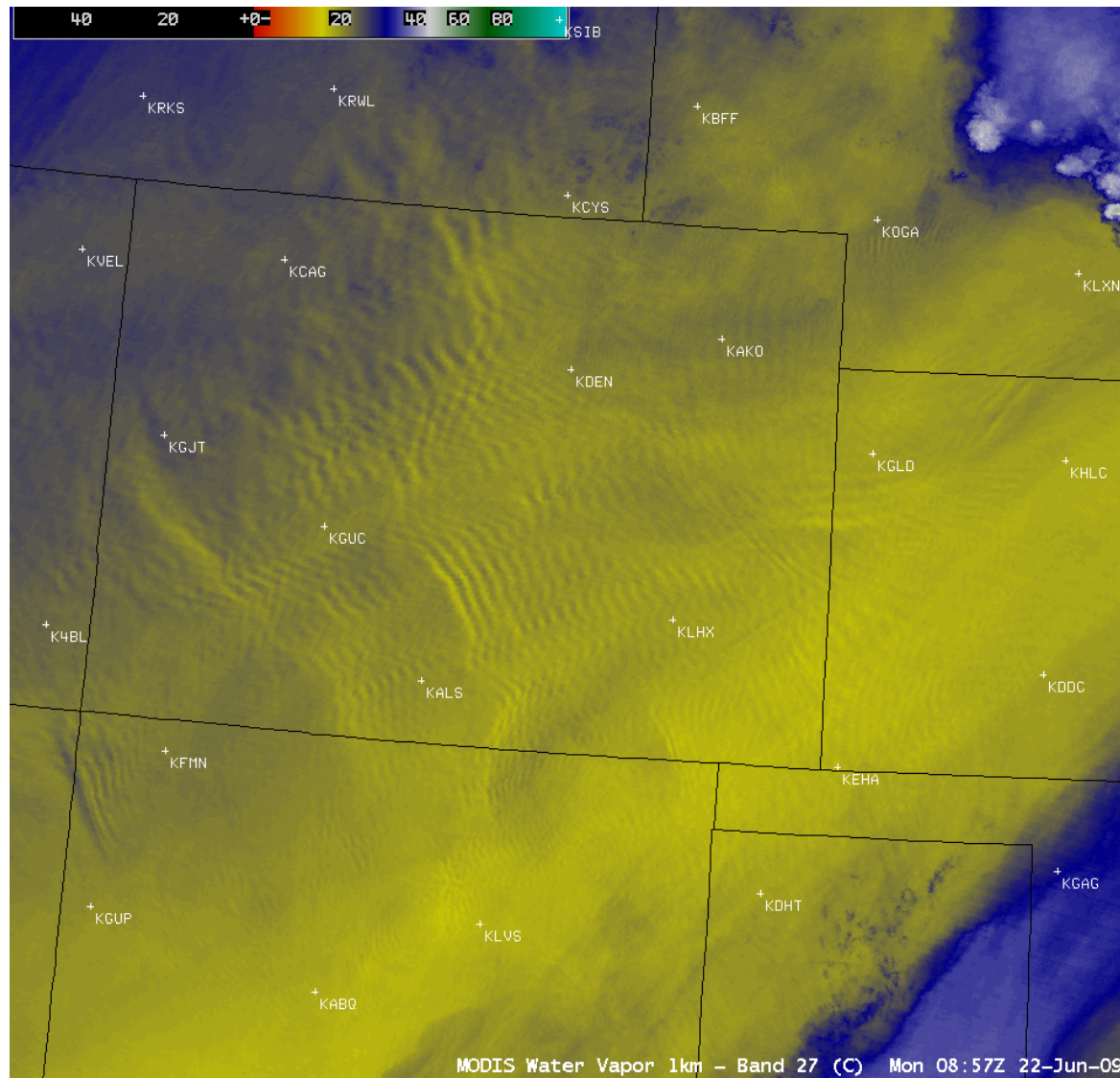


# ATMOSPHERE - THERMAL RADIATION





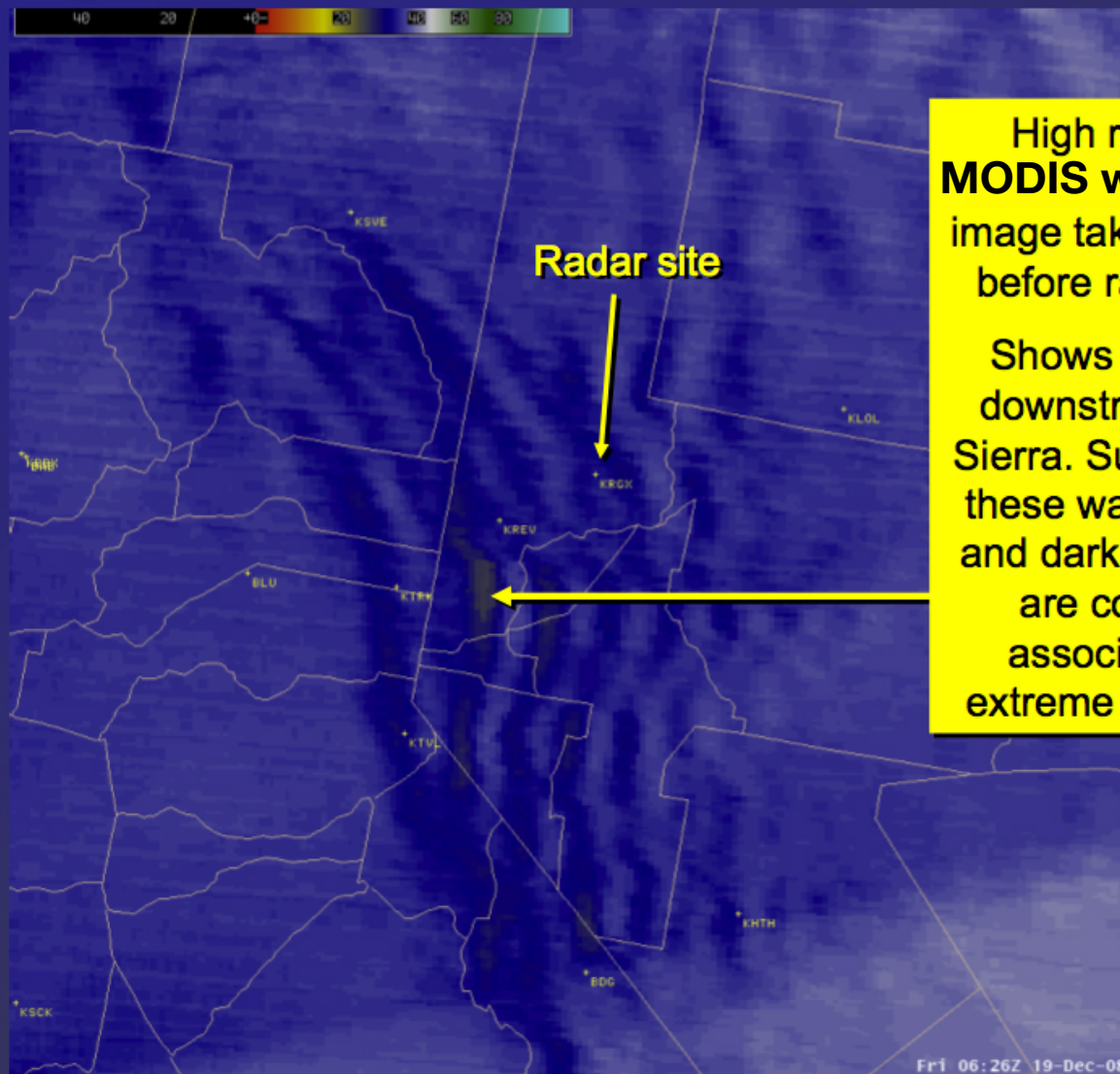
# Mountain Wave Clouds in Clear Air



MODIS and  
GOES  
08:57 UTC  
22 June 2009



# Lee Waves



High resolution  
**MODIS water vapor**  
image taken ~4 hours  
before radar failed.

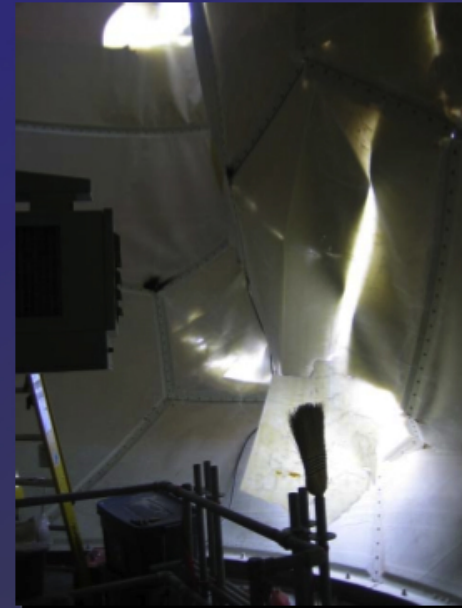
Shows lee waves  
downstream of the  
Sierra. Subsidence in  
these waves (yellow  
and dark blue areas)  
are commonly  
associated with  
extreme wind gusts.

*(credit: NWS forecast office, Reno NV)*



# Photos

Photos taken by NWS Reno electronics team, on first visit to radar after dome failure (19 Dec.).

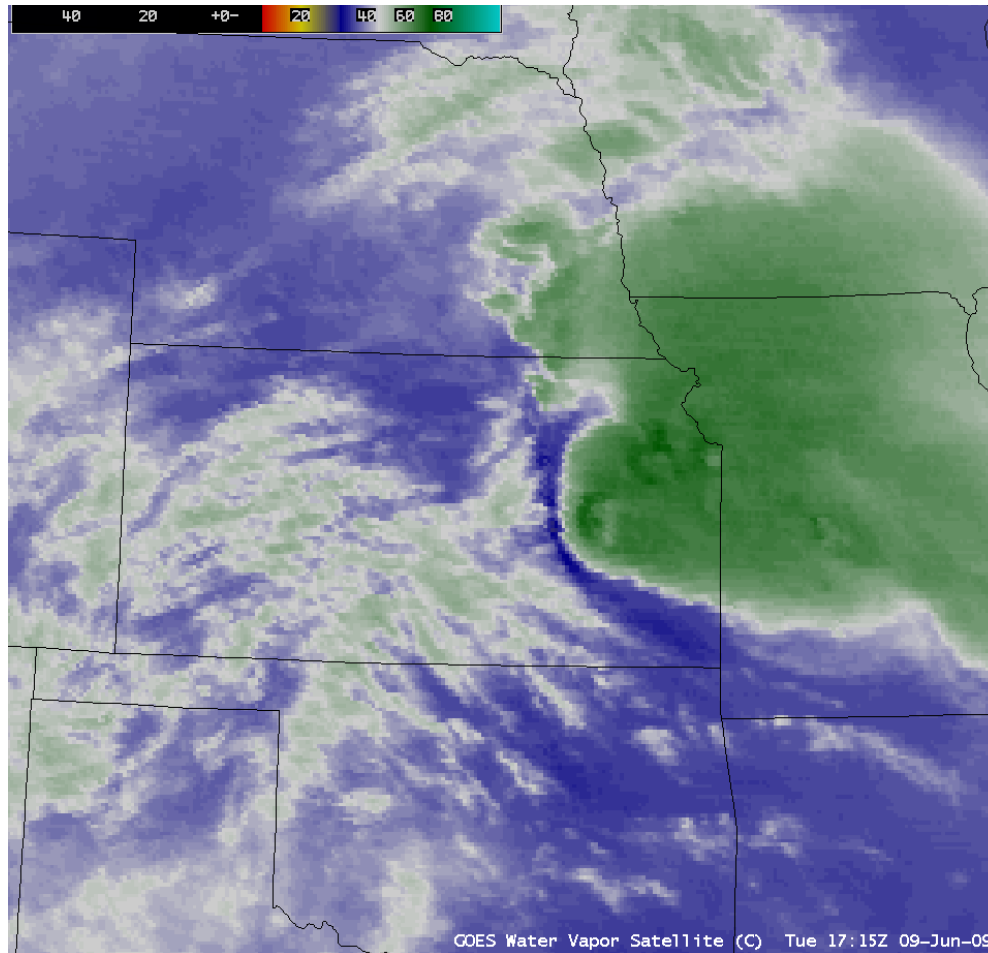


*(credit: NWS forecast office, Reno NV)*



# Turbulence Not Just from Orography

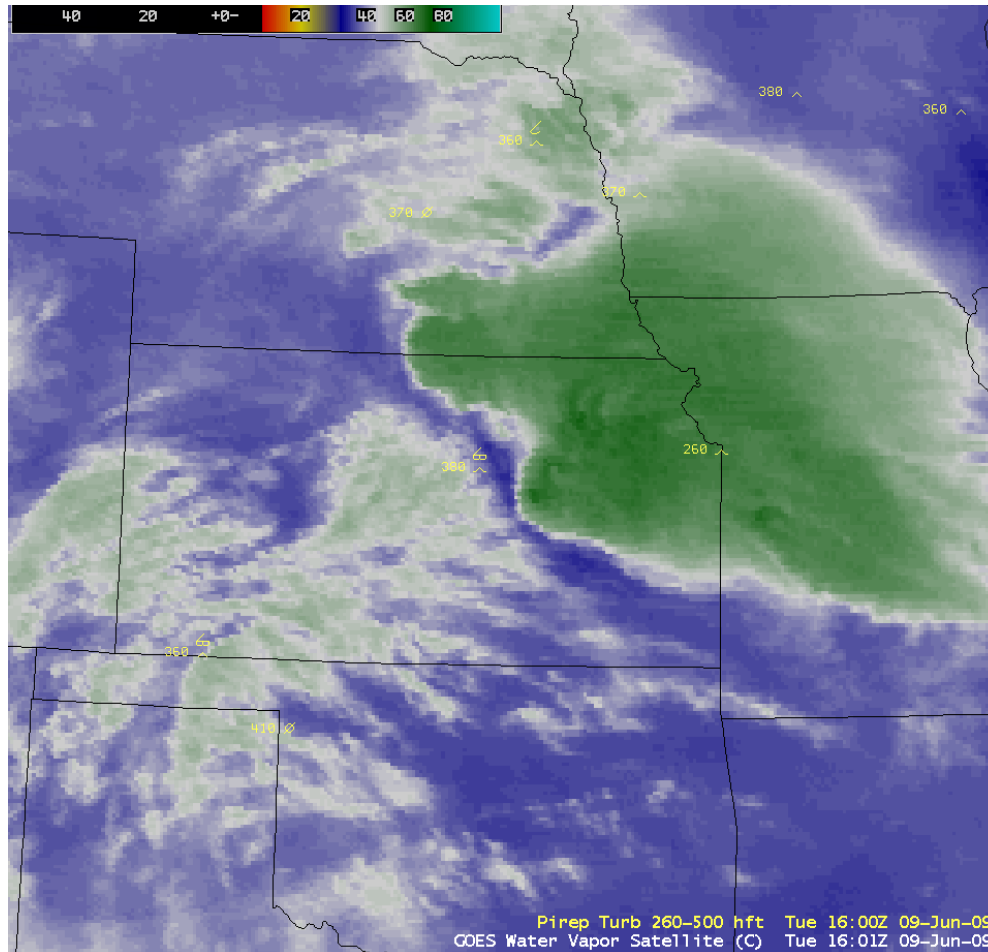
MODIS 6.7 $\mu$ m  
Water Vapor  
Band  
17:15 UTC  
9 June 2009





# Pilot Reports of Turbulence

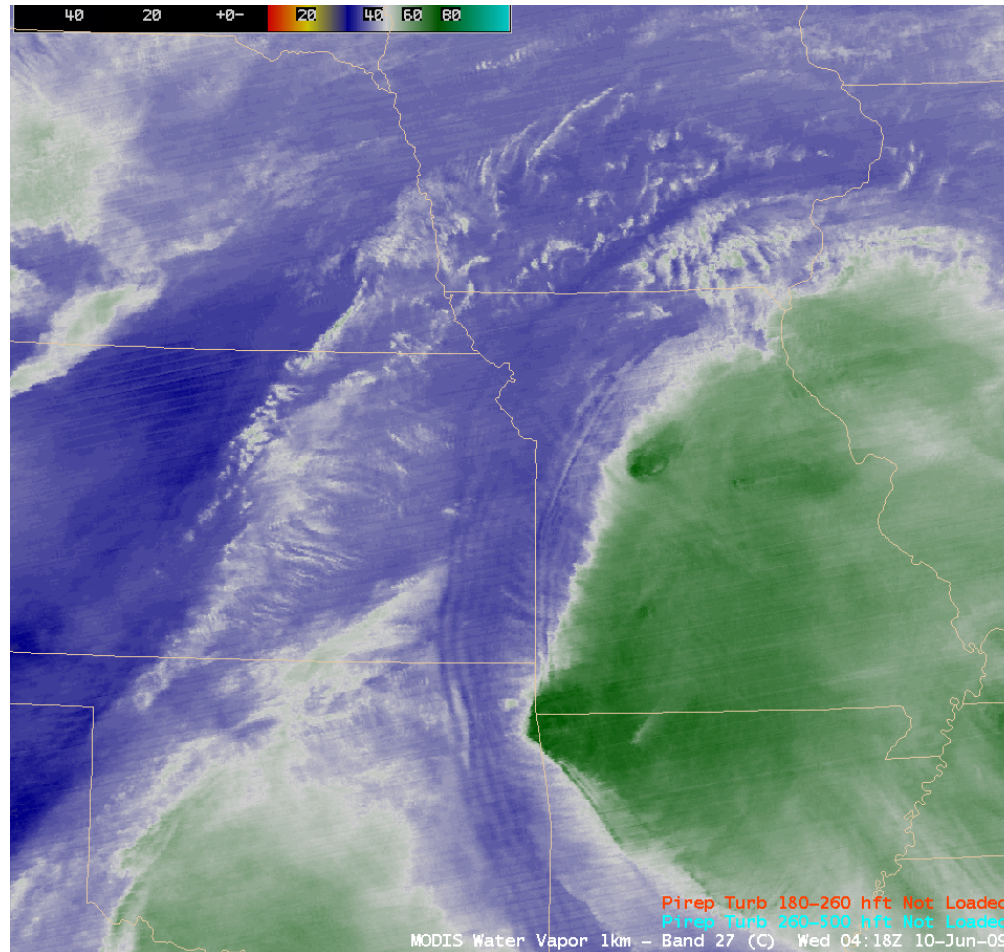
GOES 6.5 $\mu$ m  
Water Vapor  
Band  
16:00 UTC  
9 June 2009





# Turbulence Not Just from Orography

MODIS 6.7  $\mu\text{m}$   
Water Vapor  
Band  
04:18 UTC  
10 June 2009





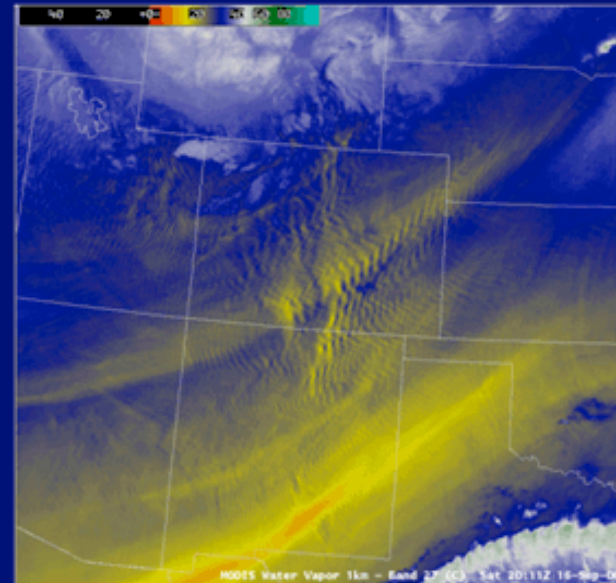
# Why is 6.7 $\mu\text{m}$ Important for the Detection of Turbulence?

## Summary

Turbulence is a significant hazard to aviation, and satellite imagery can sometimes be a helpful tool in turbulence detection.

Mountain waves are one common cause of turbulence, and water vapor channel imagery has the ability to detect areas where this type of turbulence may be present.

The typical "herringbone" signature of mountain waves often occurs in clear (cloud-free) air, making the water vapor channel the only tool for accurate turbulence detection in those cases.





# Cloud Applications Continued

- Clouds
  - Composition
  - Cloud Top Properties
  - Cloud Phase



# Clouds

- MOD06 Cloud Product

Example filename: a1.13214.2325.mod06ct.hdf

- Cloud Top Properties at 5km

- Cloud Top Pressure, Cloud Top Temperature, Cloud Fraction, Cloud Emissivity

- Cloud Phase at 5 km

- Cloud Optical Properties at 1 km (Daytime only)

- Cloud Effective Radius
    - Cloud Optical Thickness

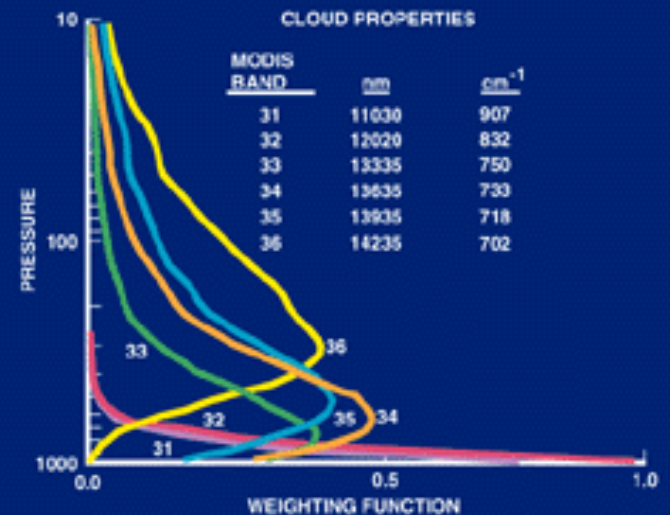
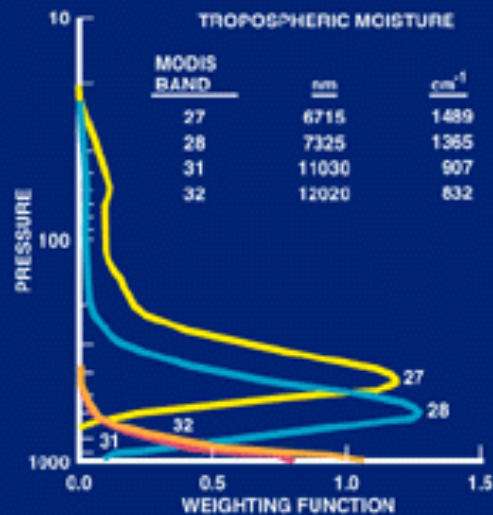
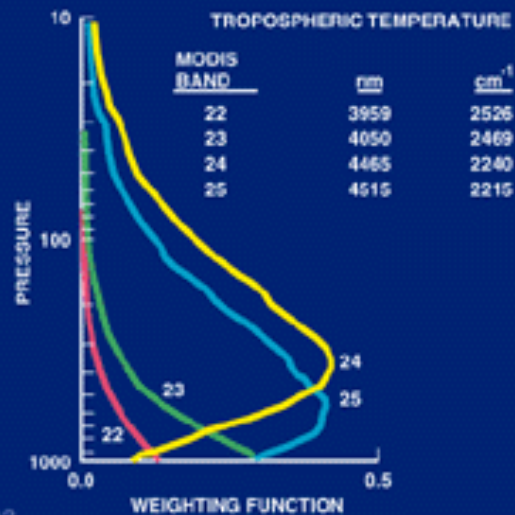
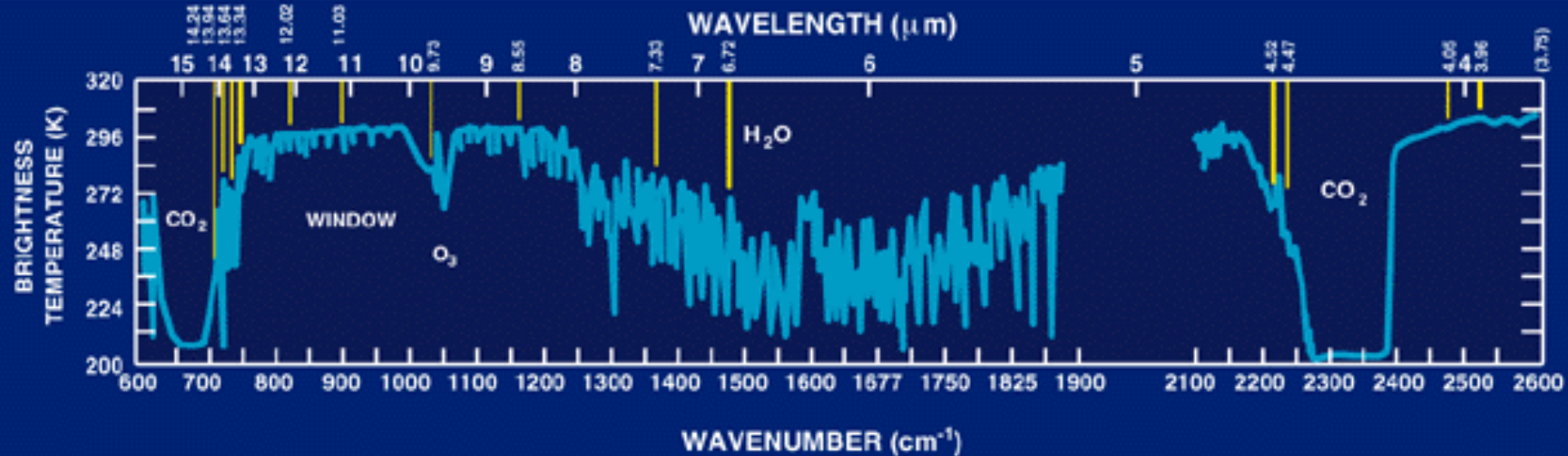


# Cloud Top Property Algorithm

- Cloud Top Pressure, Temperature, Emissivity derived using CO<sub>2</sub> “slicing”
- MODIS product utilizes 4 spectral channels in the 13 – 14  $\mu\text{m}$  region.
- 5x5 1 km pixel retrievals where at least 5 of the 1 km pixels are cloudy as determined by the cloud mask
- Cloud properties retrieved both day and night



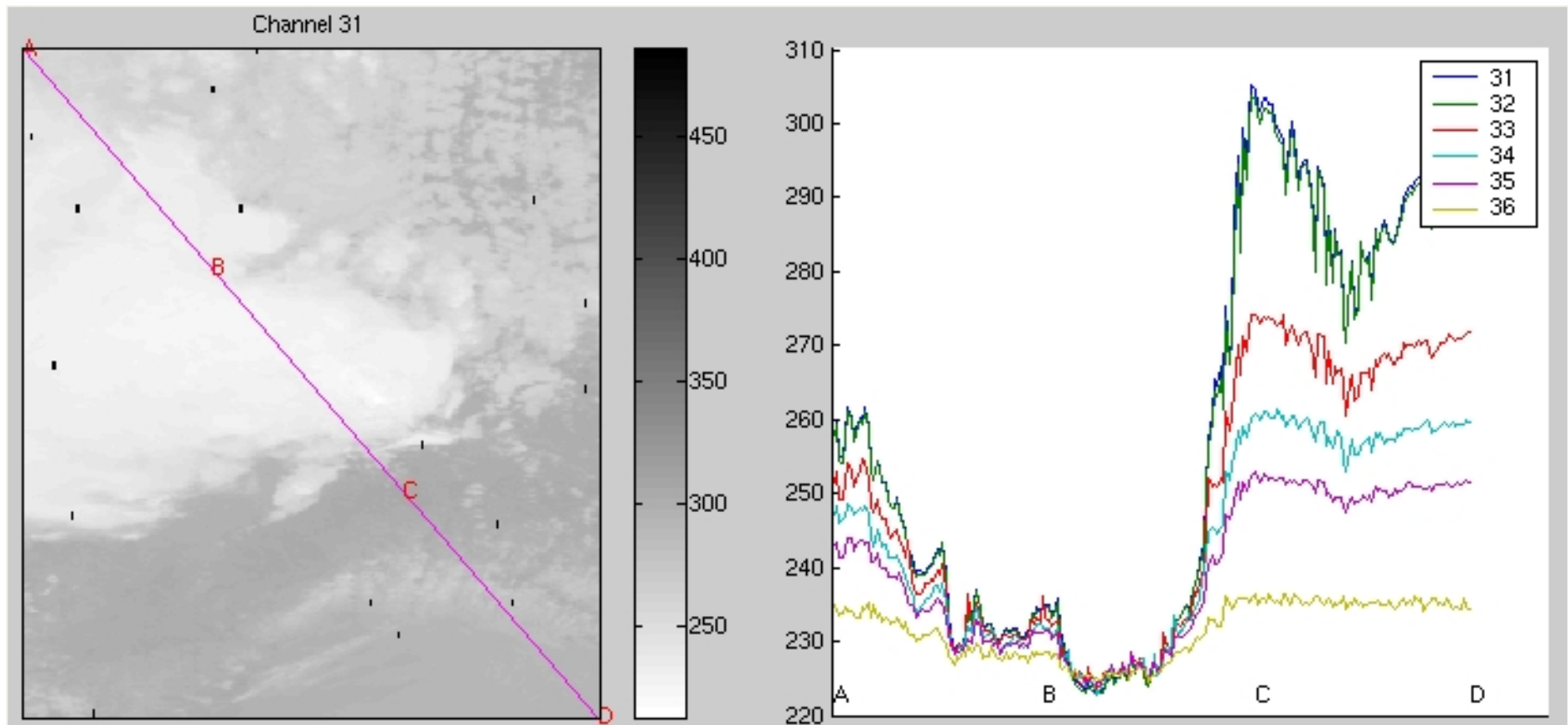
# ATMOSPHERE - THERMAL RADIATION





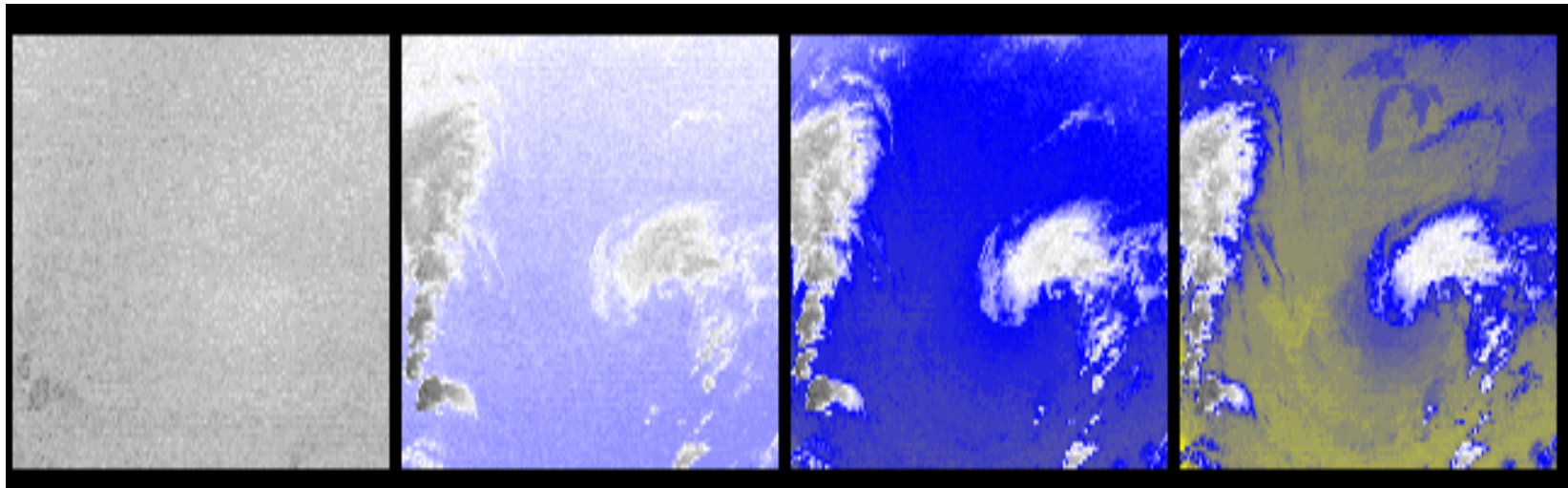
## BT in and out of clouds for MODIS CO<sub>2</sub> bands

- demonstrate weighting functions and cloud top algorithm





**CO2 channels see to different levels in the atmosphere**



14.2 um

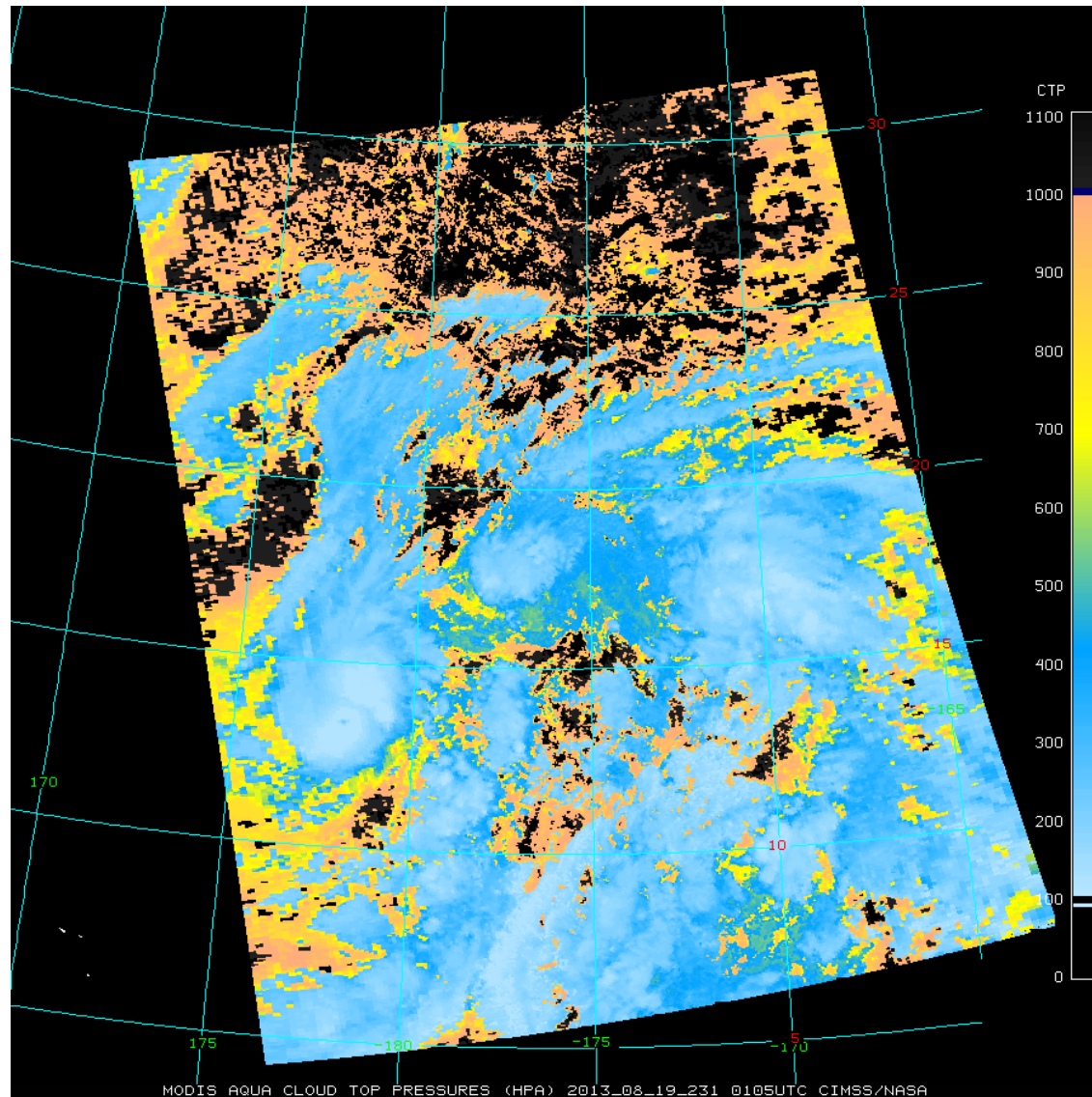
13.9 um

13.6 um

13.3 um

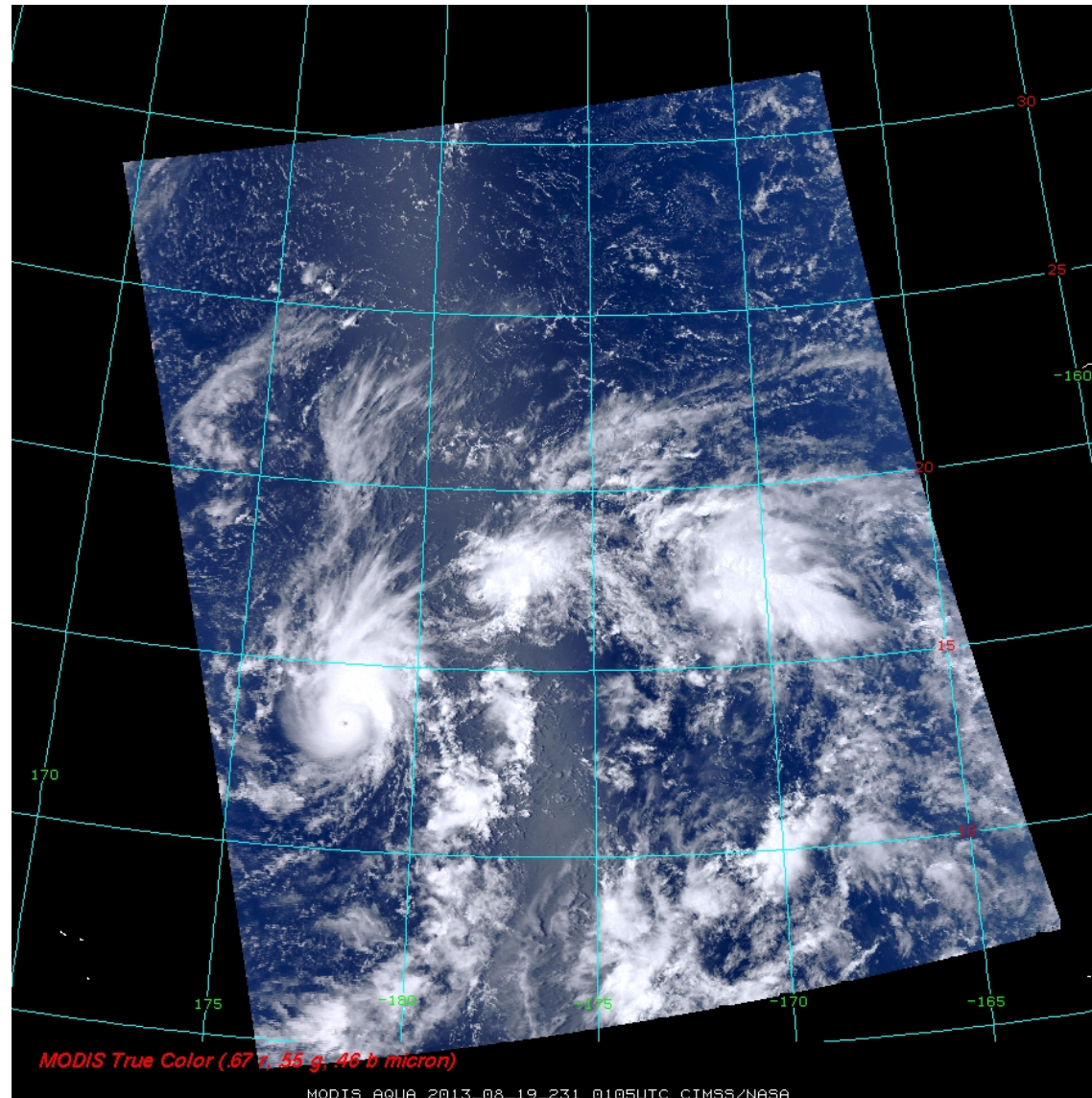


# Example Cloud Top Pressure Product





# Example Cloud Top Pressure Product

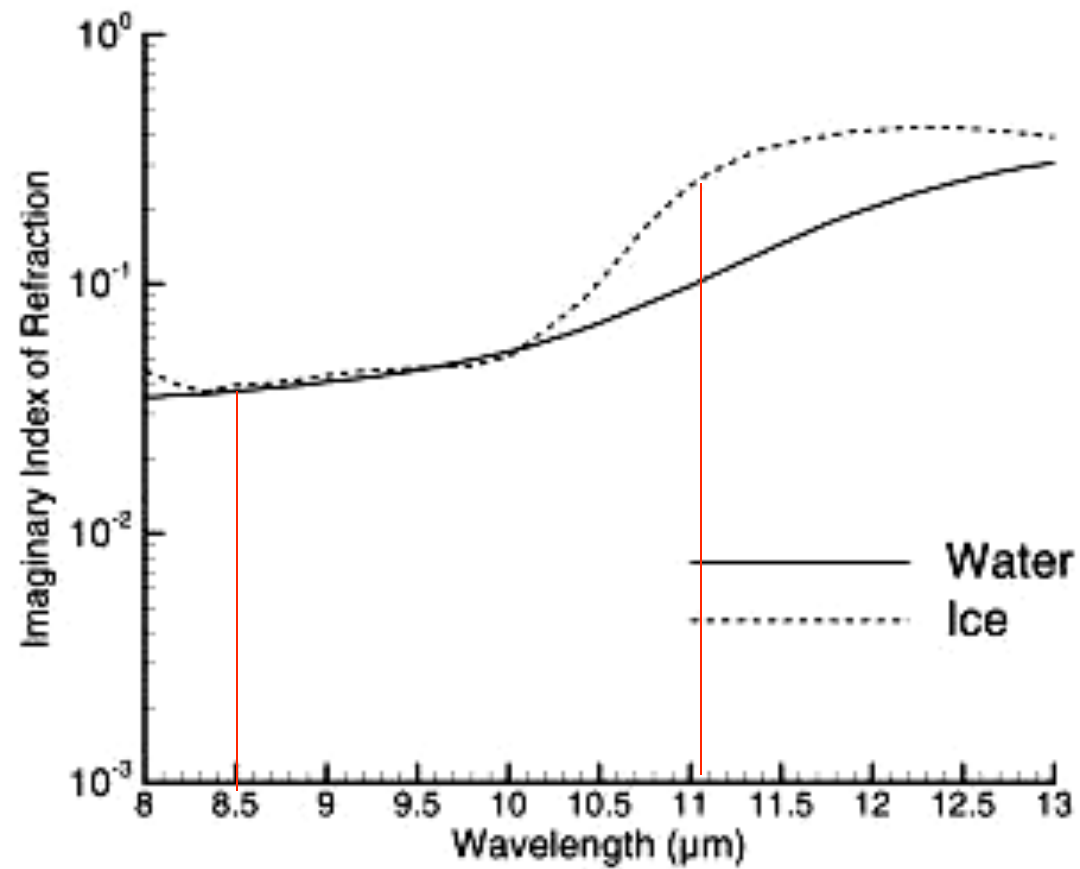




# Cloud Phase

- IR Brightness Temperature Difference Product
  - Band 29 (8.6  $\mu\text{m}$ ) – Band 31 (11  $\mu\text{m}$ )
  - Takes advantage of difference in water/ice cloud absorption in this spectral region
- Near Infrared Bands (1.6 and 2.1  $\mu\text{m}$ )
- Short Wave Infrared Bands (4  $\mu\text{m}$  region)

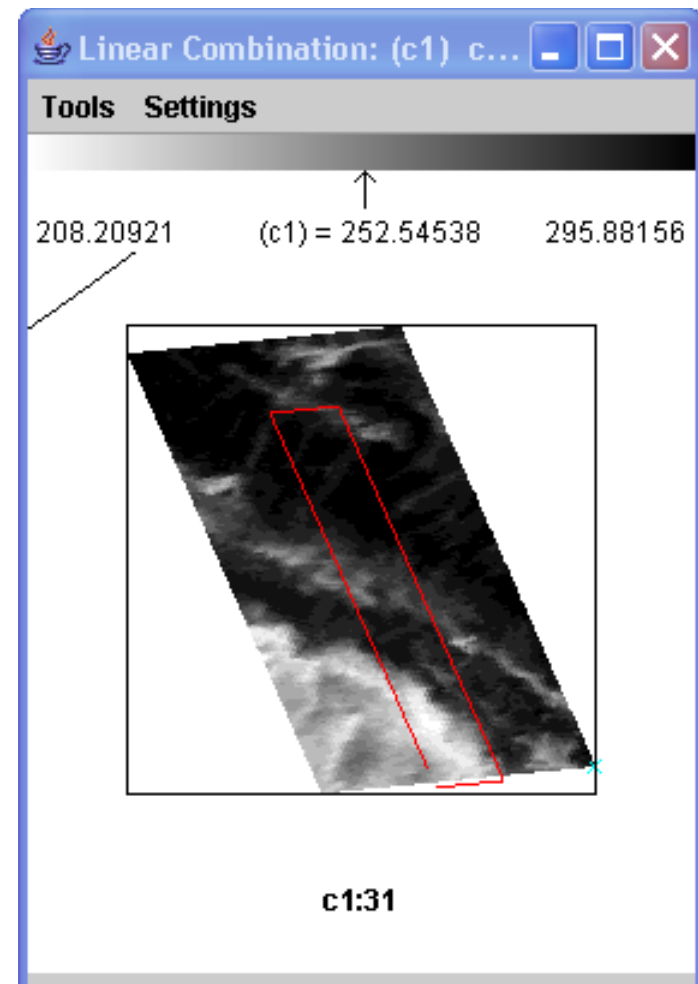
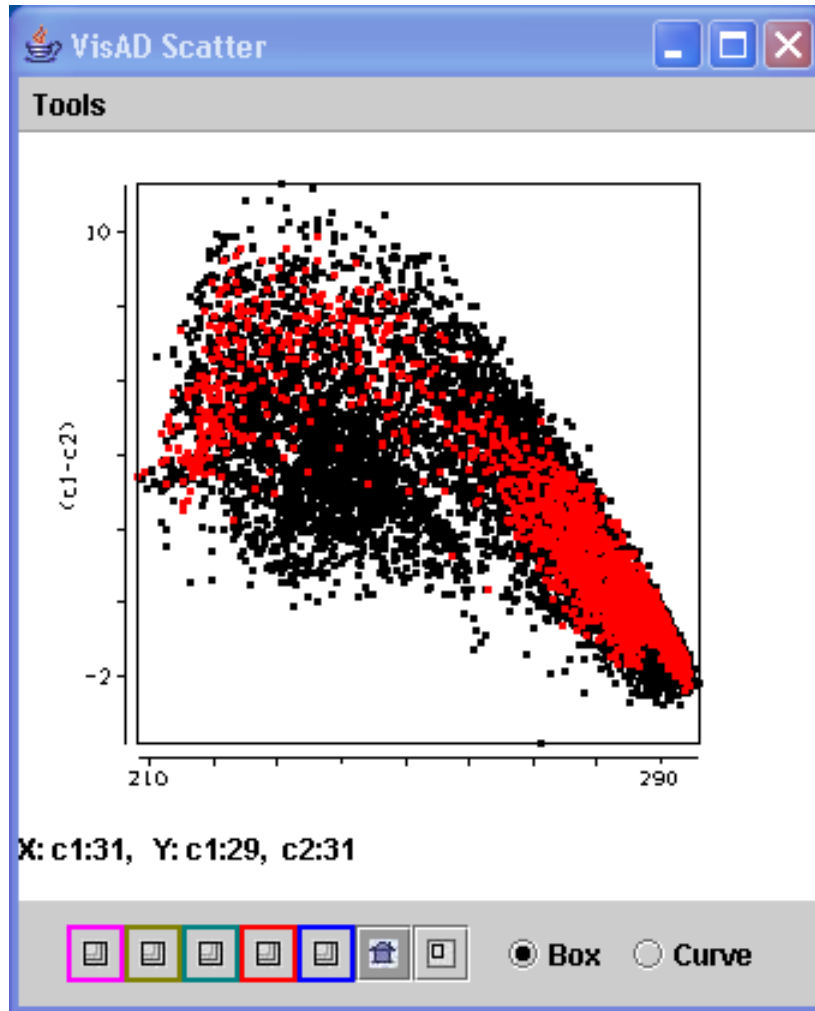




Imaginary Index of Refraction of Ice and Water 8 – 13 microns

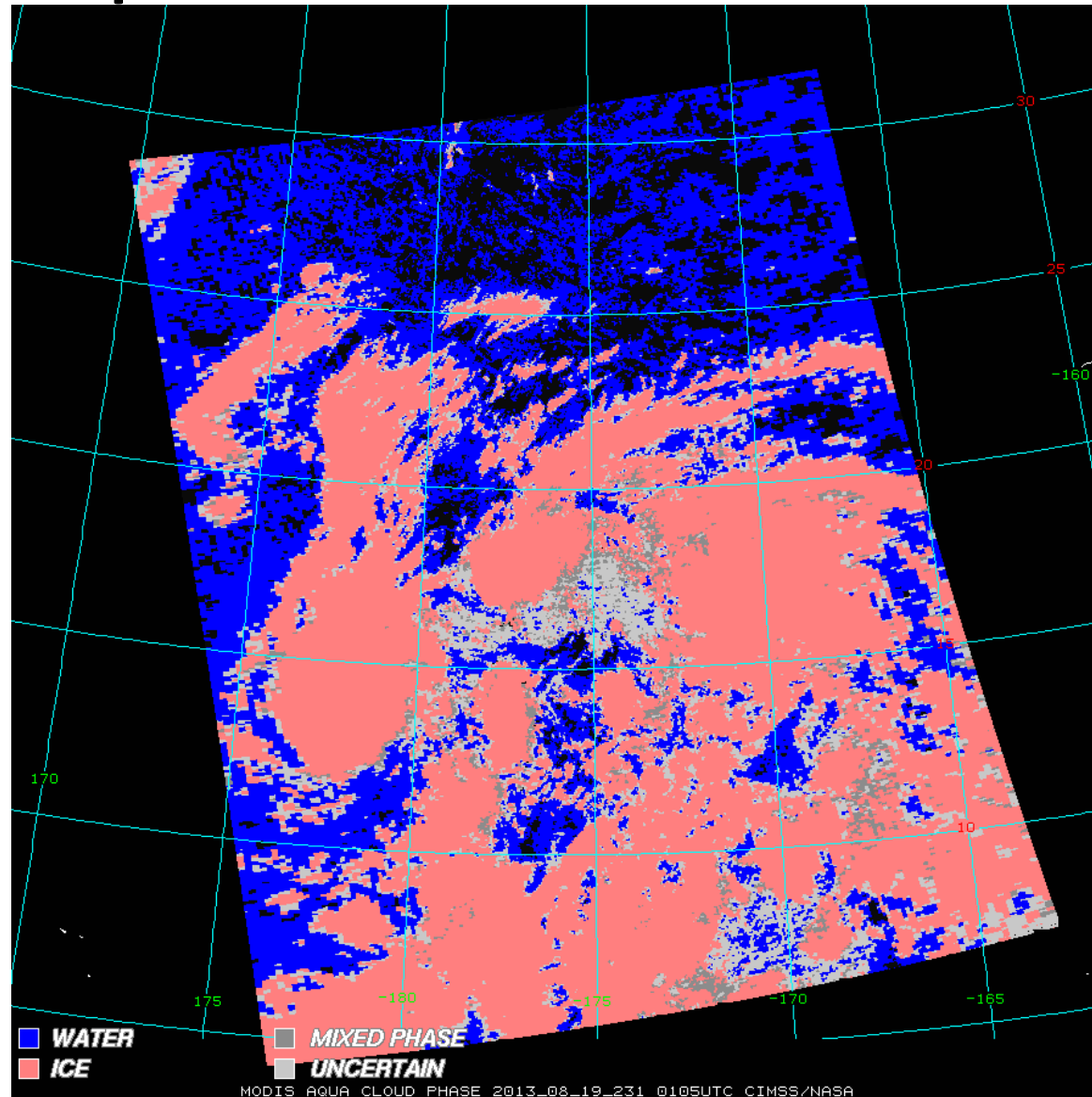


# Ice Cloud Example



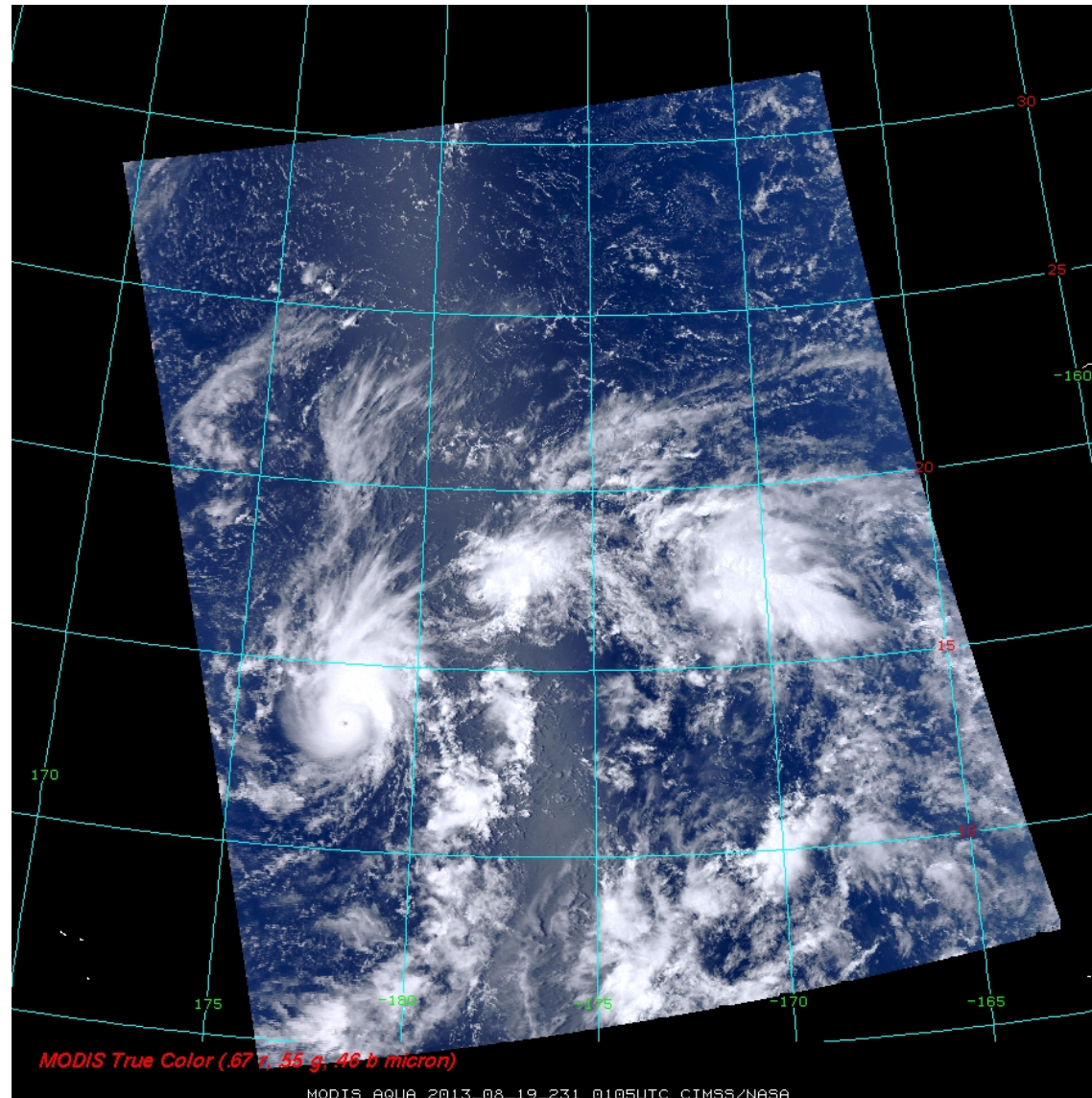


# Example Cloud Phase Product



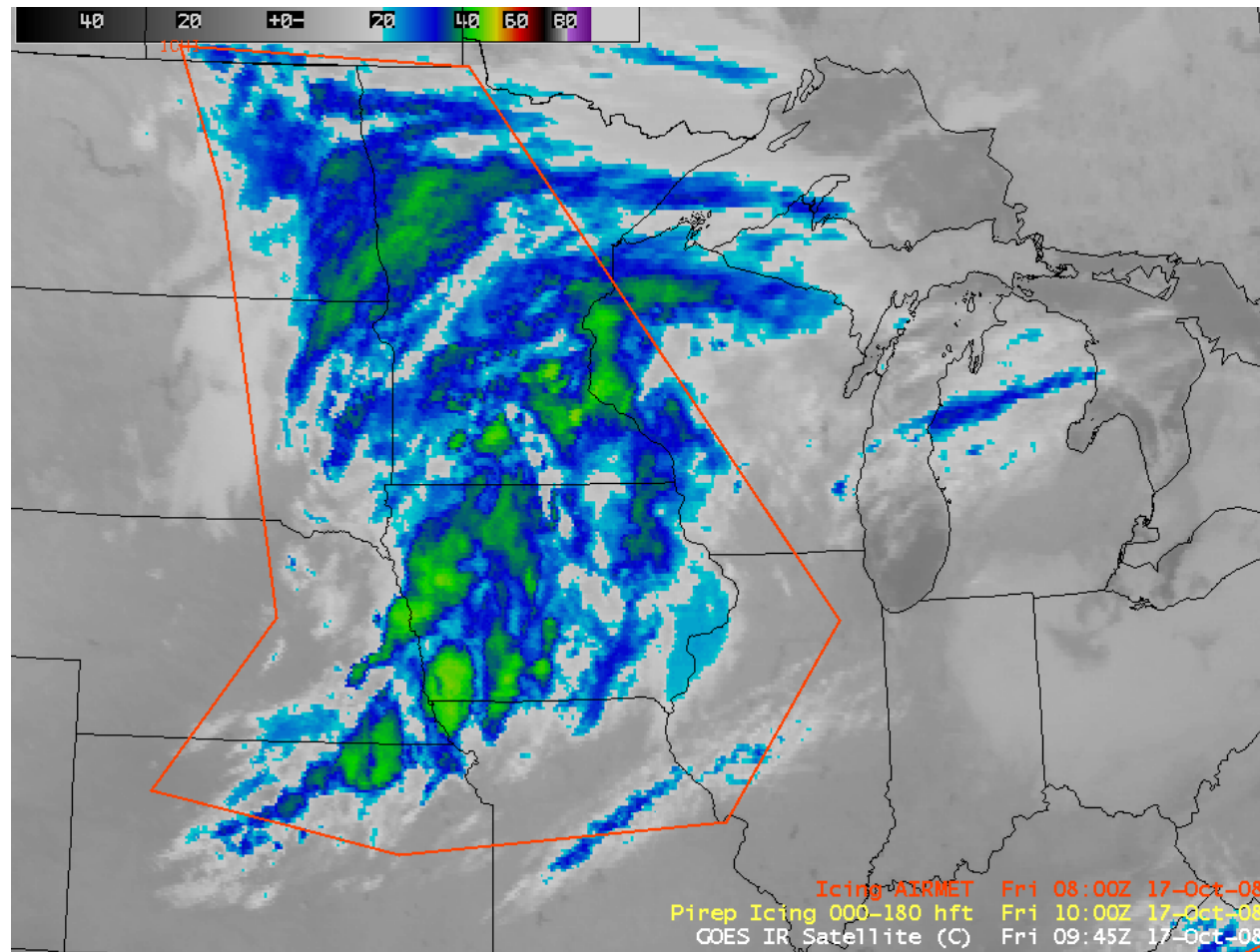


# Example Cloud Top Pressure Product





# Using Satellite Imagery to Help Diagnose Areas of Aircraft Icing Potential



GOES IR window animation 17 October 2008



# Icing

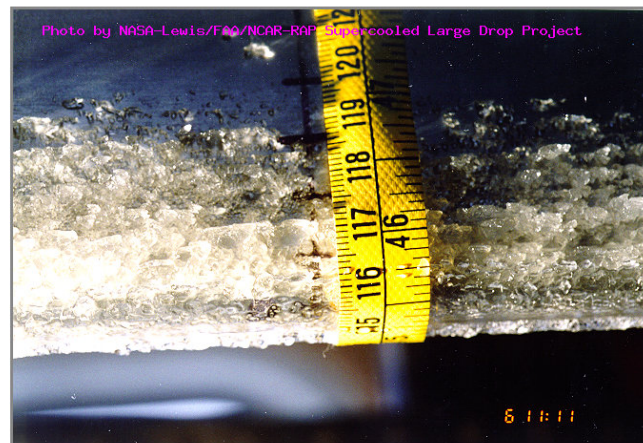
- Freezing Level
  - Altitude at which the temperature is 0 degrees C
  - Above this level the temperature is  $< 0$  C
- Water Can Exist at Temperatures Well Below Freezing
  - Supercooled Water
- An airplane whose temperature is  $<$  freezing in this environment can accrue ice



# Why is This Important?

- We worry about icing because it can adversely affect the flight characteristics of an aircraft. Icing can increase drag, decrease lift, and cause control problems. The added weight of the accreted ice is generally only a factor in light aircraft.

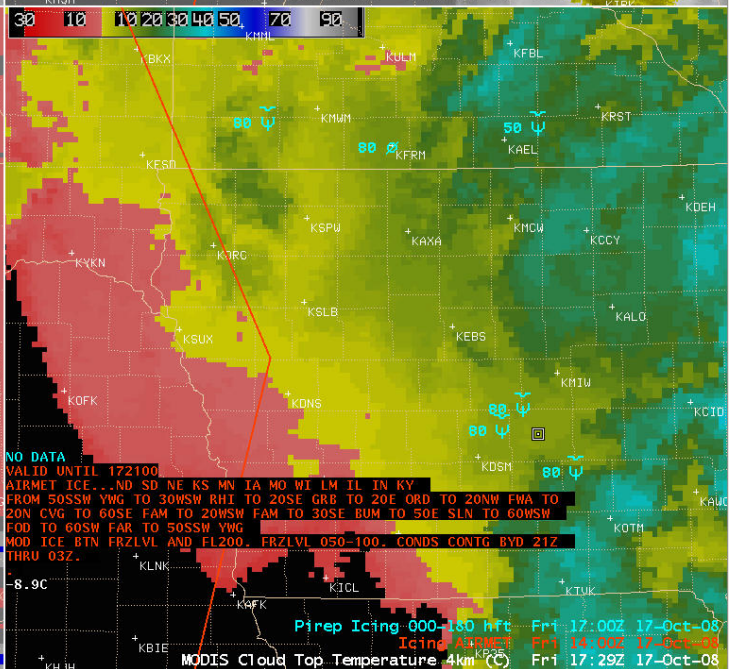
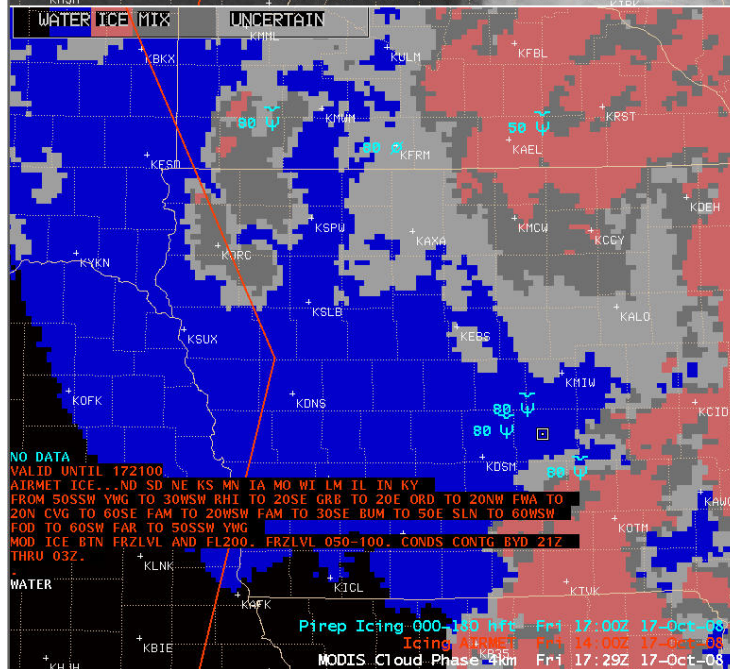
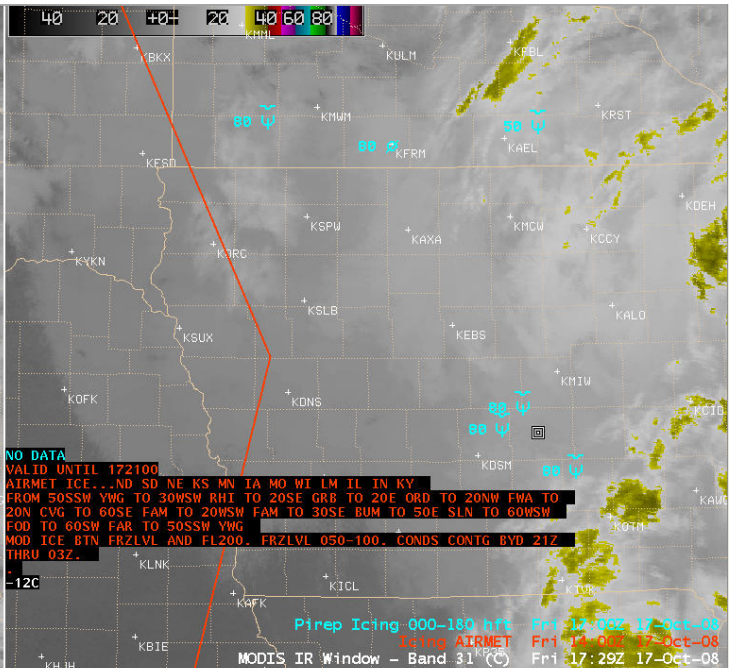
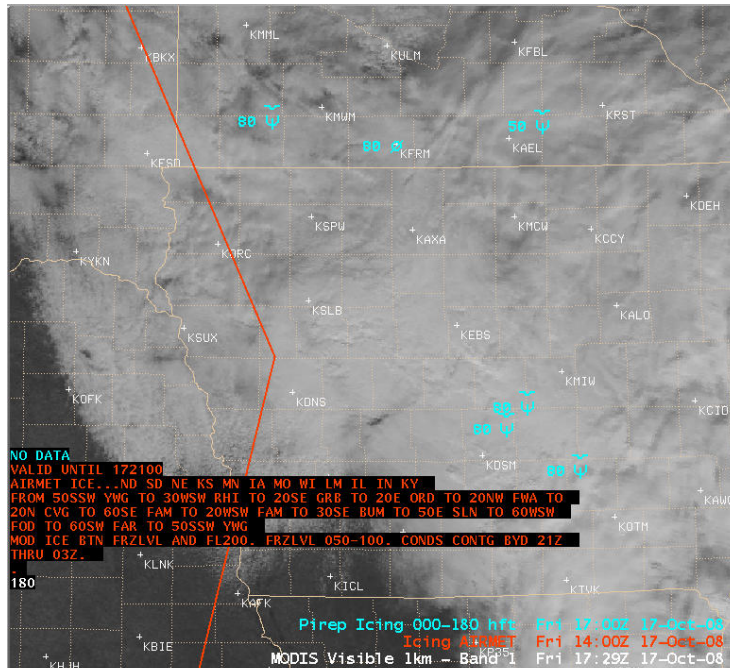
Ice Accumulation On  
The Wing of a Small Aircraft  
NASA





- A closer view using AWIPS images of the MODIS visible channel, 11.0  $\mu\text{m}$  “IR window” channel, Cloud Top Temperature (CTT) product, and Cloud Phase product at 17:29 UTC (below) indicated that much of the cloud shield along the trailing (western) edge of the shortwave over Minnesota and Iowa exhibited cloud top temperatures that were below freezing (generally in the -5 to -12° C range), but the MODIS Cloud Phase product designated those trailing edge clouds as “Water droplet” clouds (blue enhancement). Within this area of supercooled water droplet clouds were several pilot reports of icing at the 8000-foot altitude across southern Minnesota and western/central Iowa.







# Other Cloud Applications Too

- Identification of mature T-storms
  - Must glaciate, meaning tops of cell must be ice
- Accurate height of “thin” high clouds
  - Energy transmitted from below the cloud in the IR window. Can't get accurate level from window BT.
  - Important for pilots. Clouds mean more moisture, dry entrainment and potential for turbulence.



# Other Applications

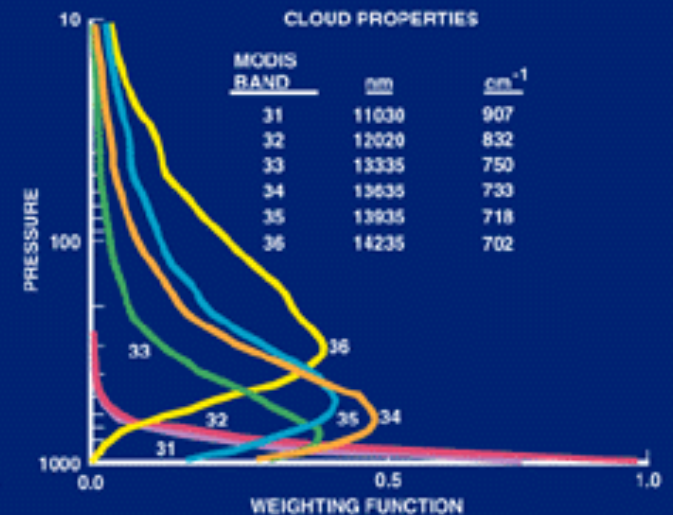
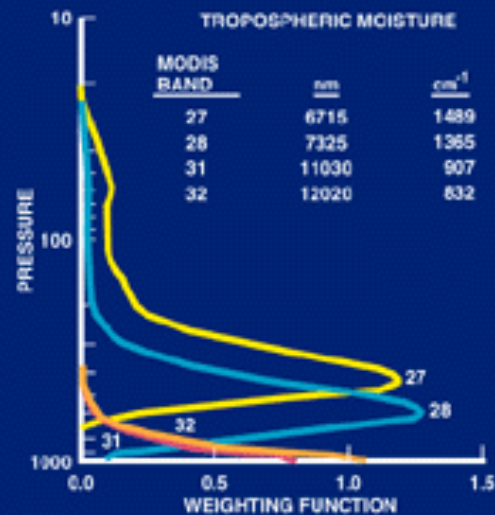
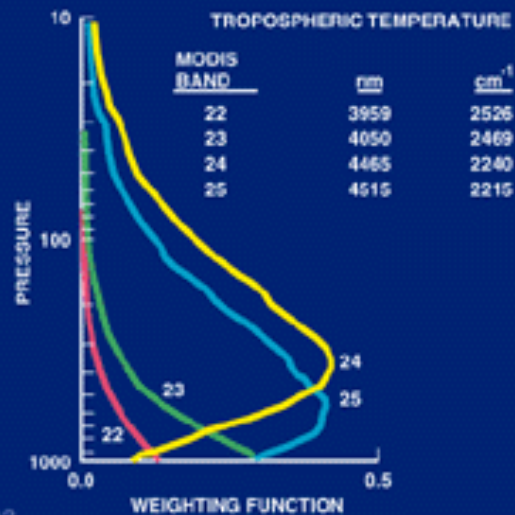
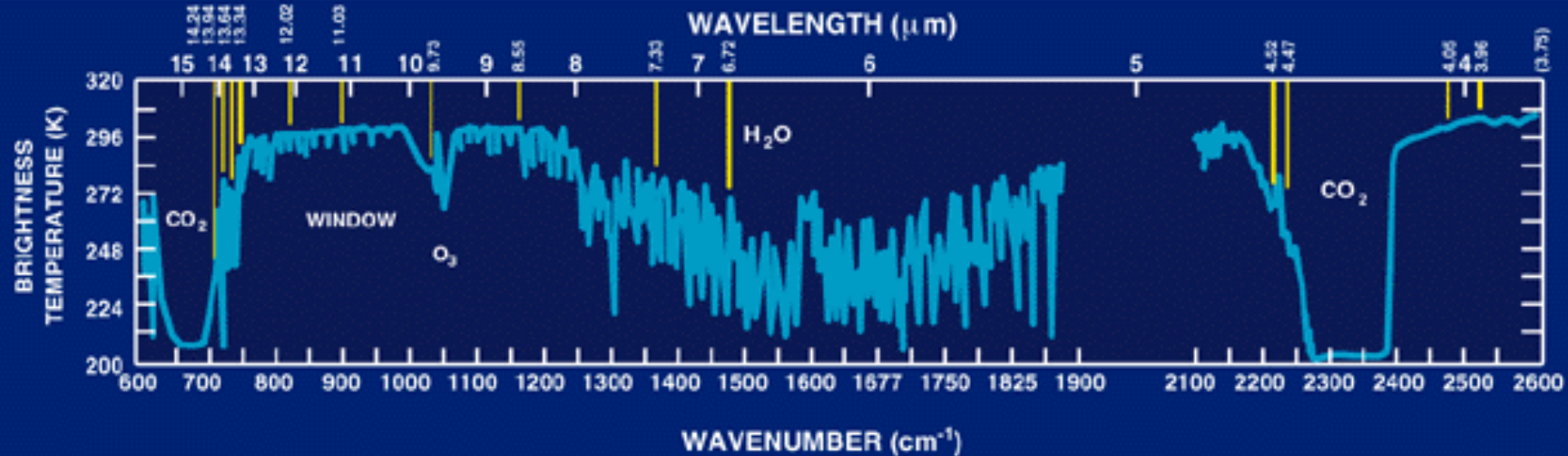


# Sea Surface Temperatures

- Simple Brightness Temperature Difference Algorithm
- “Split Window” technique
- Regression between
  - 11-12  $\mu\text{m}$  BTDIF (MODIS bands 31 and 32)
  - 4 – 11  $\mu\text{m}$  BTDIF (MODIS bands 22 and 31)
    - Must be careful in sunglint regions because of solar contamination
  - In essence, you are trying to correct for the lowering of the observed brightness temperatures by water vapor using the BTDIF between these two window channels

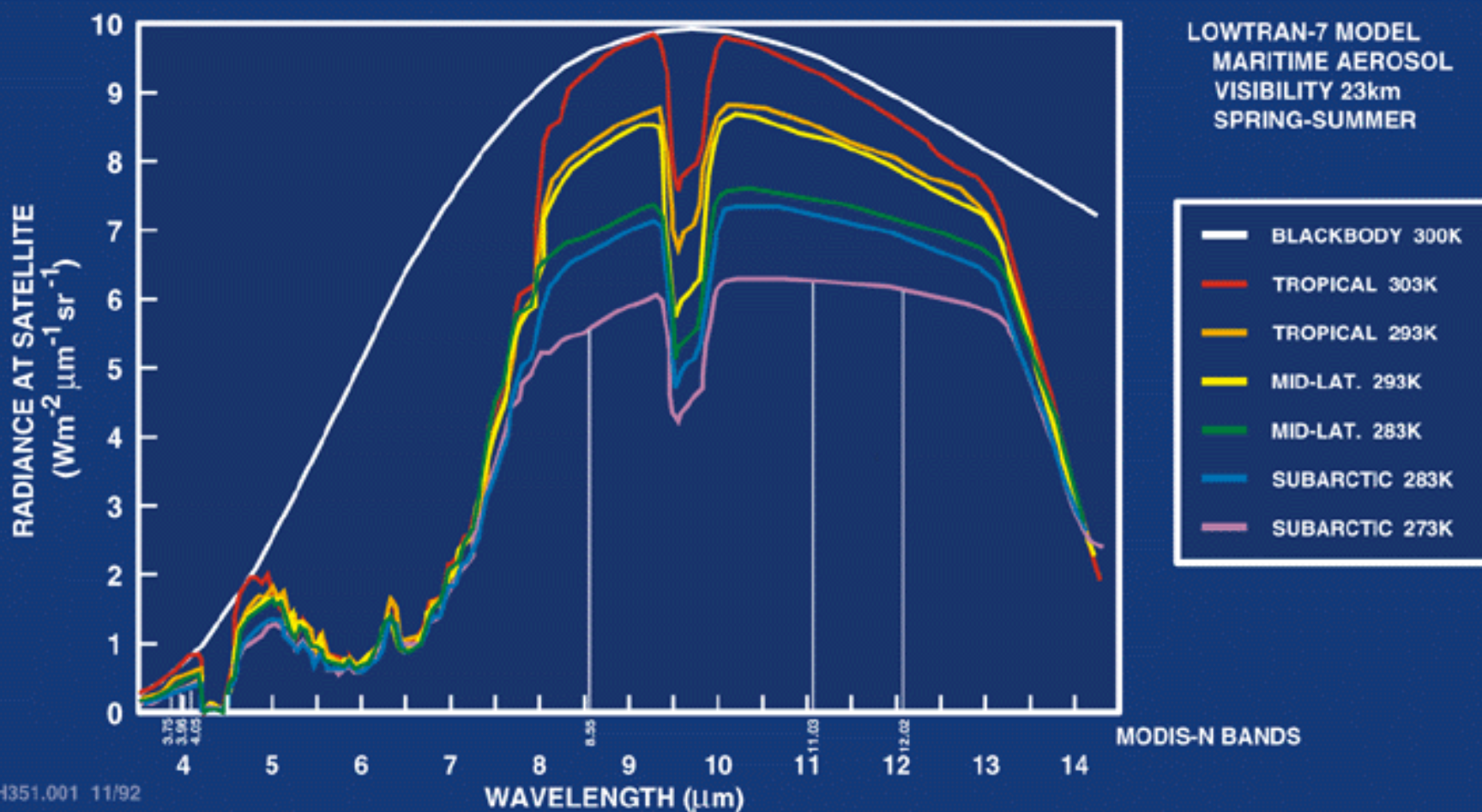


# ATMOSPHERE - THERMAL RADIATION





# MODIS SEA SURFACE TEMPERATURE





# MODIS Longwave Infrared Sea Surface Temperature (c5)

$dBT \leq 0.5$

$$sst = a00 + a01*BT11 + a02*dBT*bsst + a03*dBT*(1.0/\mu - 1.0)$$

$dBT \geq 0.9$

$$sst = a10 + a11*BT11 + a12*dBT*bsst + a13*dBT*(1.0/\mu - 1.0)$$

$0.5 < dBT < 0.9$

$$sstlo = a00 + a01*BT11 + a02*dBT*bsst + a03*dBT*(1.0/\mu - 1.0)$$

$$ssthi = a10 + a11*BT11 + a12*dBT*bsst + a13*dBT*(1.0/\mu - 1.0)$$

$$sst = sstlo + (dBT - 0.5)/(0.9 - 0.5)*(ssthi - sstlo)$$

where:

$$dBT = BT11 - BT12$$

BT11 = brightness temperature at 11  $\mu m$ , in deg-C

BT12 = brightness temperature at 12  $\mu m$ , in deg-C

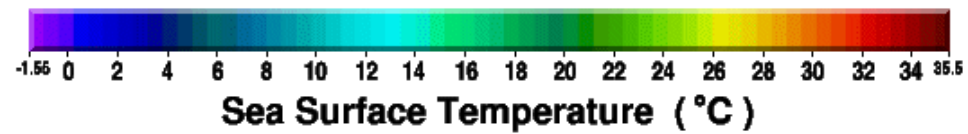
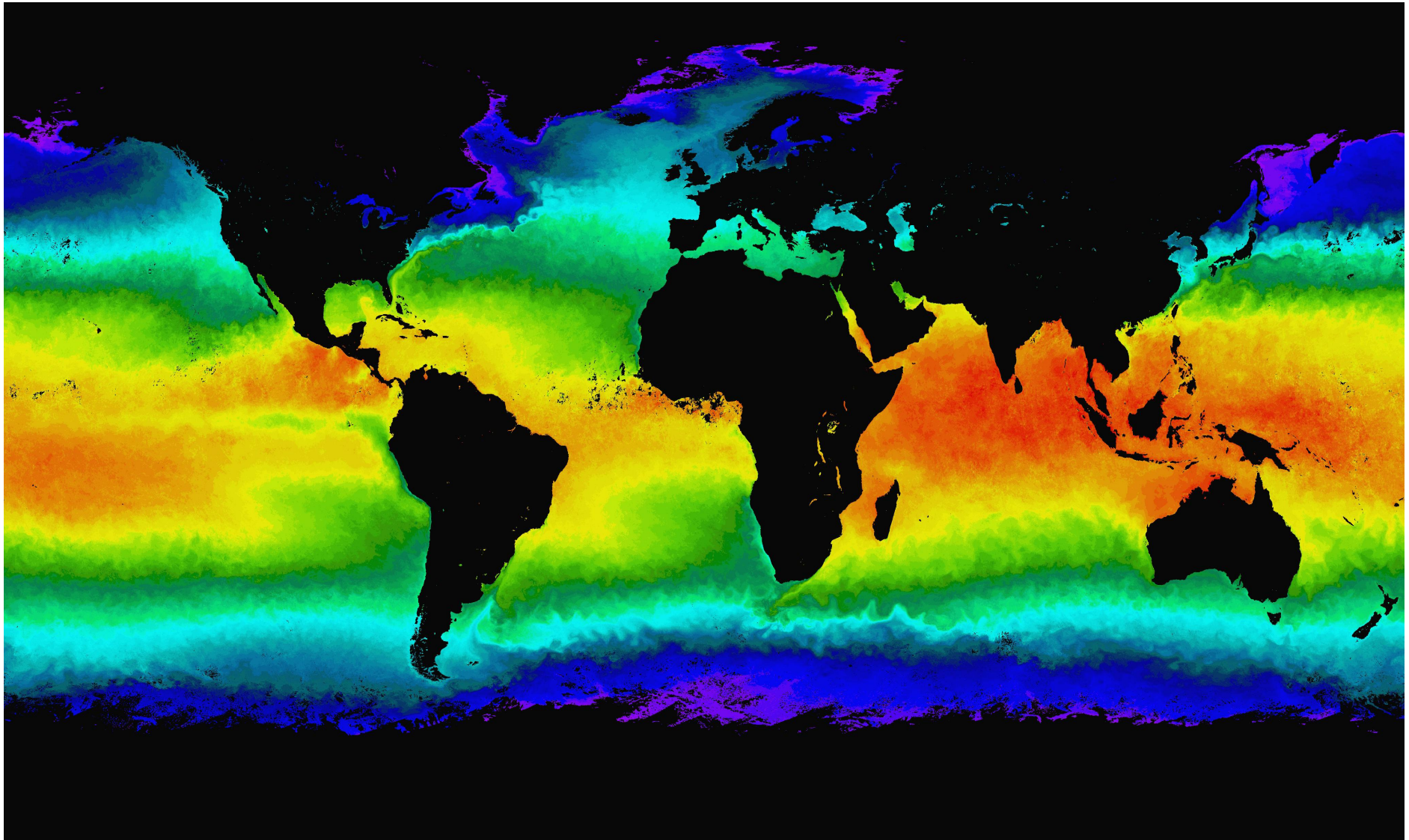
bsst = Either sst4 (if valid) or sstref (from Reynolds OISST)

$\mu$  = cosine of sensor zenith angle

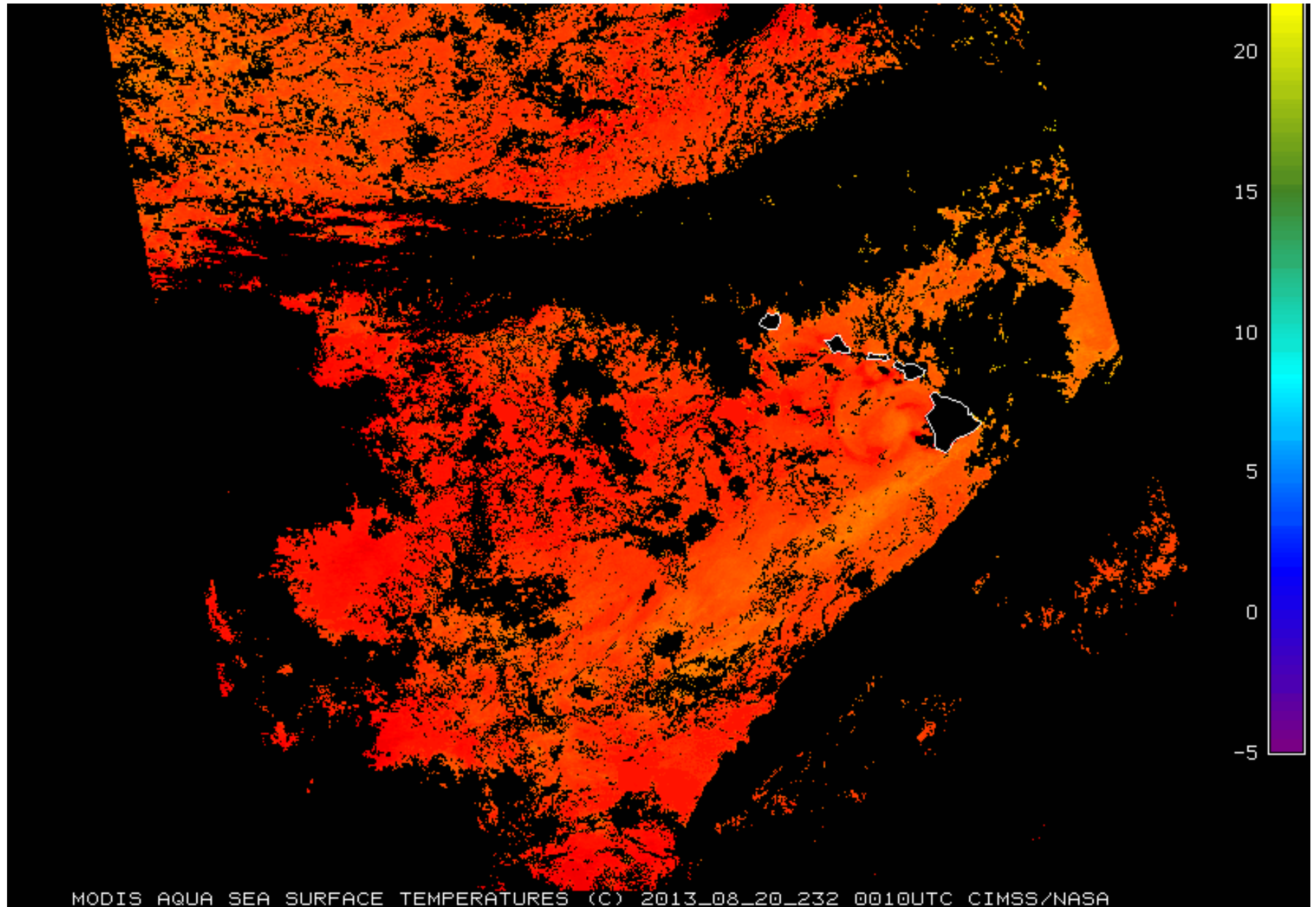
a00, a01, a02, a03, a10, a11, a12, a13 derived from match-ups



# Aqua MODIS Sea Surface Temperature, April 2004









# MODIS Sea Surface Temperature used by Forecasters

AREA FORECAST DISCUSSION...UPDATED

NATIONAL WEATHER SERVICE MILWAUKEE/SULLIVAN WI

338 AM CDT TUE MAY 31 2011

UPDATED TO ADD TODAY/TONIGHT AND AVIATION/MARINE  
SECTIONS

.MARINE...**CLEAR MODIS IMAGE FROM MONDAY EARLY AFTN  
SHOWED SHALLOWER NEAR SHORE WATERS HAD WARMED INTO  
THE LOWER 50S...WHILE MID LAKE TEMPS REMAINED IN THE MID  
40S DUE TO OVERTURNING.** TIGHTENING PRESS GRADIENT THIS  
MORNING AND SUNSHINE WILL RESULT IN STRONG MIXING  
EARLY THIS MRNG. HENCE WL BUMP UP START OF SMALL CRAFT  
ADVY SEVERAL HOURS...AND RUN INTO THE EVE. FEW GUSTS  
NEAR THE SHORE MAY REACH 30-35 KNOTS LATER THIS MRNG/  
EARLY AFTN.

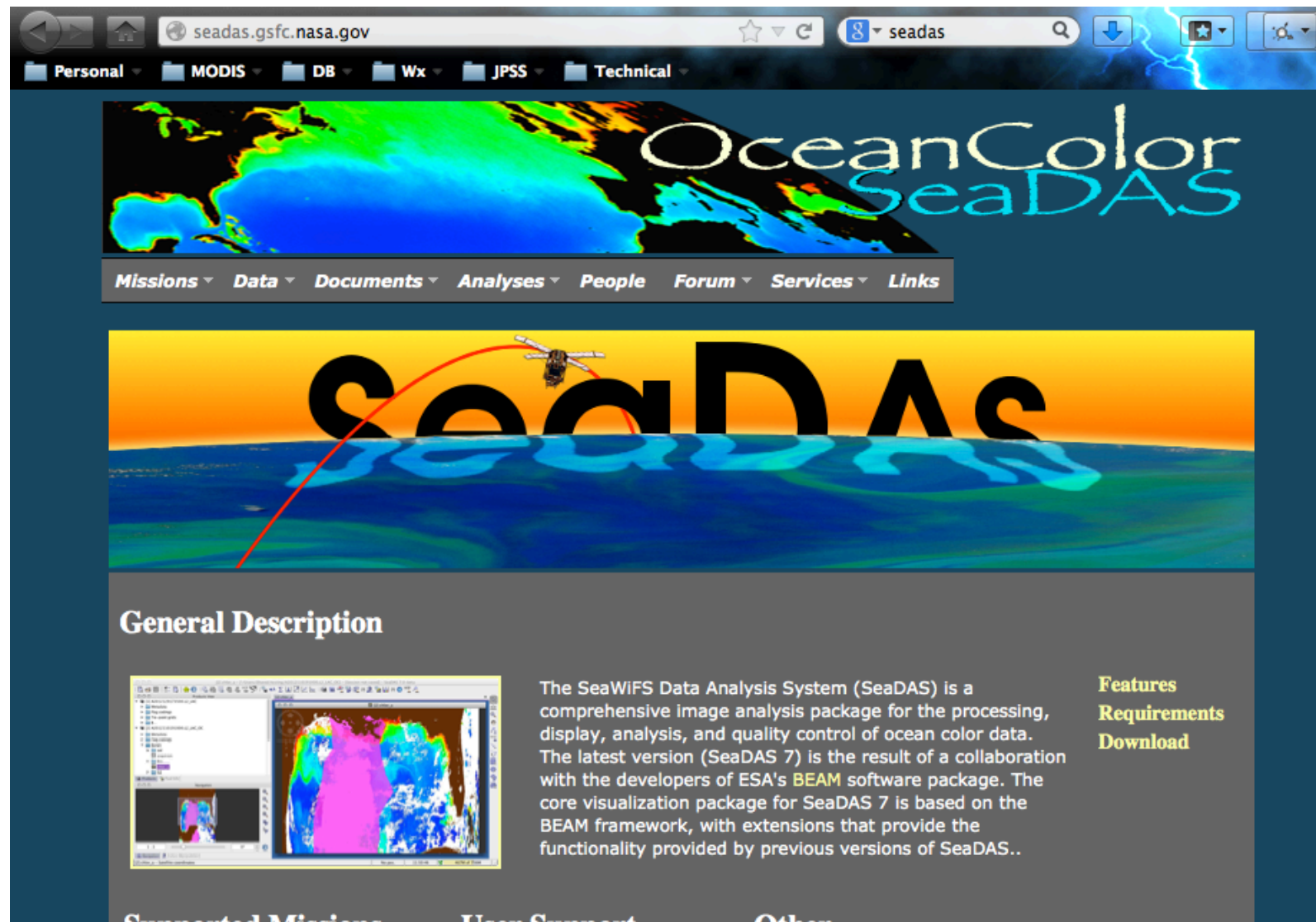


## MODIS Ocean Standard Products

Geophysical Parameter Name	Description
nLw_412	Normalized water-leaving radiance at 412 nm
nLw_443	Normalized water-leaving radiance at 443 nm
nLw_488	Normalized water-leaving radiance at 488 nm
nLw_531	Normalized water-leaving radiance at 531 nm
nLw_551	Normalized water-leaving radiance at 551 nm
nLw_667	Normalized water-leaving radiance at 667 nm
Tau_869	Aerosol optical thickness at 869 nm
Eps_78	Epsilon of aerosol correction at 748 and 869 nm
Chlor_a	OC3 Chlorophyll a concentration
K490	Diffuse attenuation coefficient at 490nm
Angstrom_531	Angstrom coefficient, 531-869 nm
SST	Sea Surface Temperature: 11 micron
SST4	Sea Surface Temperature: 4 micron (night only)



# Ocean Color - SeaDAS



The screenshot shows the SeaDAS website in a web browser. The address bar displays 'seadas.gsfc.nasa.gov'. A navigation menu at the top includes 'Personal', 'MODIS', 'DB', 'Wx', 'JPSS', and 'Technical'. The main header features a satellite map of the Atlantic Ocean with the text 'OceanColor SeaDAS'. Below this is a secondary navigation bar with links: 'Missions', 'Data', 'Documents', 'Analyses', 'People', 'Forum', 'Services', and 'Links'. A large banner image shows a satellite in orbit over the ocean with the text 'SeaDAS' overlaid. The 'General Description' section contains a screenshot of the SeaDAS software interface, which displays a satellite image of a coastal area. To the right of the screenshot is a text block describing SeaDAS as a comprehensive image analysis package for ocean color data, mentioning its collaboration with ESA's BEAM software. To the right of the text are three links: 'Features', 'Requirements', and 'Download'. At the bottom of the page, there are three tabs: 'Supported Missions', 'User Support', and 'Other'.

seadas.gsfc.nasa.gov

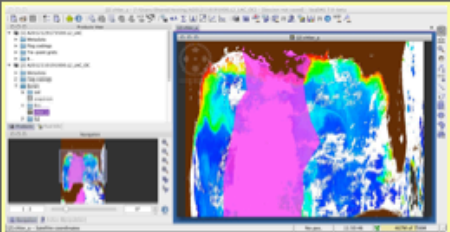
Personal MODIS DB Wx JPSS Technical

OceanColor SeaDAS

Missions Data Documents Analyses People Forum Services Links

SeaDAS

### General Description



The SeaWiFS Data Analysis System (SeaDAS) is a comprehensive image analysis package for the processing, display, analysis, and quality control of ocean color data. The latest version (SeaDAS 7) is the result of a collaboration with the developers of ESA's **BEAM** software package. The core visualization package for SeaDAS 7 is based on the BEAM framework, with extensions that provide the functionality provided by previous versions of SeaDAS..

**Features**  
**Requirements**  
**Download**

Supported Missions User Support Other



# MODIS Atmospheric Correction for Ocean Bands

Statement of the problem:

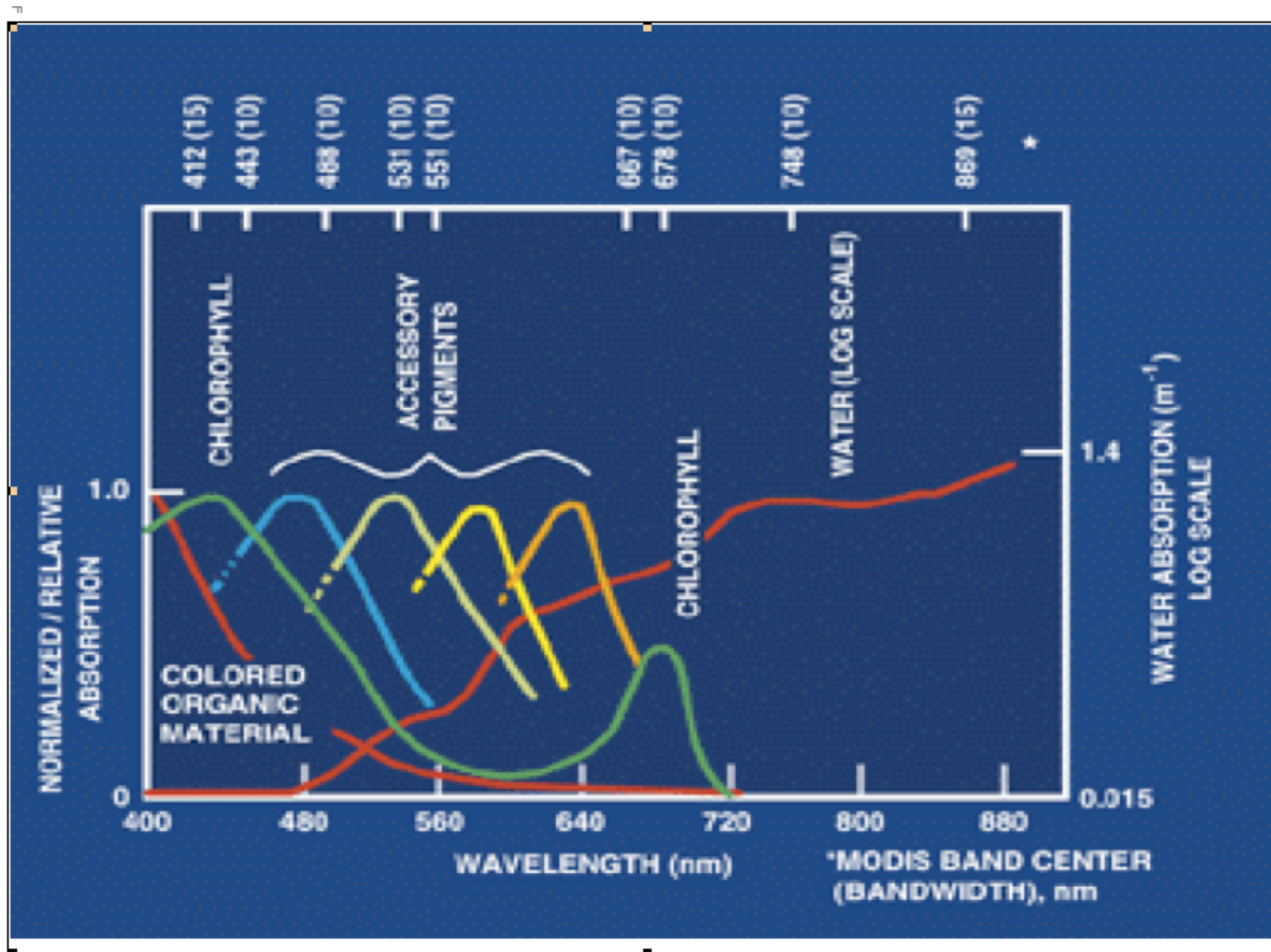
- Total radiance observed by the satellite is composed of 5-10% ocean signal and 90-95% atmosphere signal.
- The atmospheric and ocean surface scattering effects must be accurately modelled and removed.
- Desired parameter is normalized water leaving radiance (nLw) for MODIS bands 8, 9, 10, 11, 12, 13 (0.412, 0.443, 0.488, 0.531, 0.551, 0.667 microns)

Aerosol model selection:

- Assume zero (or negligible) water leaving radiance in the NIR bands (15 and 16; 0.750 and 0.865 microns); remainder is from aerosols.
- This is extrapolated to visible wavelengths using aerosol models.
- For case 1 waters, NIR bands are used to select aerosol model.
- Where this assumption is not valid, water-leaving radiance in NIR bands is estimated and removed prior to aerosol model selection.

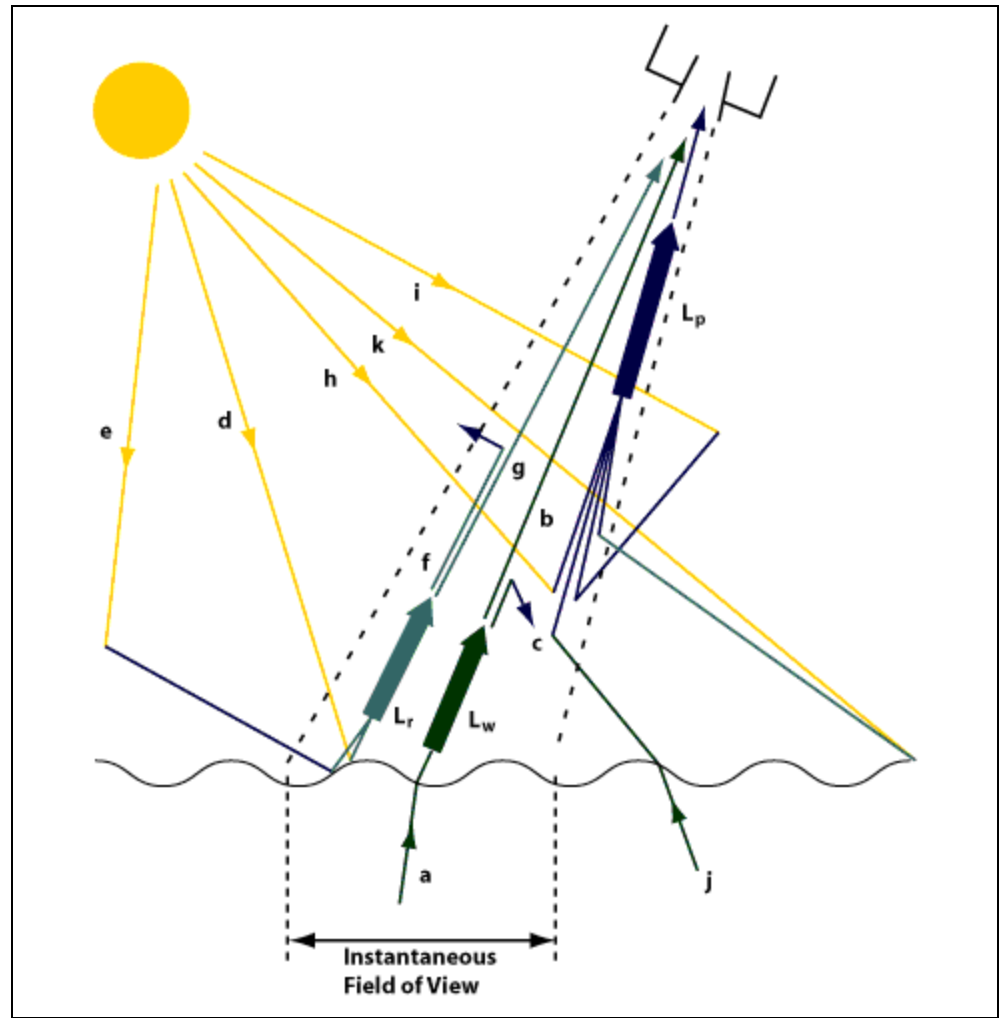
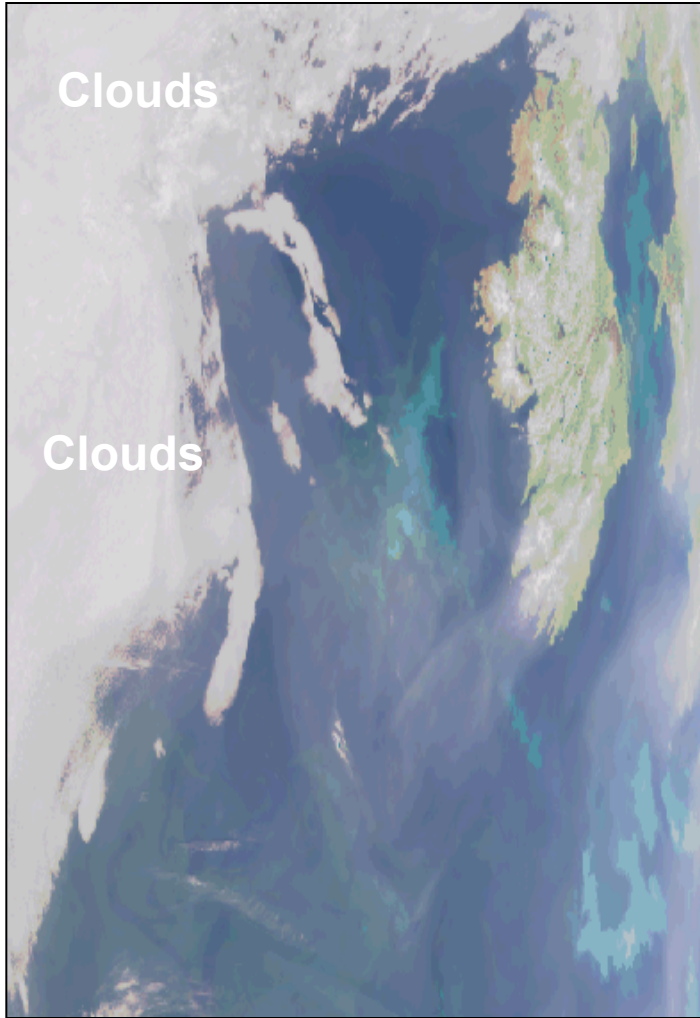


## Visible light absorption by water



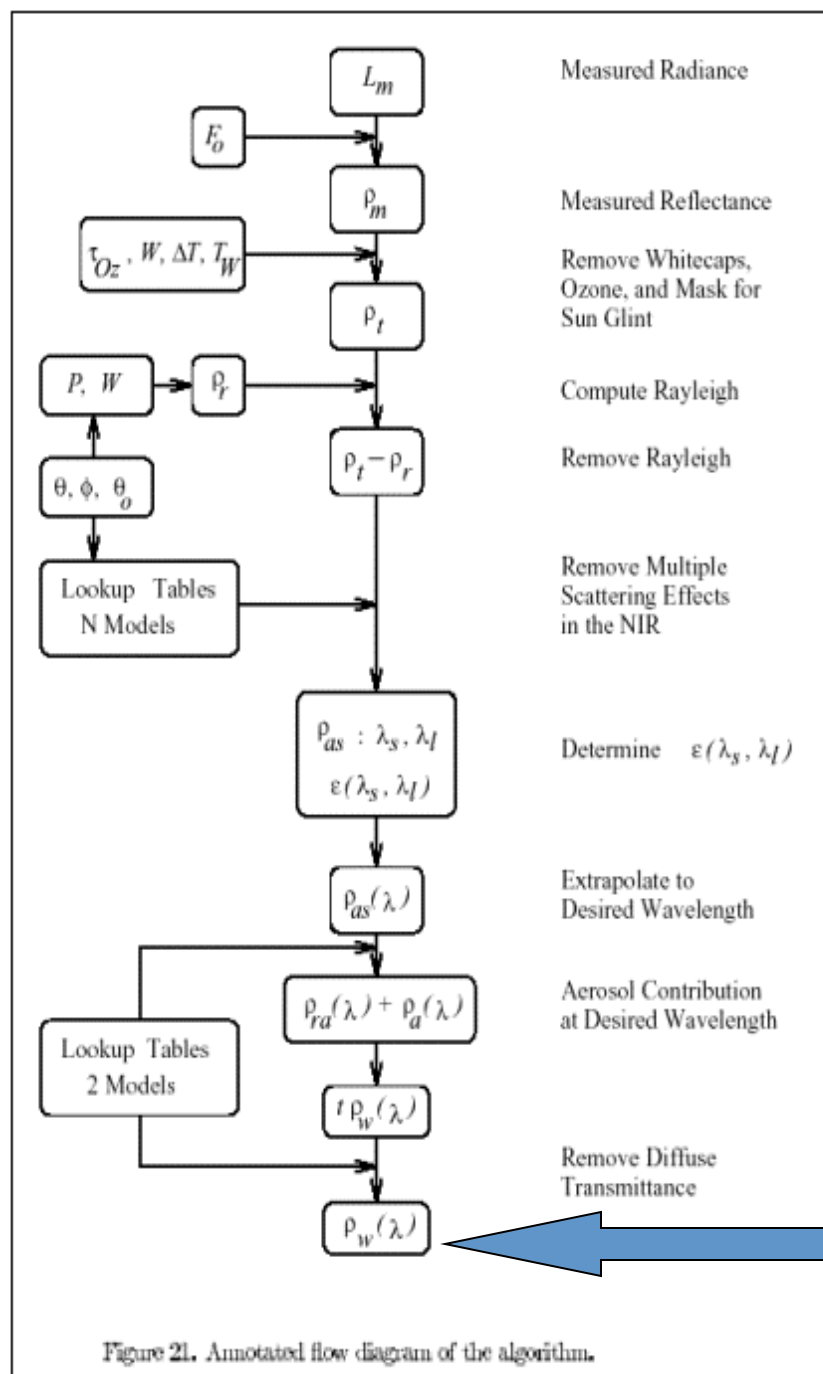


## Atmospheric correction is critical for ocean color retrievals



- $L_w$  is only 5-10% of signal reaching satellite: rest due to  $L_p$
- $L_p$  components: molecular (Rayleigh) & aerosols





$$\rho_t = \rho_r + (\rho_a + \rho_{ra}) + t\rho_{wc} + t\rho_g + t\rho_w$$

\*  $\rho_w$  is the quantity we wish to retrieve at each wavelength.

\*  $\rho_g$  is Sun glint, the direct + diffuse reflectance of the solar radiance from the sea surface. This effect for SeaWiFS is minimized by tilting the sensor. MODIS does not tilt and the sun glint must be removed, depends on vector winds and polarization.

\*  $\rho_{wc}$  is the contribution due to "white"-capping, estimated from statistical relationship with wind speed.

\*  $\rho_r$  is the contribution due to molecular (Rayleigh) scattering, which can be accurately modeled. MODIS requires accurate measurement of change in mirror reflectivity with angle of incidence, depends on polarization, winds, atmospheric pressure

\*  $\rho_a + \rho_{ra}$  is the contribution due to aerosol and Rayleigh-aerosol scattering, estimated in NIR from measured radiances and extrapolated to visible using aerosol models.

\*  $\rho_t$  is the total reflectance measured at the satellite

Reflectance at water surface



# MODIS Chlorophyll Algorithm (OC3)

Semi-analytical algorithm<sup>(1)</sup>

$$\text{Chl\_a} = 10^{(0.283 - 2.753 \cdot R + 1.457 \cdot R^2 + 0.659 \cdot R^3 - 1.403 \cdot R^4)}$$

where:

$$R = \log_{10}((R_{rs443} - R_{rs488}) / R_{rs551})$$

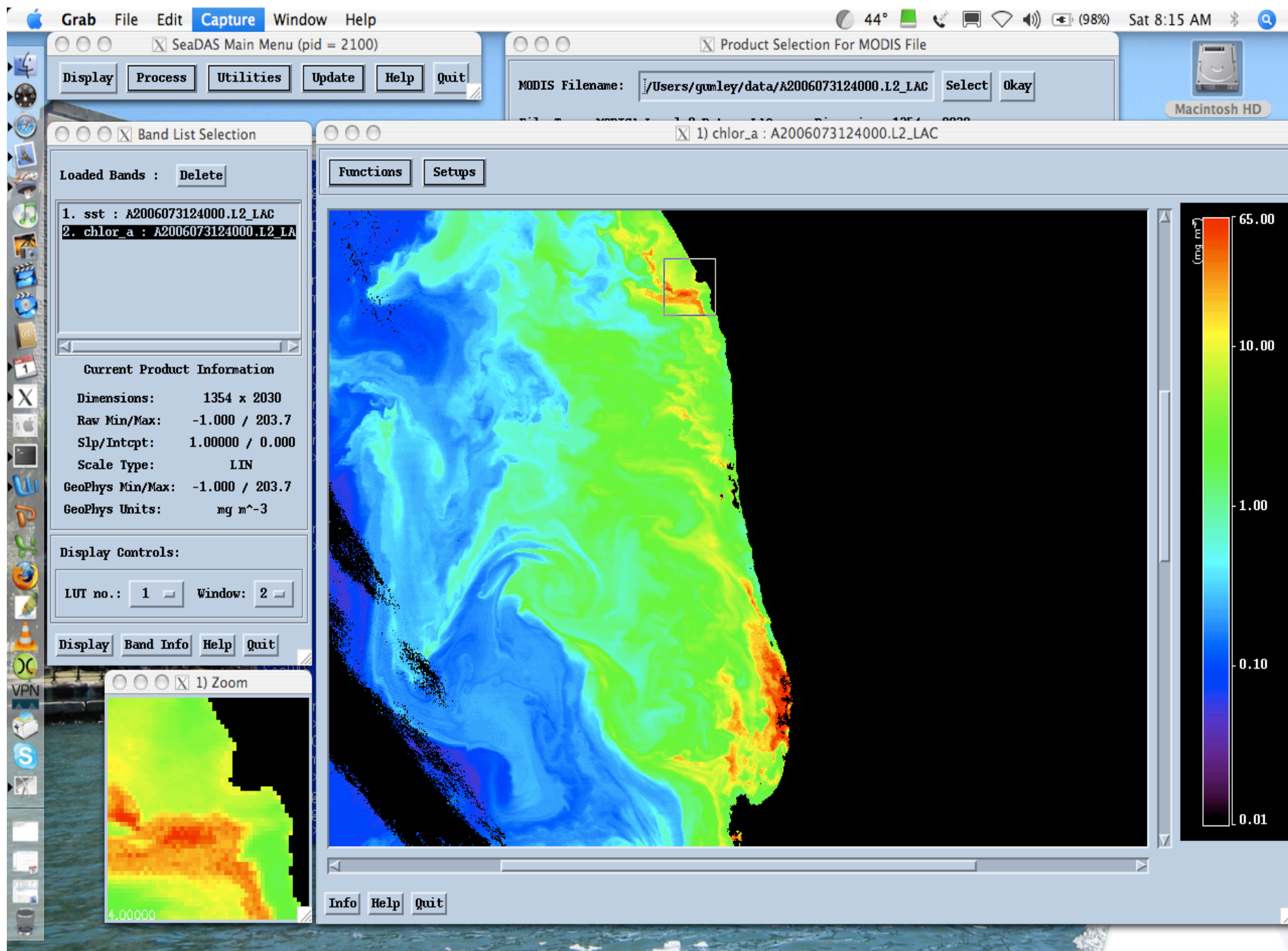
$R_{rs} = nLw / F_0$ ; remote sensing reflectance

$F_0$  = extraterrestrial solar irradiance

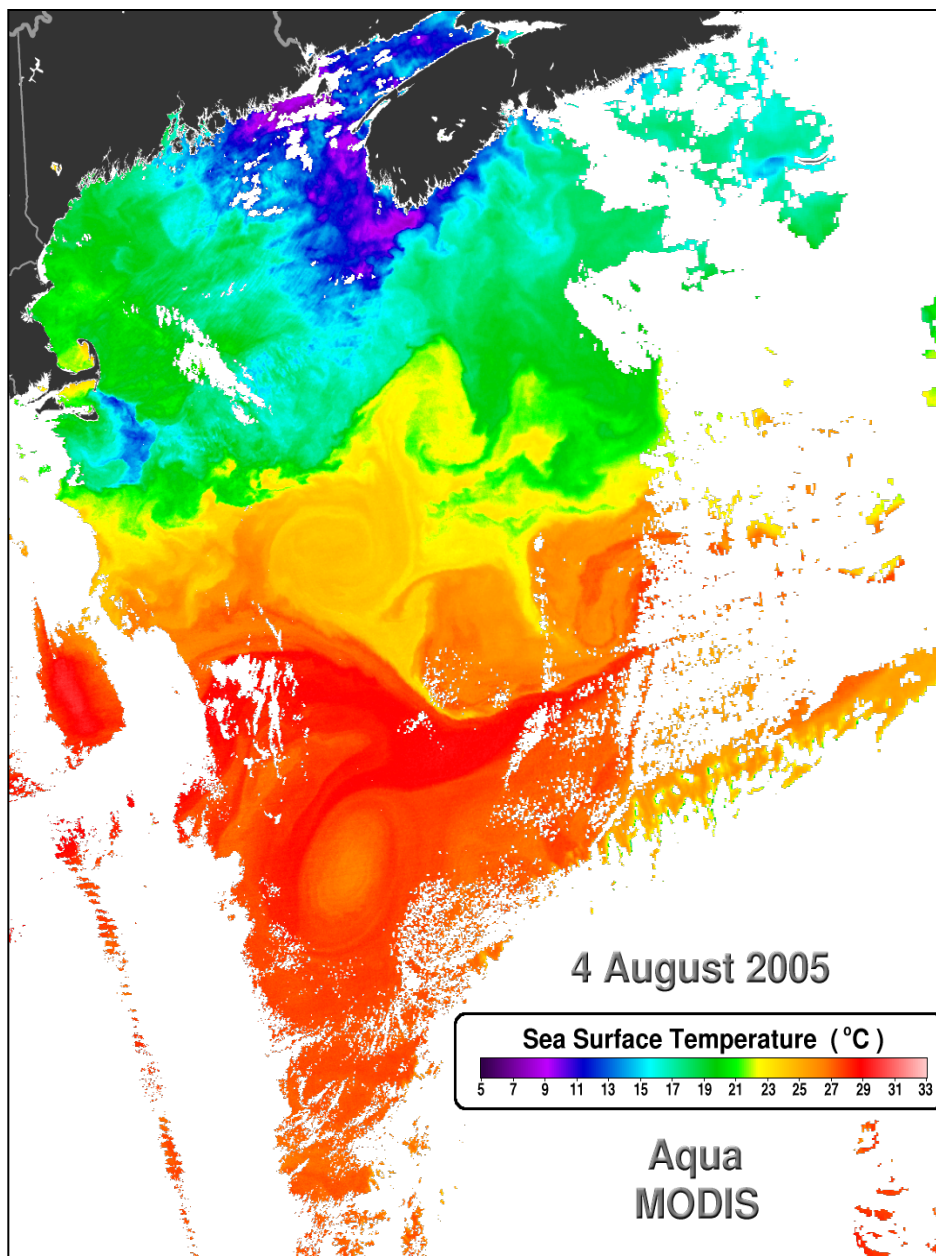
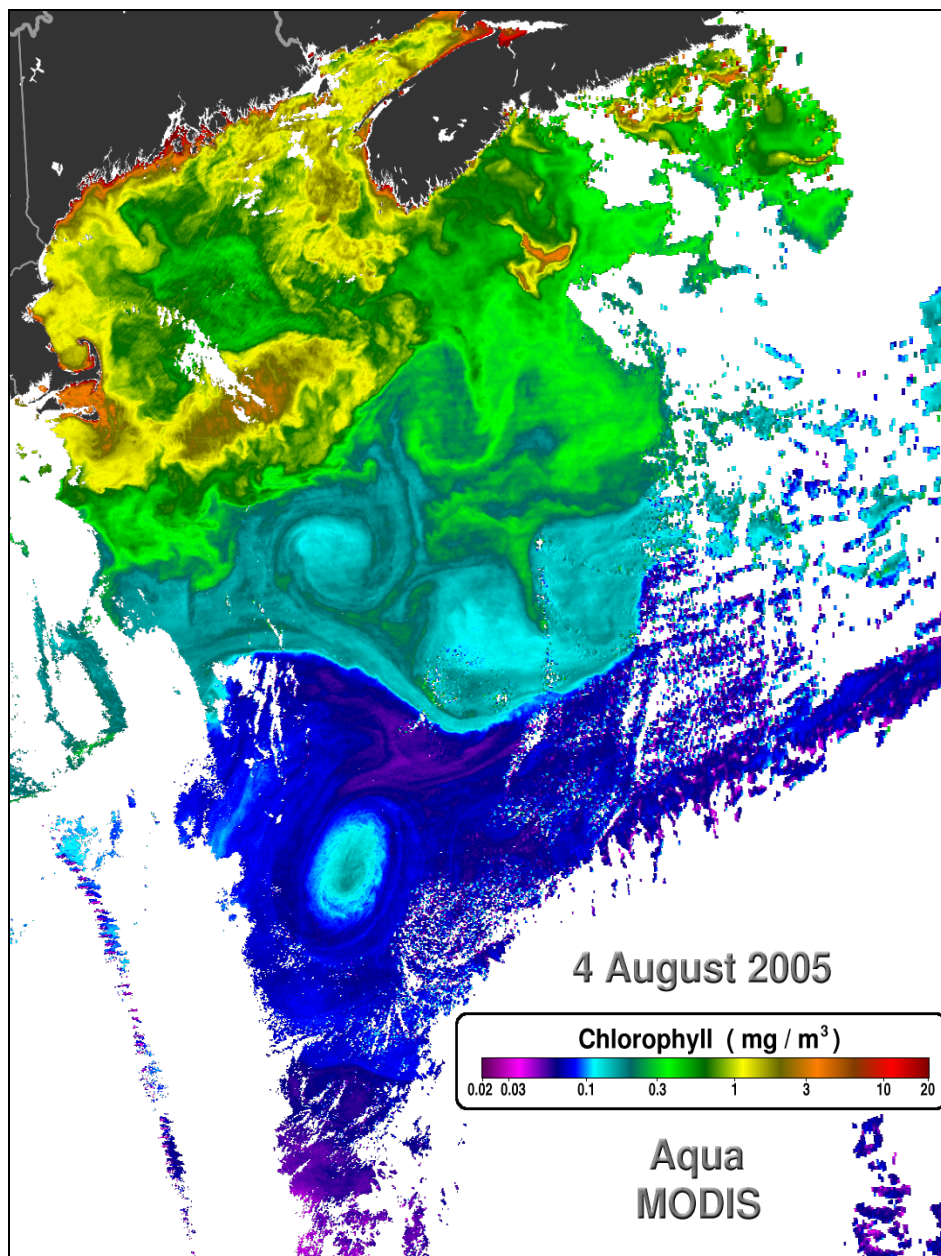
$nLw$  = water leaving radiance at 443, 488, 551

<sup>(1)</sup> Performance of the MODIS Semi-analytical Ocean Color Algorithm for Chlorophyll-a Carder, K.L.; Chen, F.R.; Cannizzaro, J.P.; Campbell, J.W.; Mitchell, B.G. Advances in Space Research. Vol. 33, no. 7, pp. 1152-1159. 2004









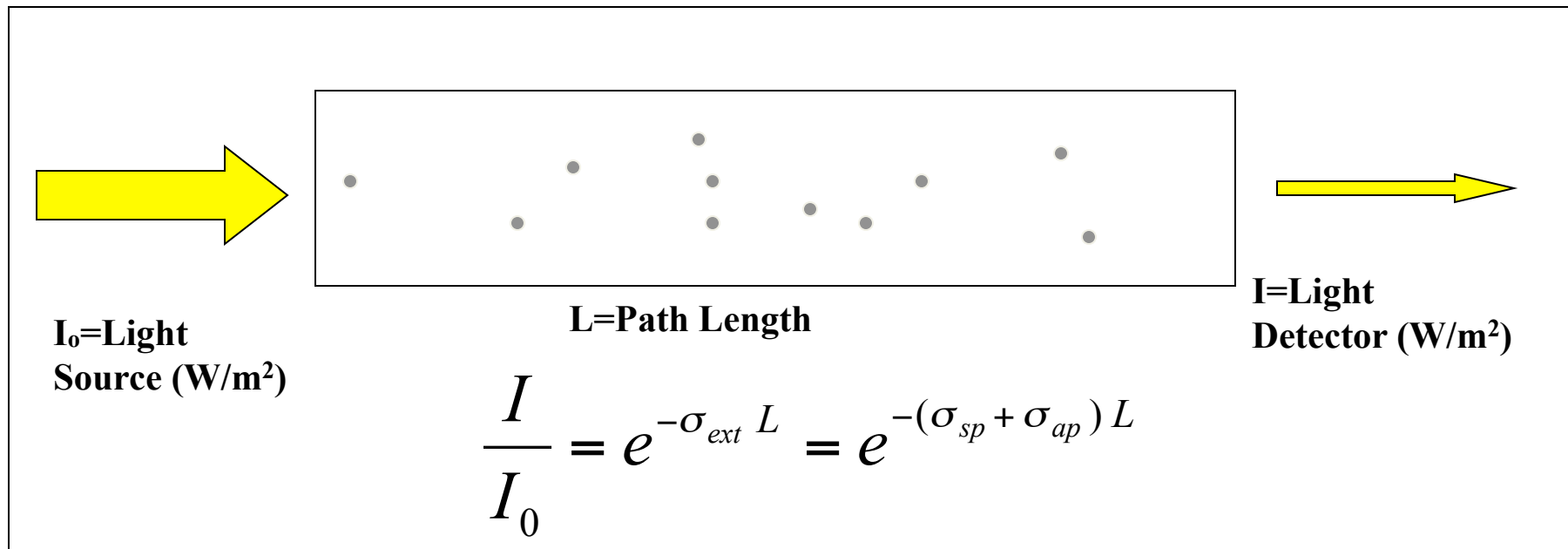


# Air Quality Applications

Aerosol Detection



# Scattering and Absorption of Light by Aerosols



$$\tau = (\sigma_{sp} + \sigma_{ap}) * L \quad \varpi = \sigma_{sp} / (\sigma_{sp} + \sigma_{ap})$$

The quantity  $L$  is called the density weighted path length.  $\sigma_{ext(\lambda)} L$  is a measure of the cumulative depletion that the beam of radiation has experienced as a result of its passage through the layer and is often called the optical depth  $\tau_\lambda$ .



## Wide Spectral Range makes land retrieval possible

- Mid-IR is used to observe the surface brightness
- Then aerosol is derived from estimated surface reflectance in the visible and actual reflectance

$$t_{0.66} \sim [r_{0.66}^* - 0.5r_{2.1}^*]$$

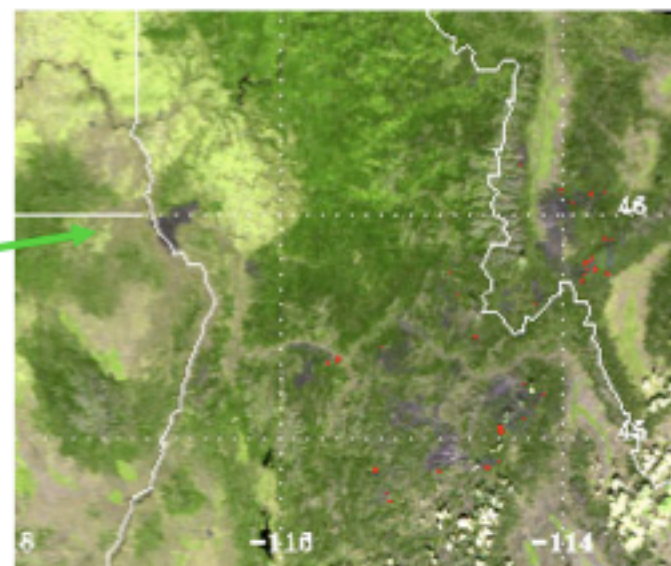
$$t_{0.47} \sim [r_{0.47}^* - 0.25r_{2.1}^*]$$

$\lambda(\mu\text{m})$

1.2

1.6

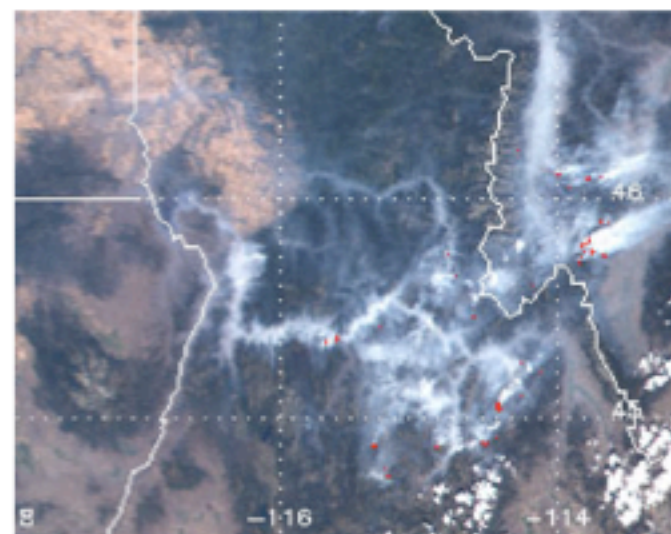
2.1



0.47

0.55

0.66



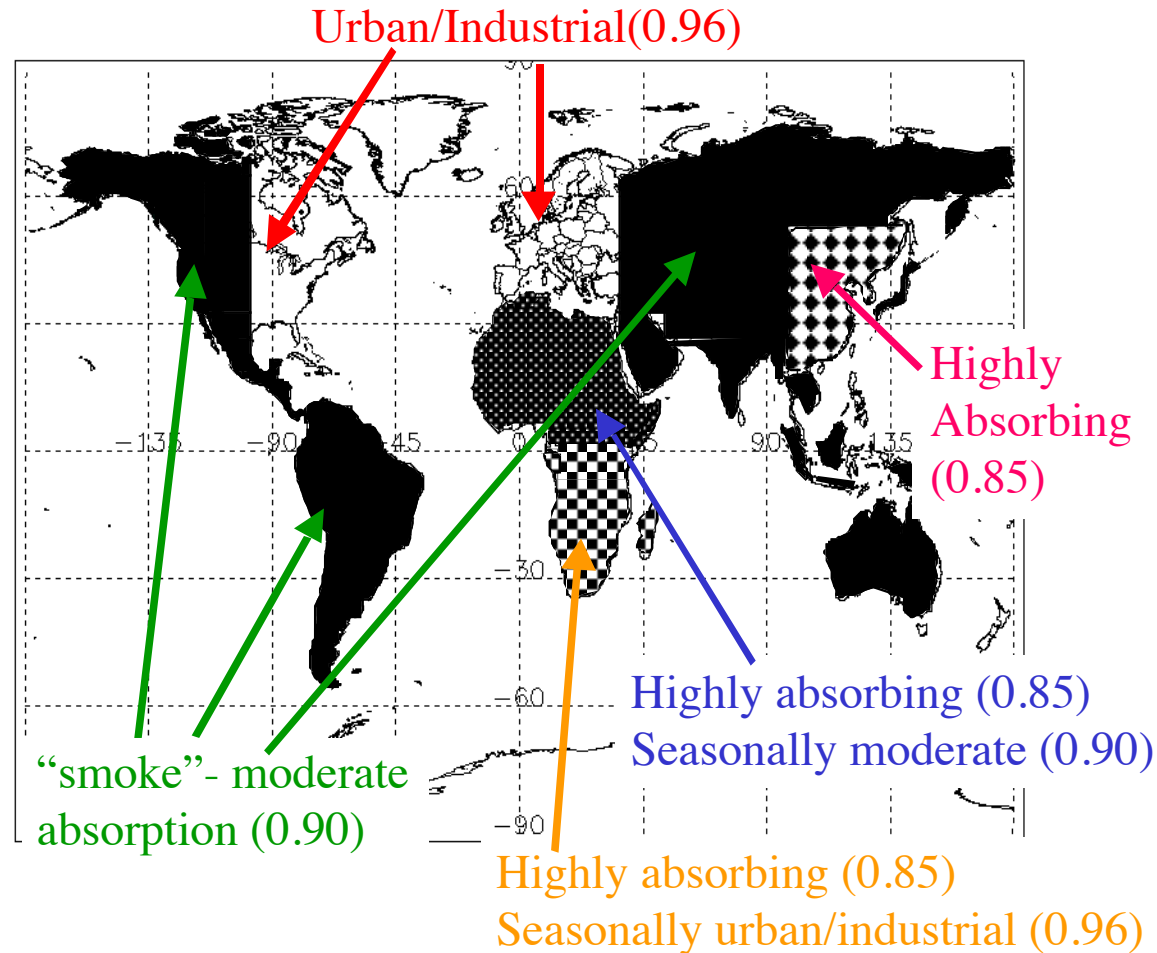


3 non-dust models

plus dust

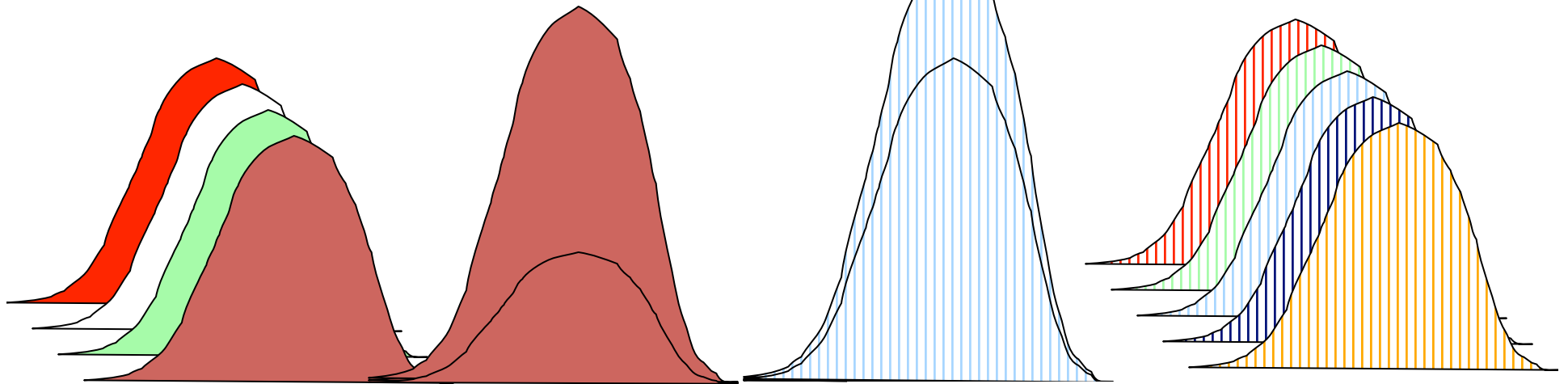
Set by geography and  
season

Models are dynamic  $f(\tau)$



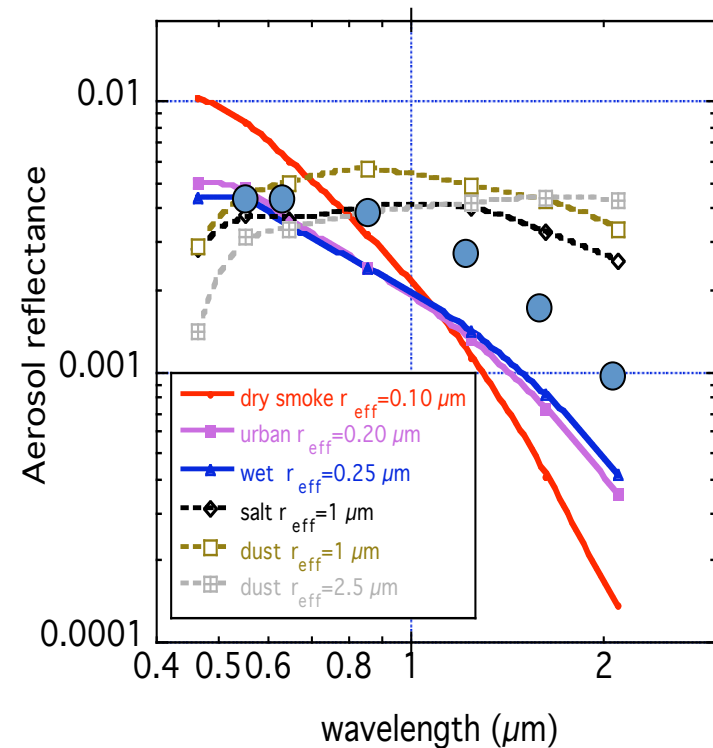


# The Ocean Algorithm



Choice of 4 fine modes  
and 5 coarse modes

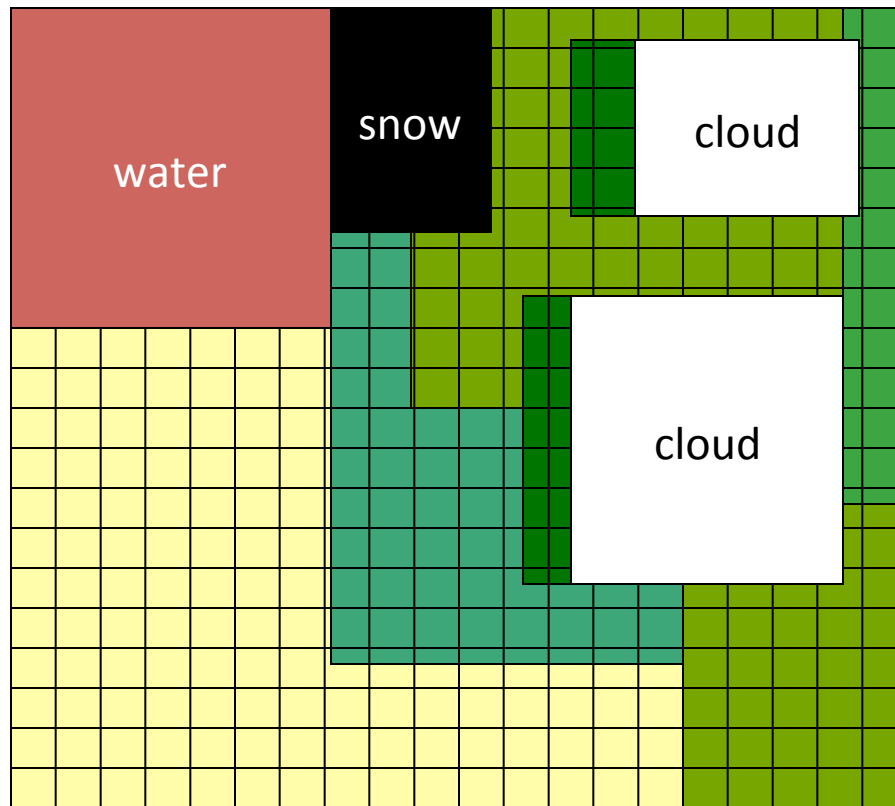
In order to minimize  
 $(\rho_{\text{meas}} - \rho_{\text{LUT}})$  over 6 wavelengths





# MODIS Over Land Algorithm

20 x 20 pixels at 500 m resolution  
(10 km at nadir)



← 10 km →

400 total  
- 56 water

---

344  
- 24 snow

---

320  
- 55 cloud

---

265  
- 116 "bright"

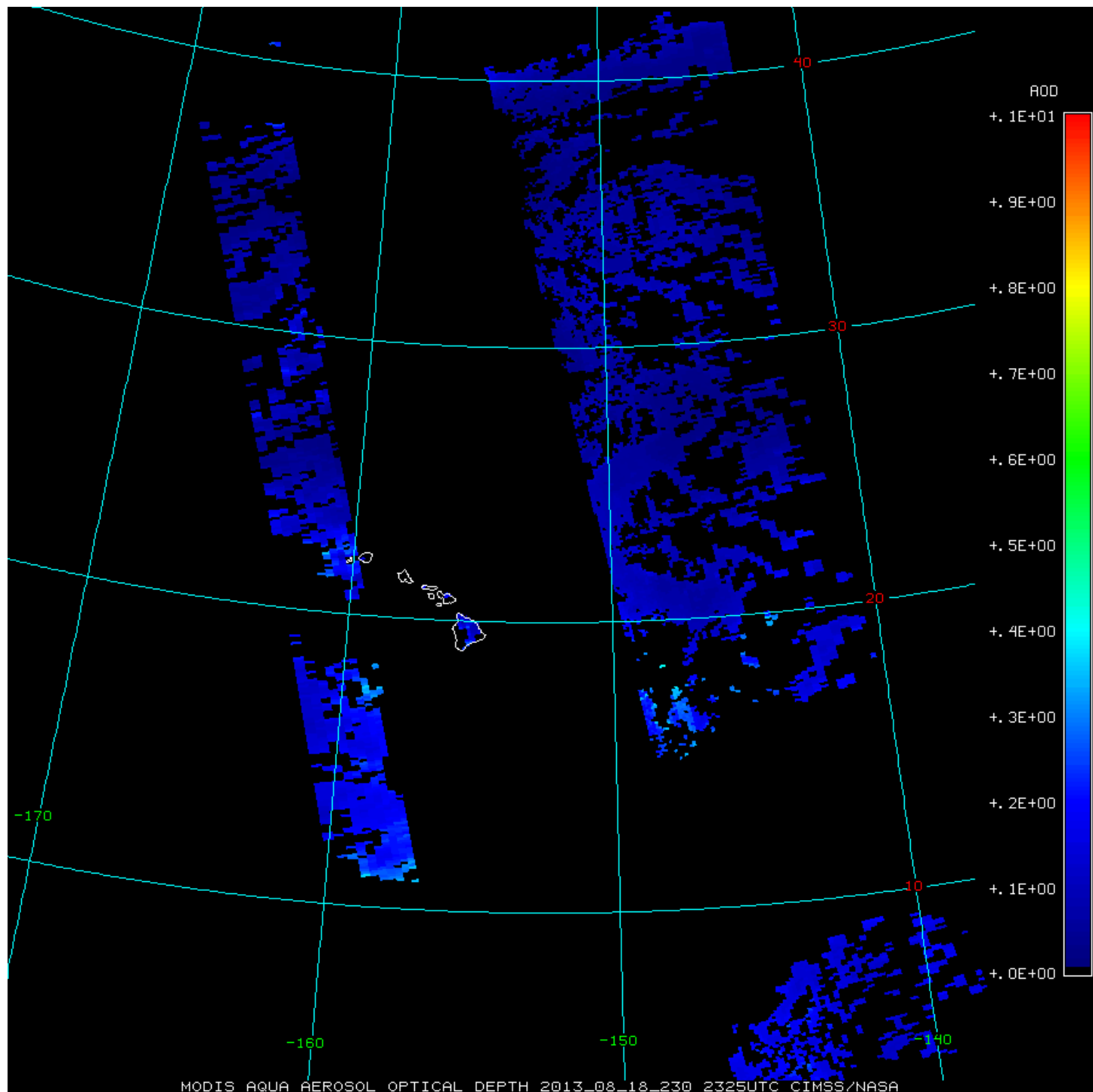
---

149 "good"

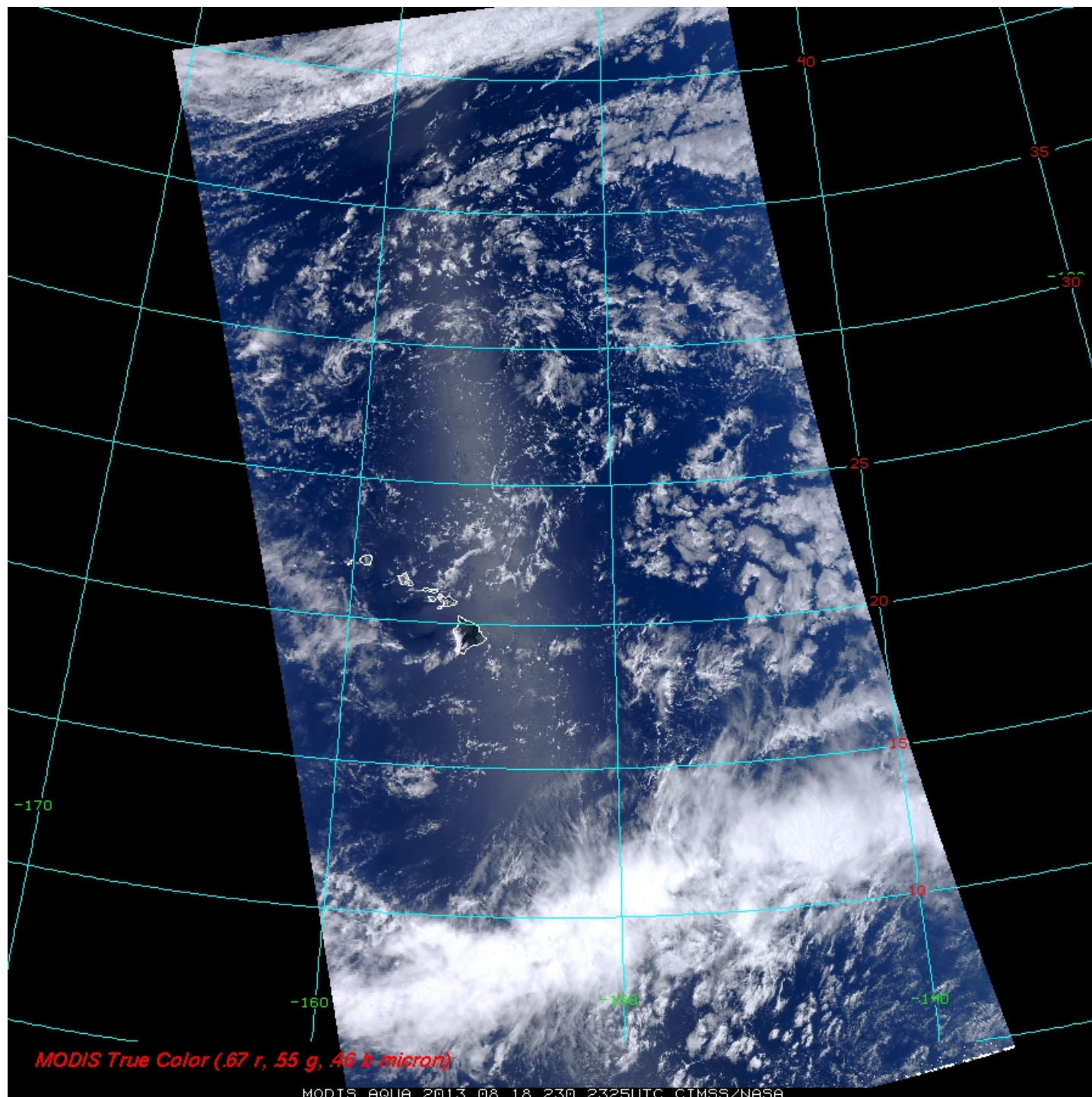
Discard brightest 50%  
and darkest 20% of the  
149 good pixels.

→ 44 pixels











# References

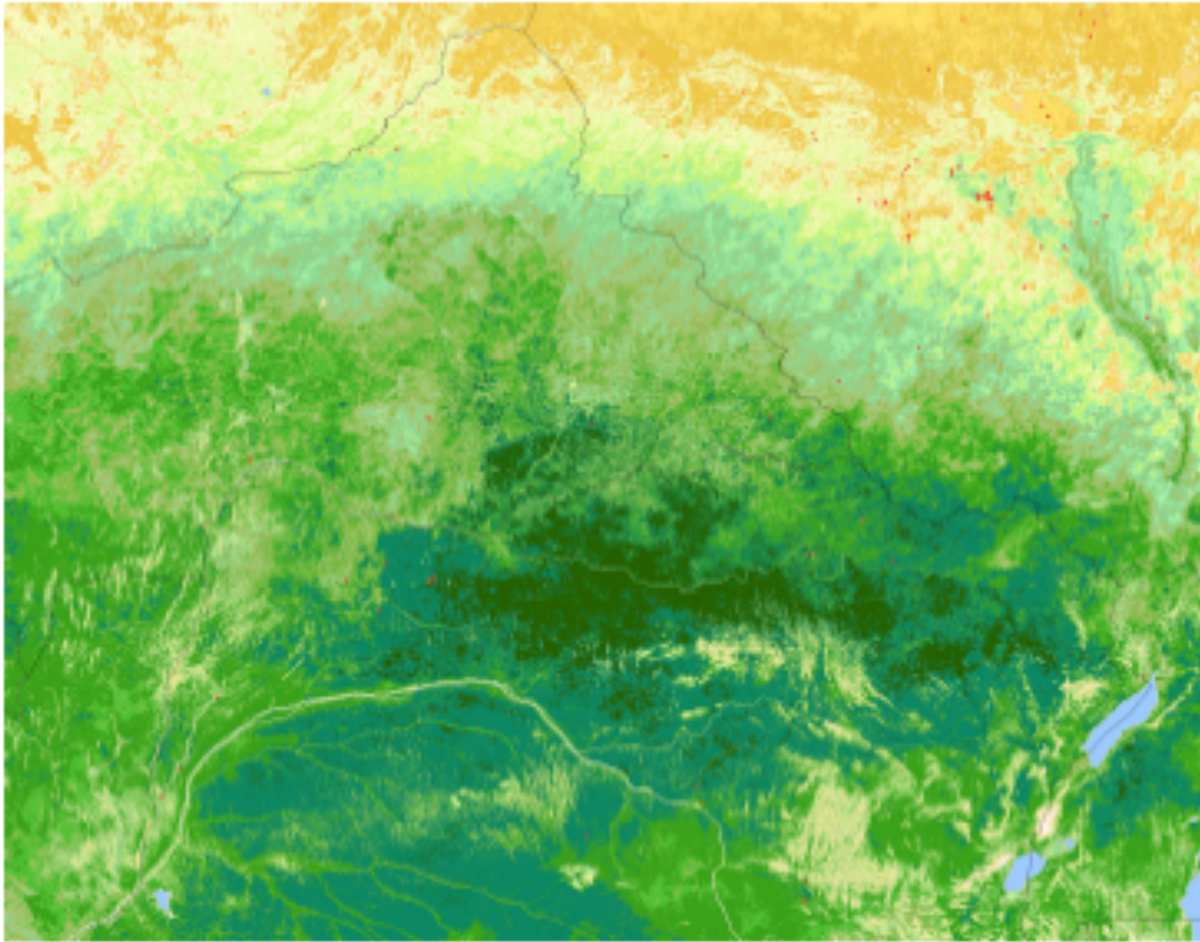
Levy, R. C., L. A. Remer, and O. Dubovik, 2007: Global aerosol optical properties and application to Moderate Resolution Imaging Spectroradiometer aerosol retrieval over land. J. Geophys. Res., 112, D13210

Levy, R. C., L. Remer, S. Mattoo, E. Vermote, and Y. J. Kaufman, 2007: Second-generation algorithm for retrieving aerosol properties over land from MODIS spectral reflectance. J. Geophys. Res., 112, D13211, 22 pages.

Remer, L. A., Y. J. Kaufman, D. Tanre, S. Mattoo, D. A. Chu, J. V. Martins, R-R. Li, C. Ichoku, R. C. Levy, R. G. Kleidman, T. F. Eck, E. Vermote, & B. N. Holben, 2004: The MODIS Aerosol Algorithm, Products and Validation. Journal of Atmospheric Sciences, 64, 4, 947-973.



# Vegetation Index

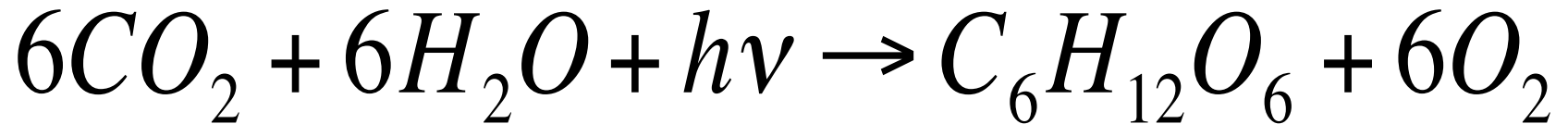


Normalized Difference Vegetation Index (NDVI) image of Central Africa  
<http://rapidfire.sci.gsfc.nasa.gov/>



# Photo-Chemistry

- Light may be absorbed and participate (drive) a chemical reaction. Example: Photosynthesis in plants



- Only certain wavelengths are absorbed by some participant(s) in the reaction
- Some structure must be present to allow the reaction to occur –Chlorophyll
- Combination of chemical and structural properties of plants

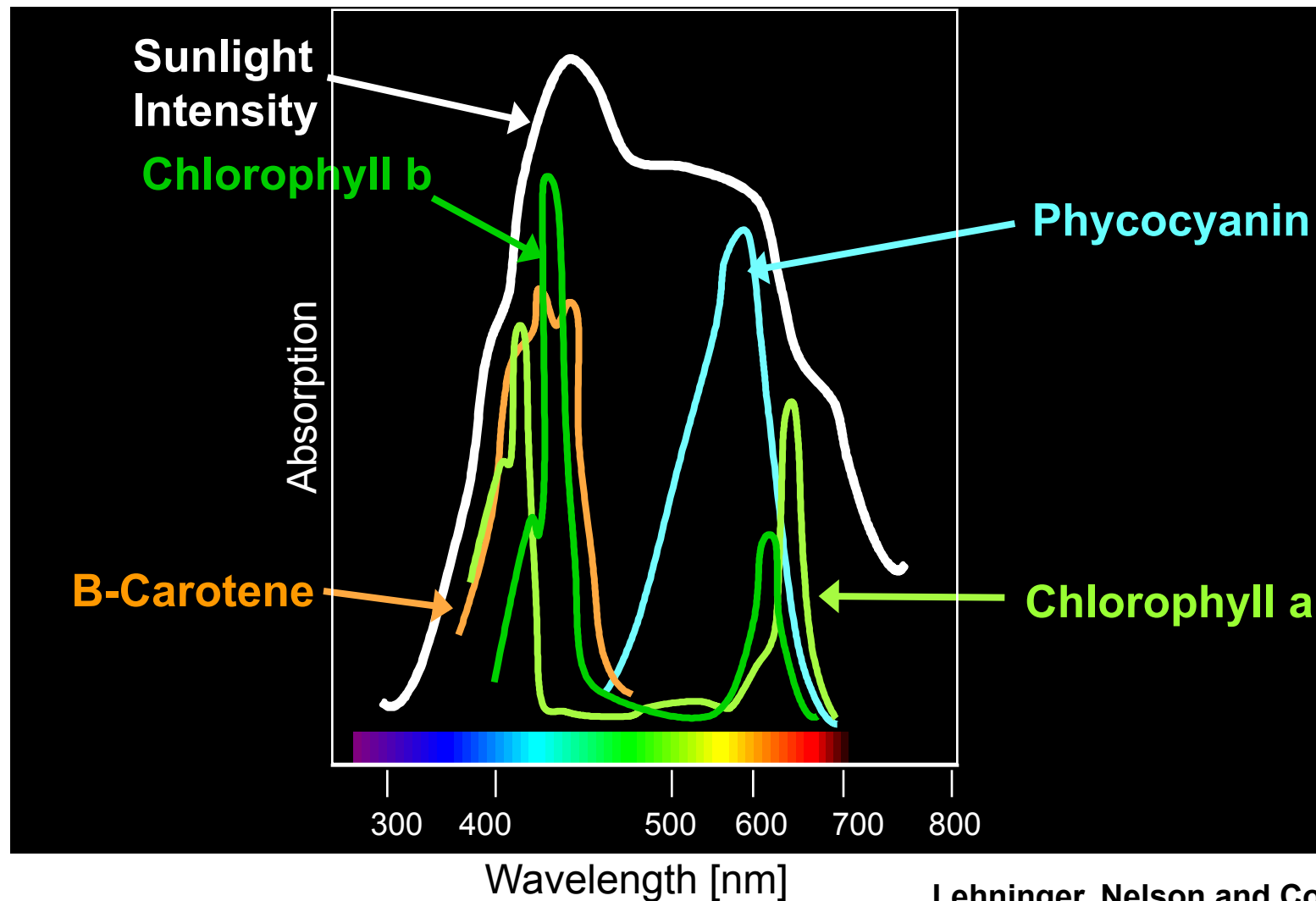


# Primary and secondary absorbers in plants

- Primary
  - Chlorophyll-a
  - Chlorophyll-b
- Secondary
  - Carotenoids
  - Phycobilins
  - Anthocyanins



# Absorption of Visible Light by Photo-pigments



Lehninger, Nelson and Cox



# Theoretical description

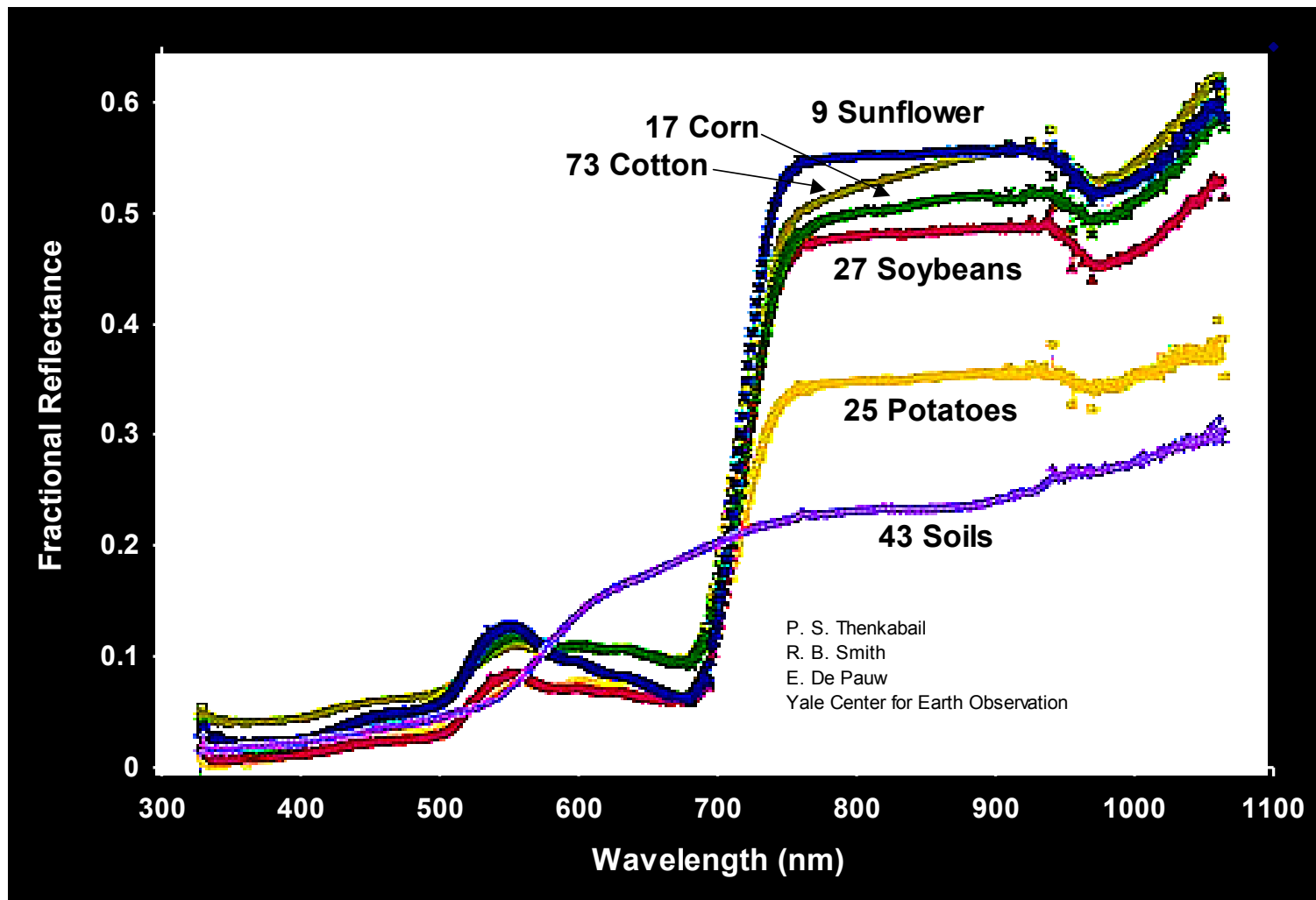
VISIBLE radiation is highly absorbed by vegetation in the red (0.68 micron) and in the blue (0.47 micron). The absorption is mainly due to photosynthetically active pigments

NIR radiation is reflected and transmitted with very little absorption by vegetation

Contrast between RED and NIR responses is correlated to vegetation amount



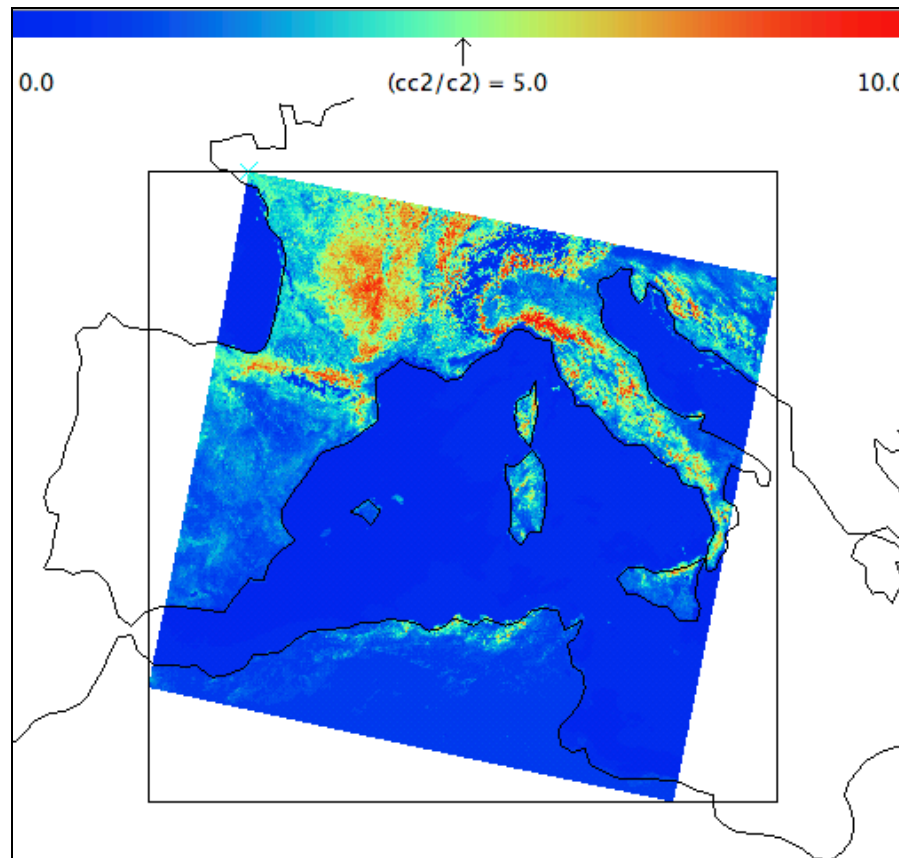
# Soil and crop reflectance





# Simple Ratio (SR)

- It was the first index to be used (Jordan, 1969)
- Defined as the ratio  $X_{\text{nir}}/X_{\text{red}}$
- For densely vegetated areas  $X_{\text{red}}$  tends to 0 and SR increases without bounds





# Normalized Difference Vegetation Index (NDVI)

Defined as the ratio

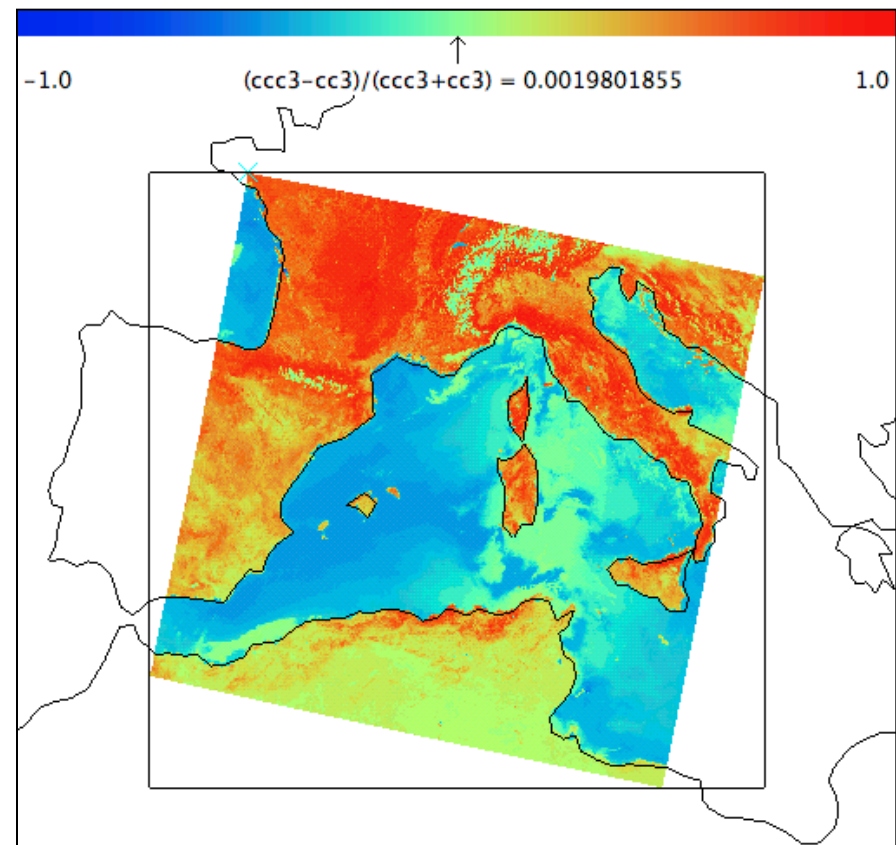
$$(X_{\text{nir}} - X_{\text{red}})/(X_{\text{nir}} + X_{\text{red}})$$

## Correlated with:

Plant Biomass	Crop Yield
Plant Nitrogen	Plant Chlorophyll
Water Stress	Plant Diseases
Insect Damage	

## Applications:

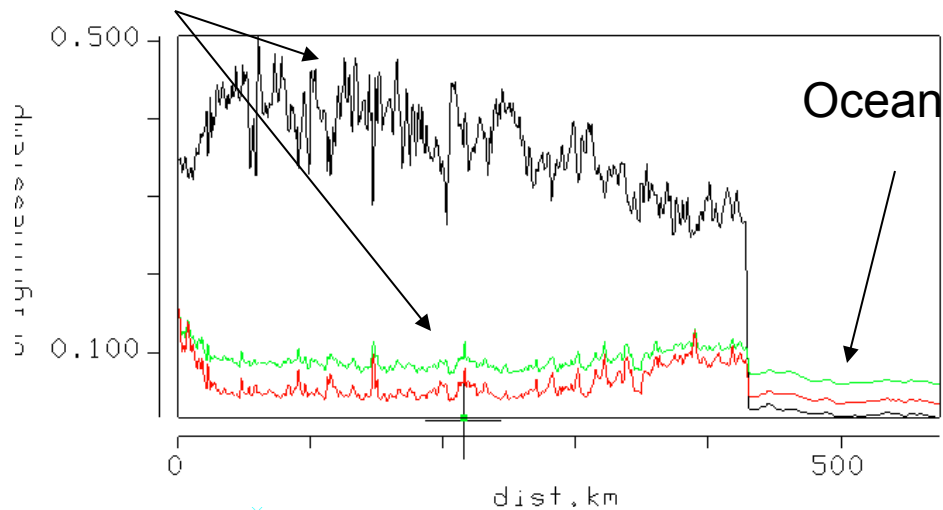
Vegetation Monitoring	Agricultural Activities
Drought studies	Landcover Change
Public Health Issues (mosquitos)	Climate Change Detection
Net Primary Production	Carbon Balance



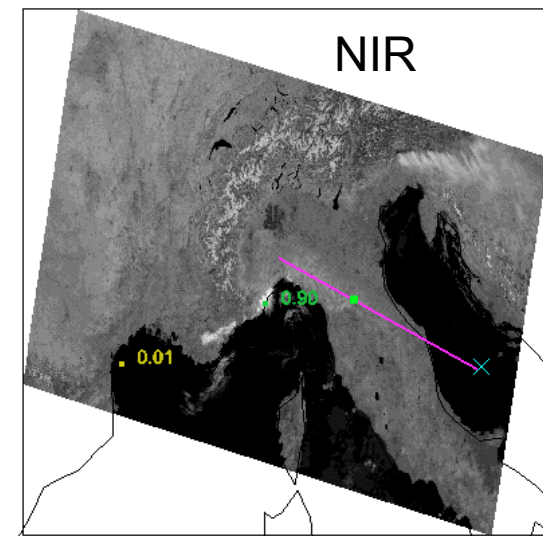
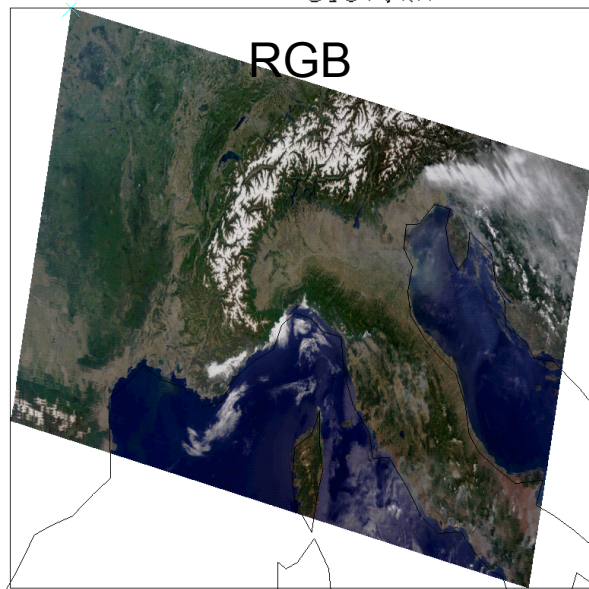


# NIR and VIS over Vegetation and Ocean

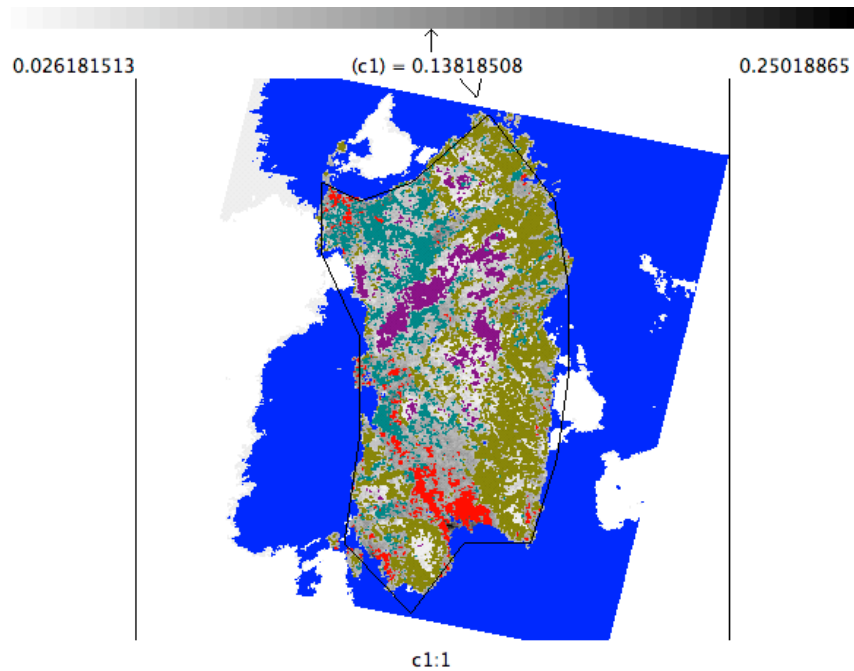
## Vegetation



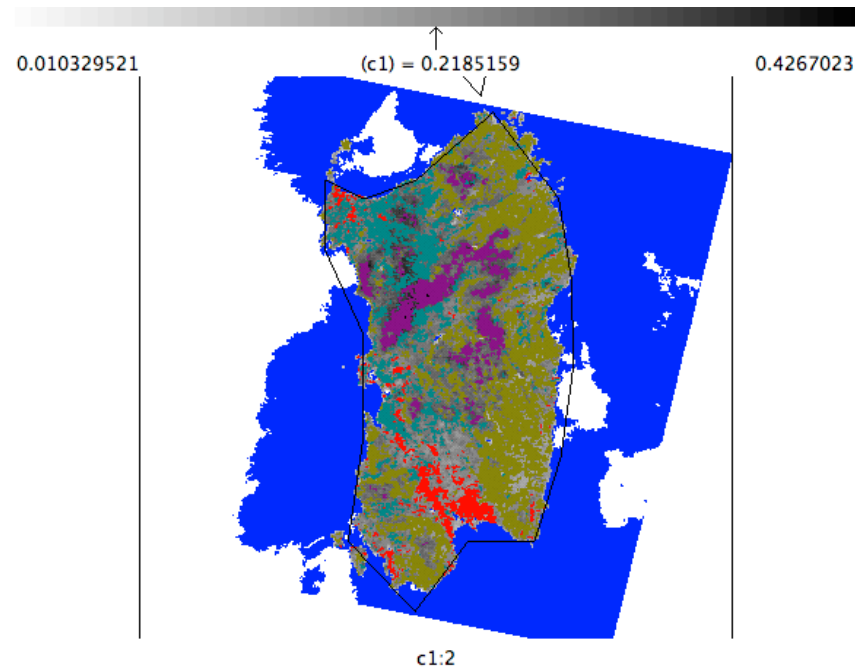
NIR (.86 micron)  
Green (.55 micron)  
Red (0.68 micron)



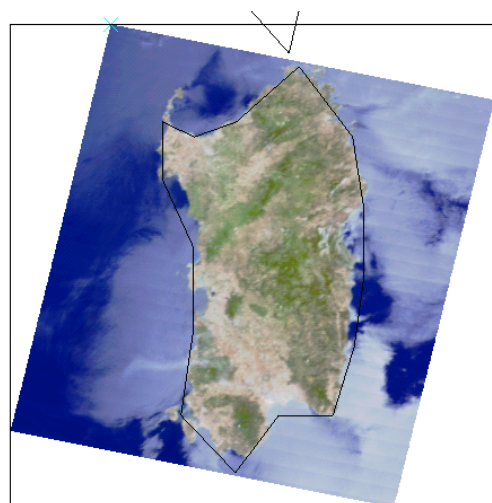




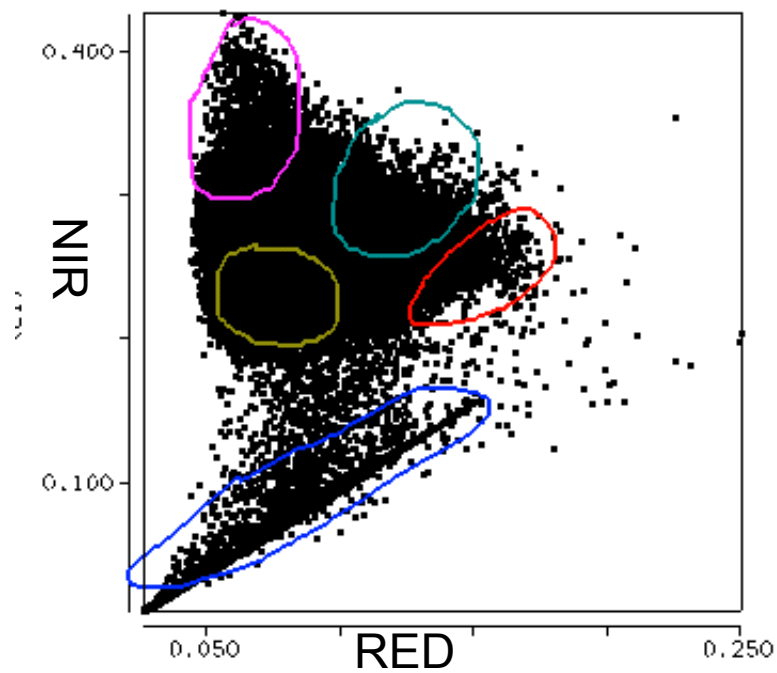
RED



NIR



True Color



Dense Vegetation

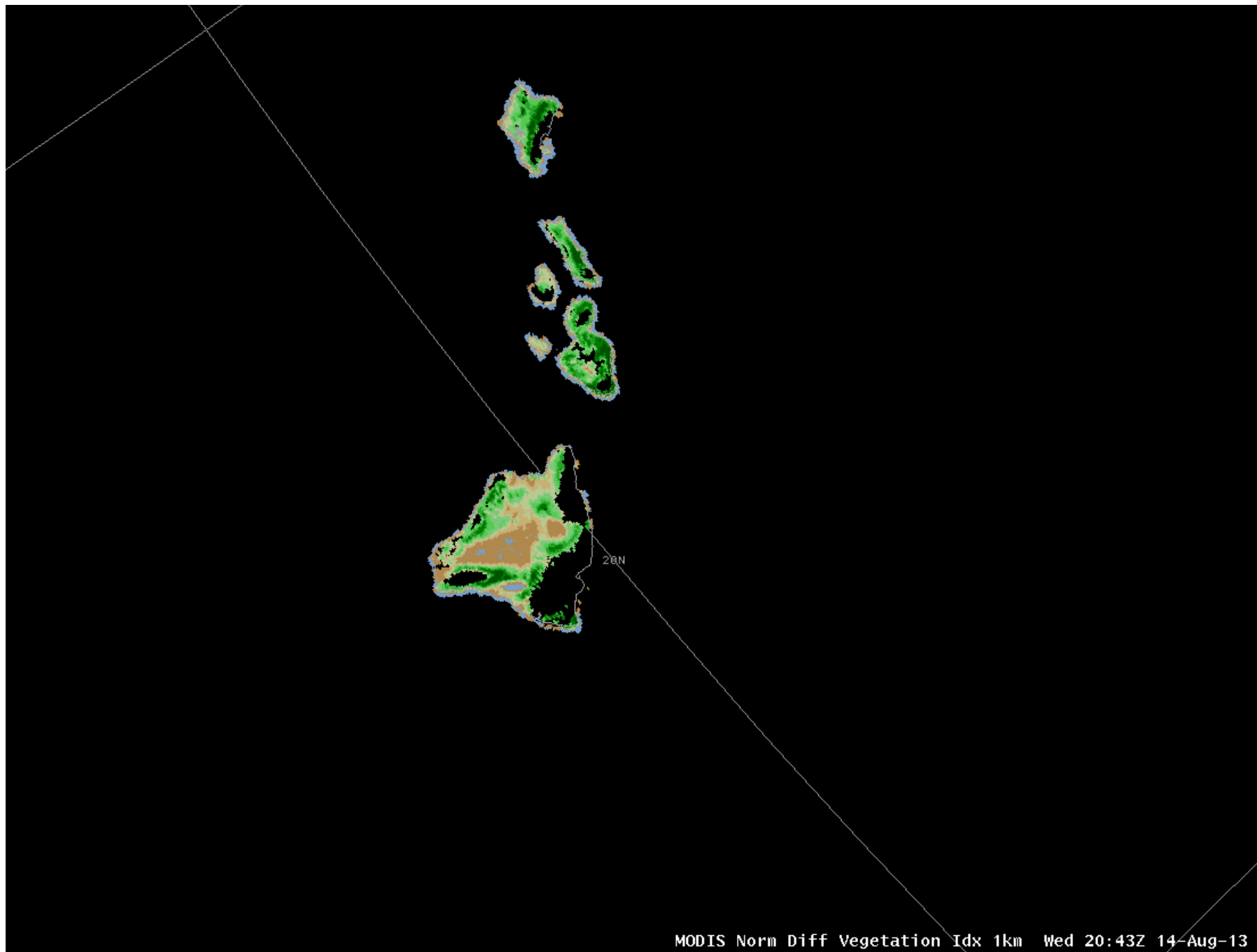
Less Vegetation

Less Vegetation

Barren Soil

Ocean







# Using MODIS Sun Glint Patterns

- What is sun glint?
- Application
  - Identifying regions of calm waters
  - Relationship of calm waters and sea surface temperatures





“Mirror” reflection of sunlight off calm water.

Sun Glint Ellipse Defined by:  $\theta_r < 36$

$$\cos \theta_r = \sin \theta_v \cos \theta_s \cos \Delta\Phi + \sin \theta_v \cos \theta_s$$

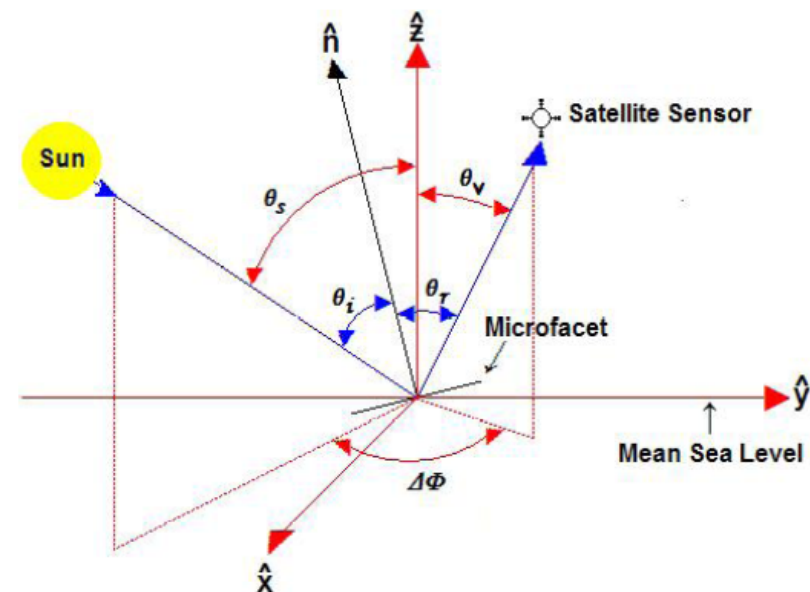
Where  $\theta_v$  = Viewing Zenith Angle

$\theta_s$  = Solar Zenith Angle

$\Delta\Phi$  = Relative Angle –  
difference between the Solar and  
Viewing azimuth angles.

## Sun Glint

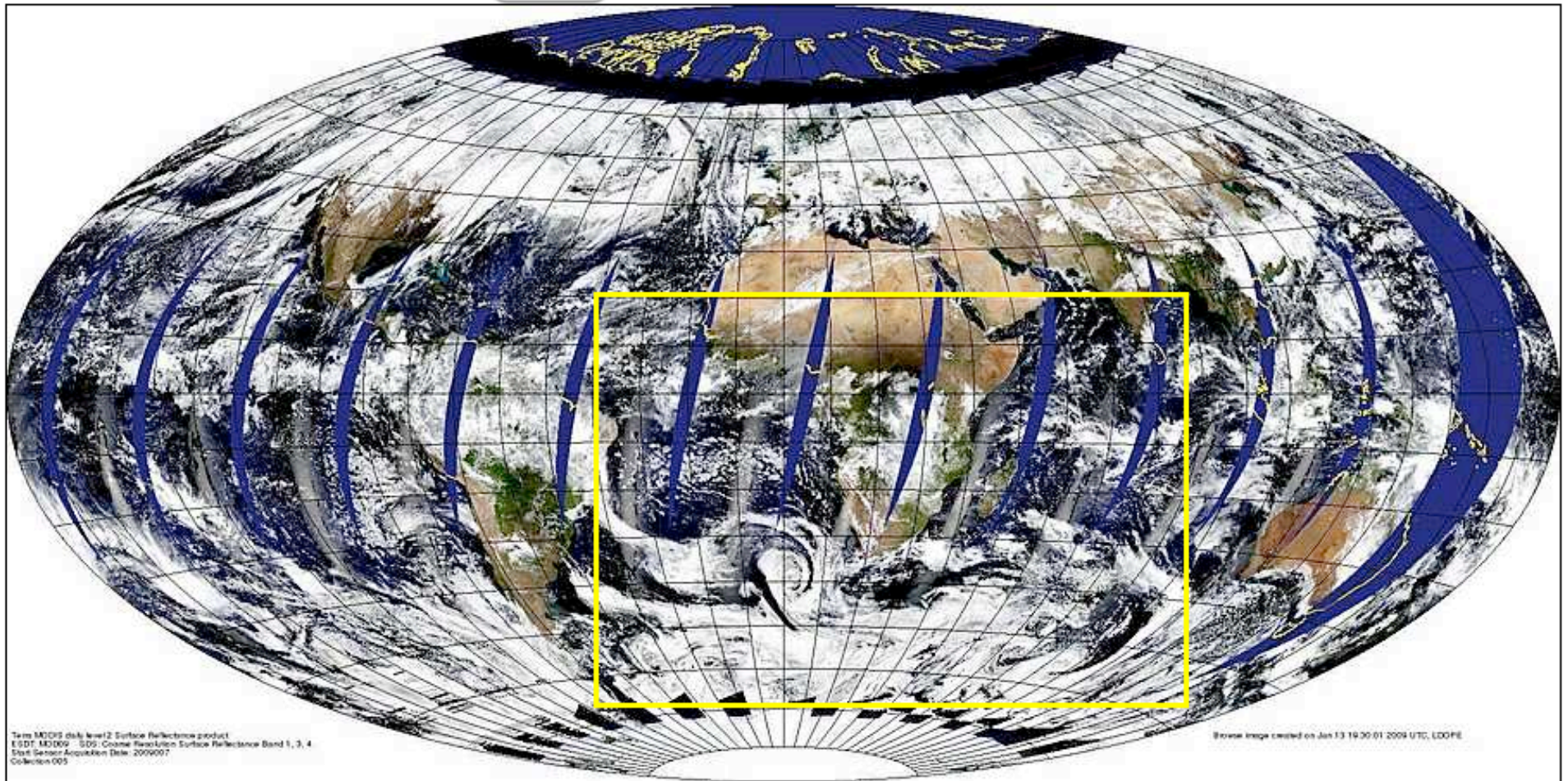
Simple example where your  
eye is the sensor





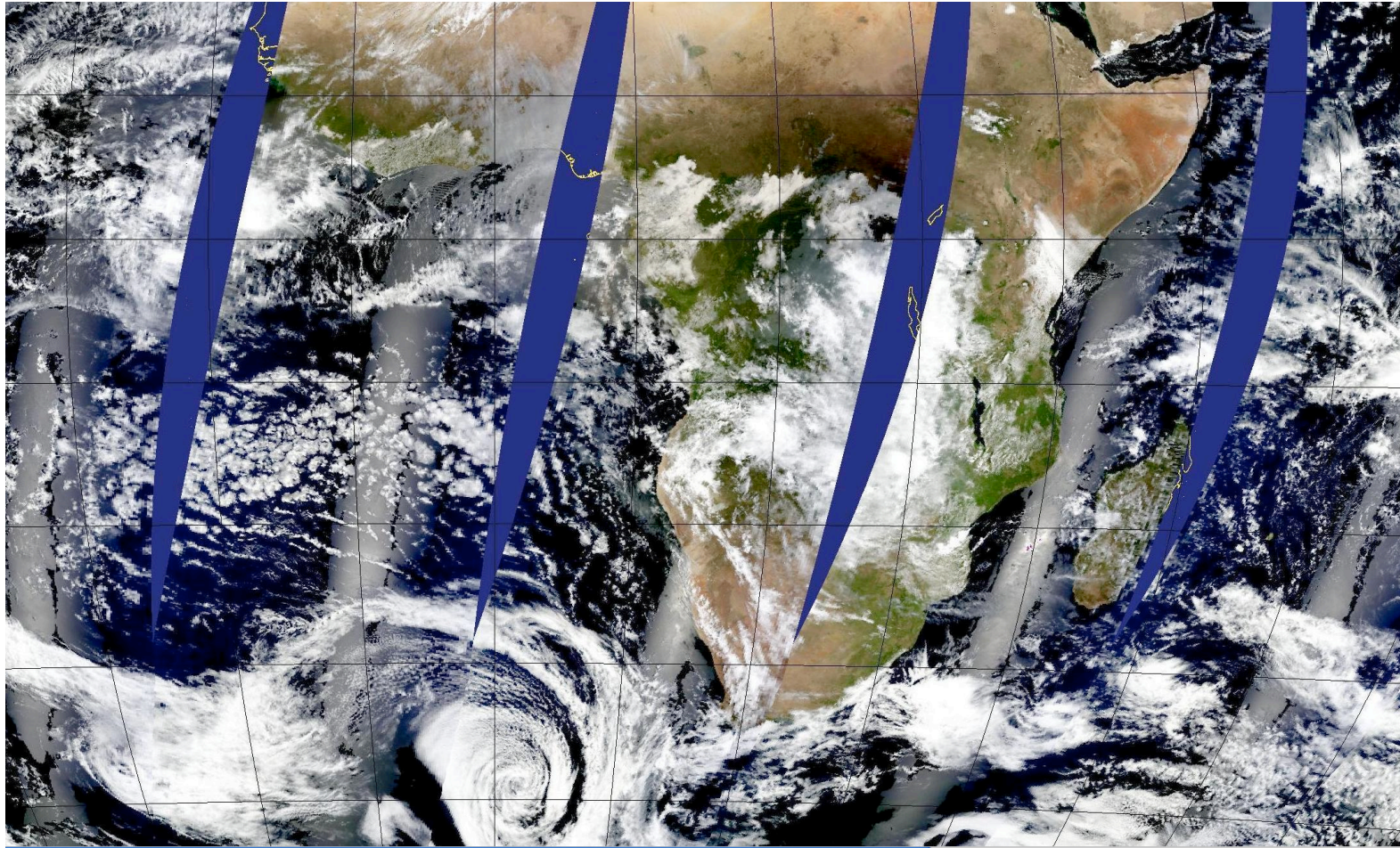
# Aqua MODIS Sun Glint Example

7 January 2009



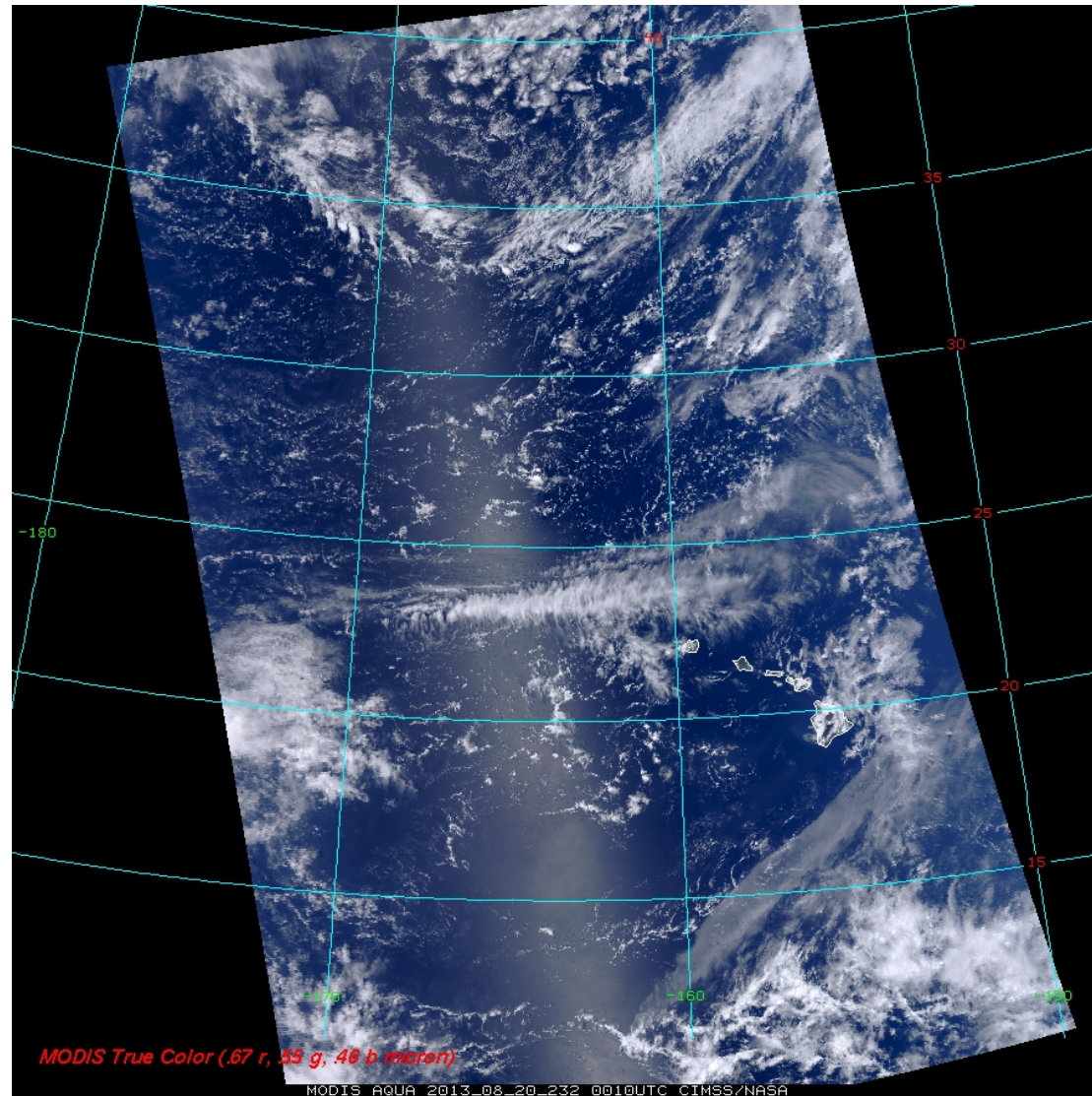


# Sun Glint Patterns





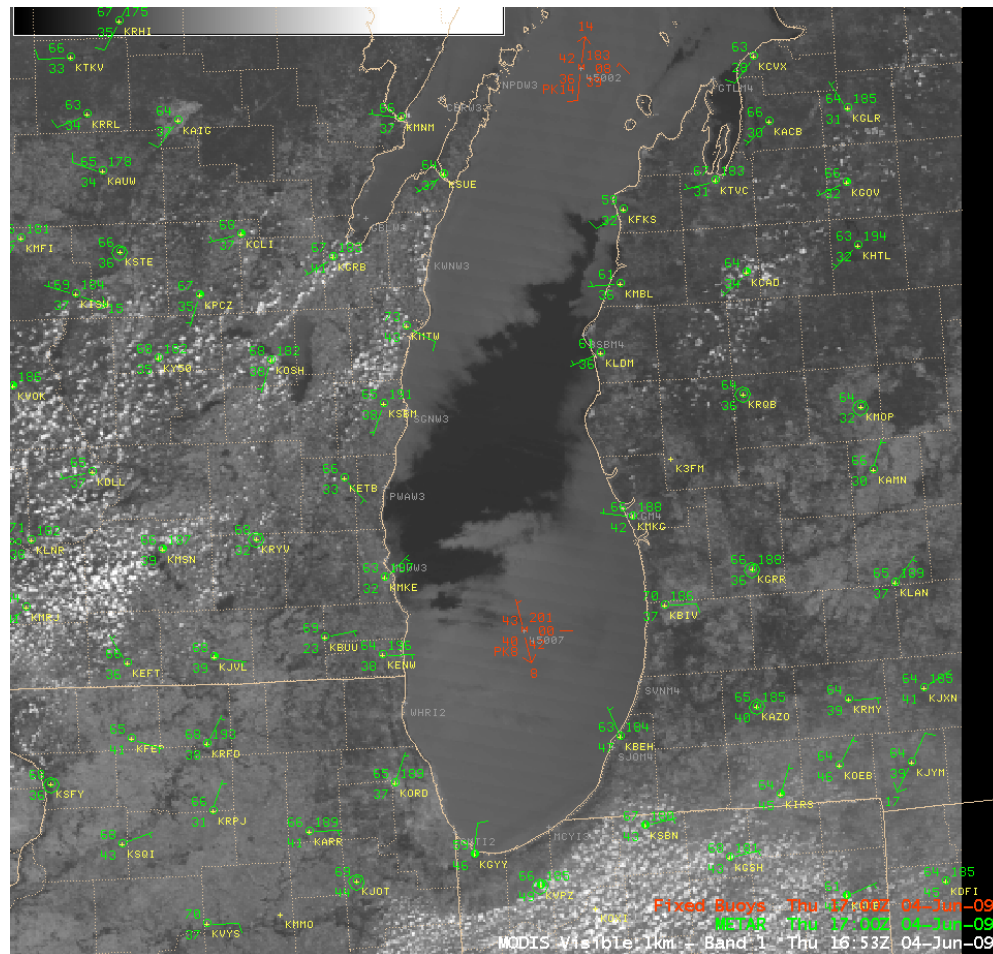
# Sunglint Pattern – HCC Antenna Pass





# Example From Lake Michigan

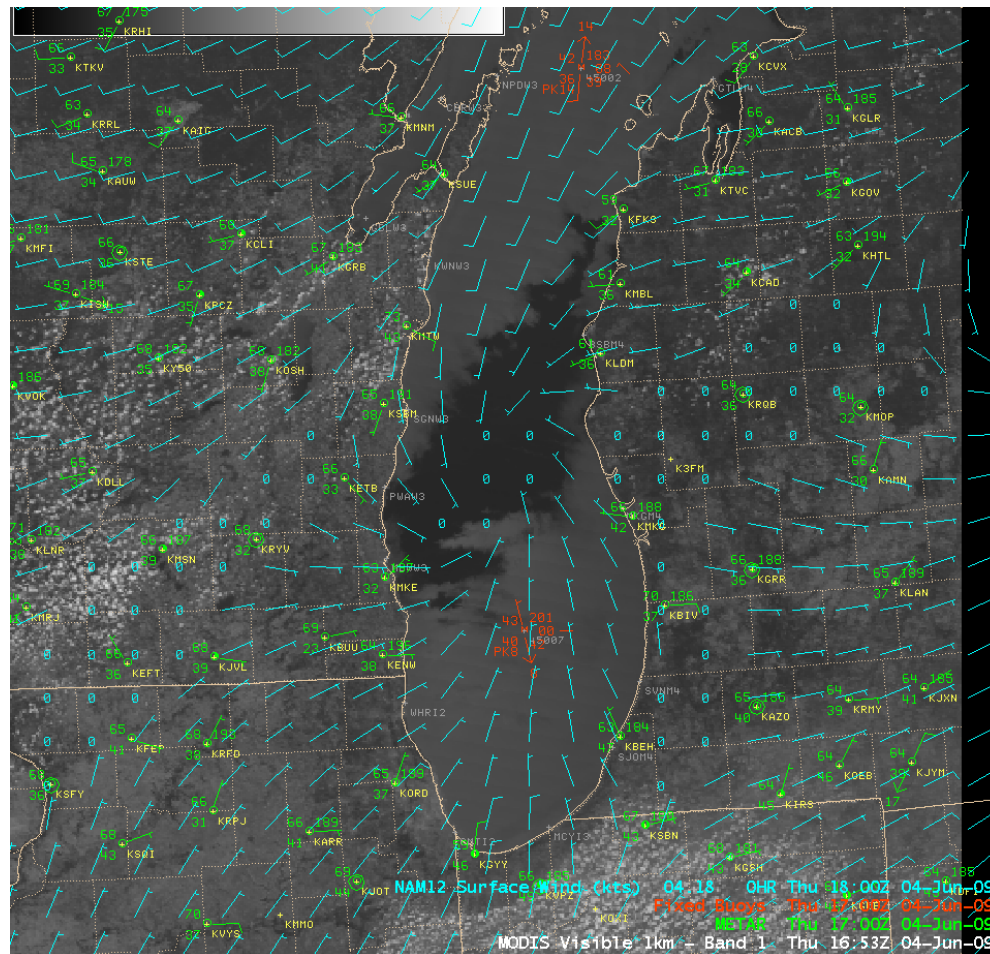
4 June 2009





# Numerical Weather Prediction

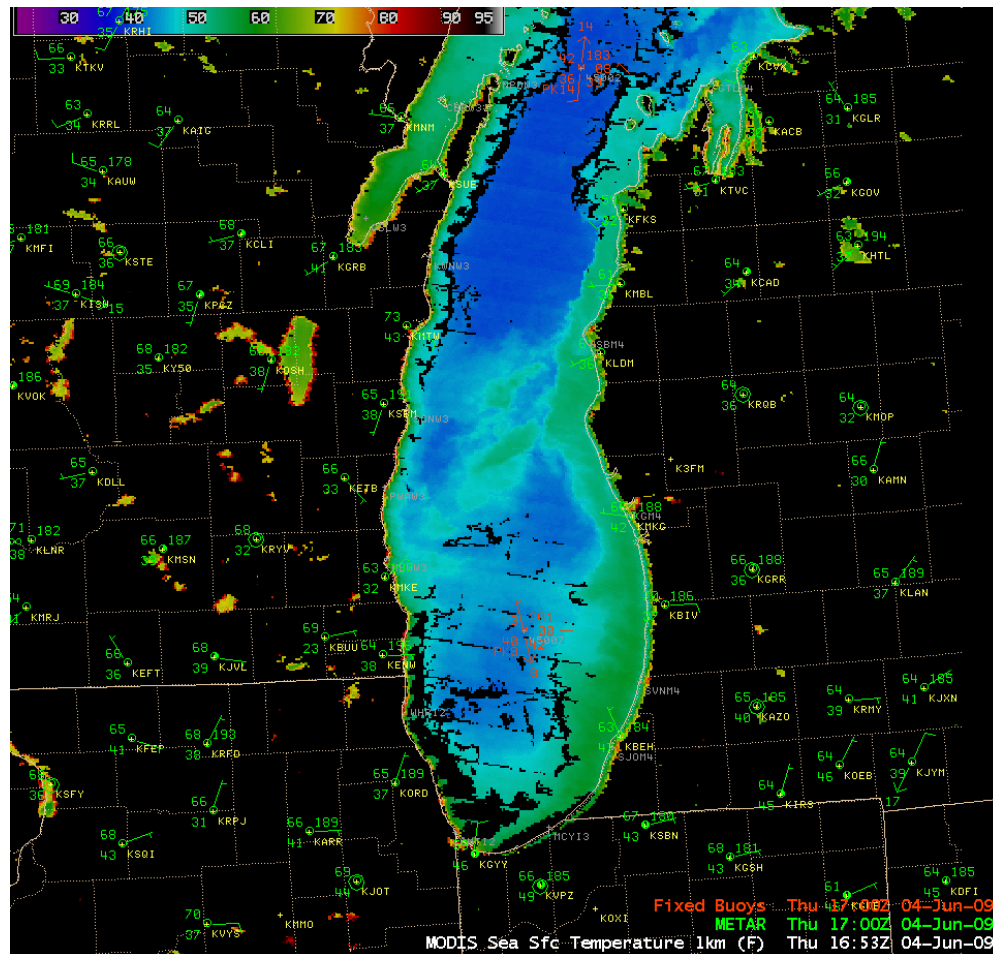
Wind analysis 18 UTC 4 June 2009





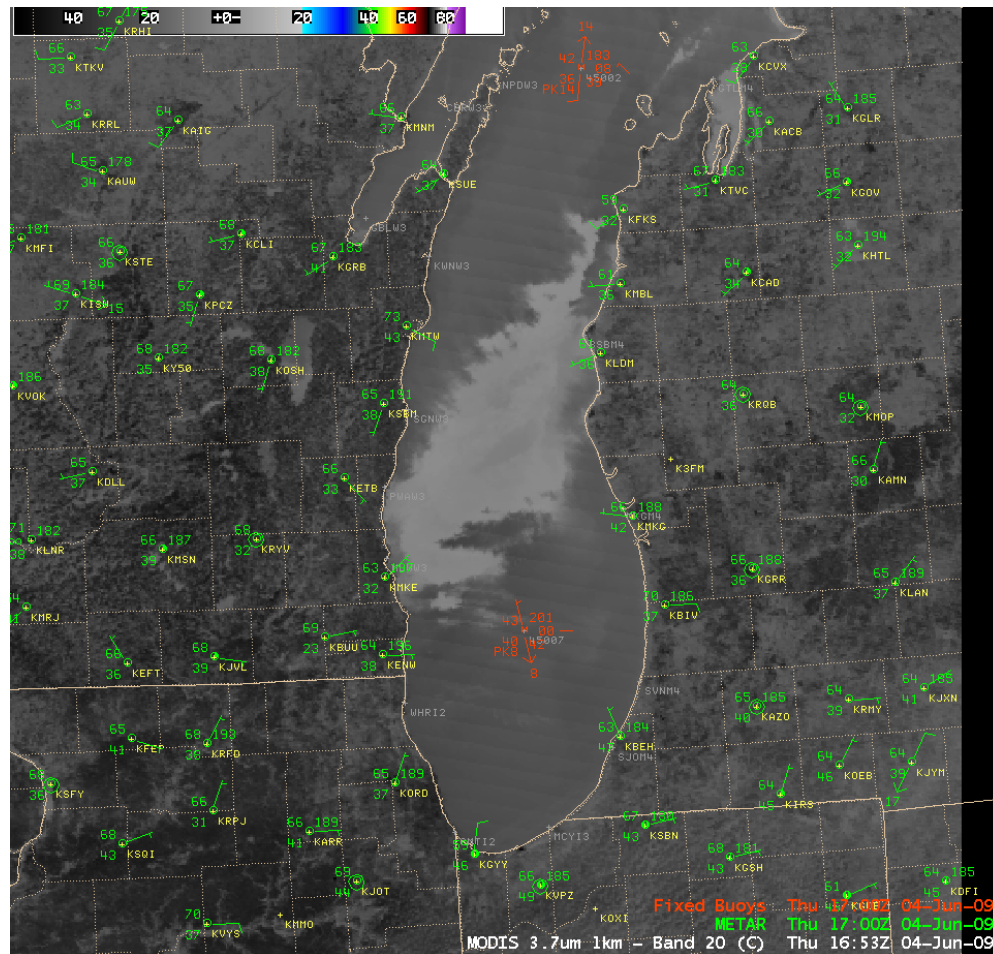
# MODIS Sea Surface Temperatures

4 June 2009





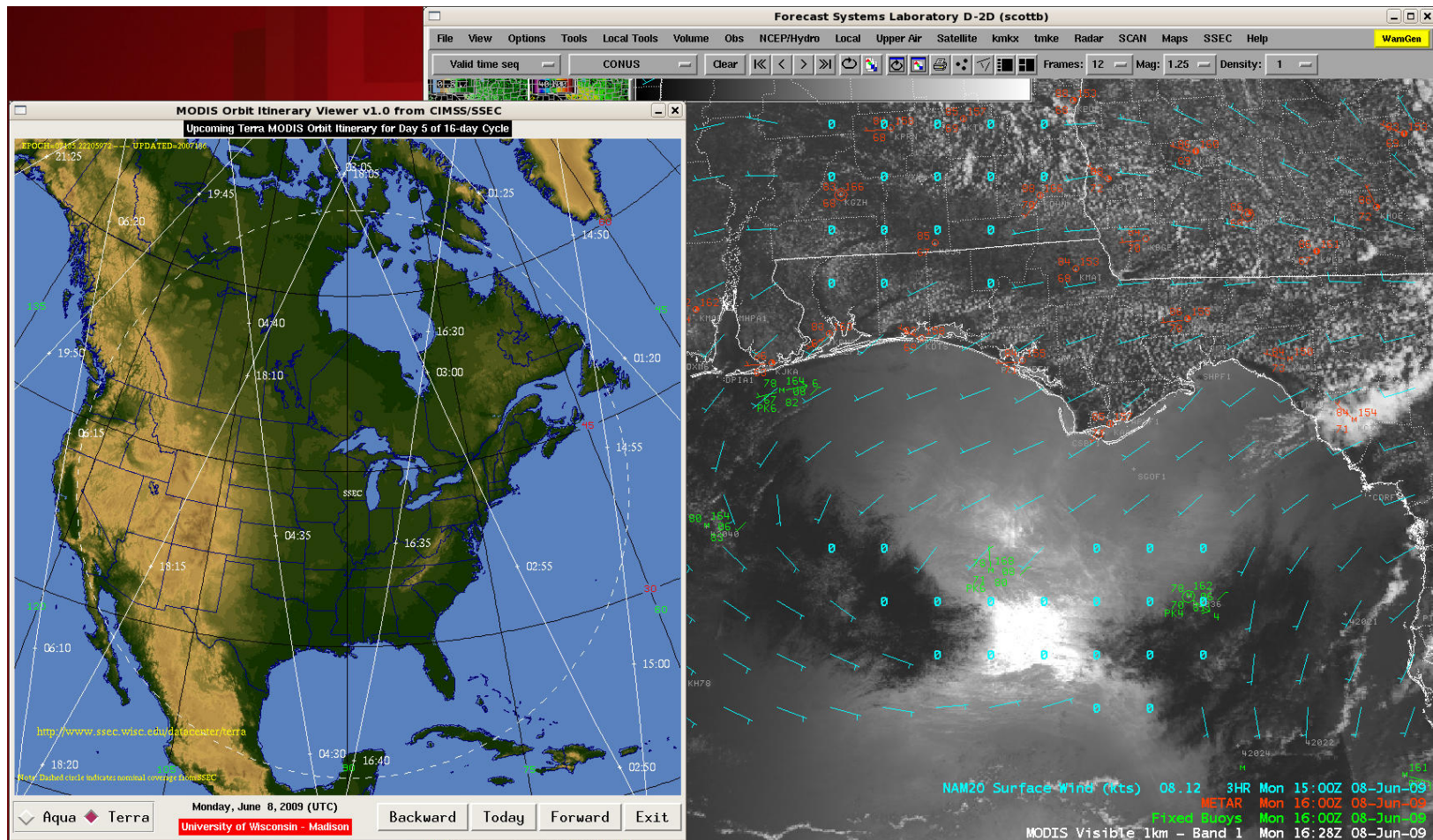
# MODIS 4 $\mu\text{m}$ Brightness Temperatures





# MODIS Sunglint Pattern

8 June 2009





# MODIS LST and buggers

1064

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Vol. 43, no. 5

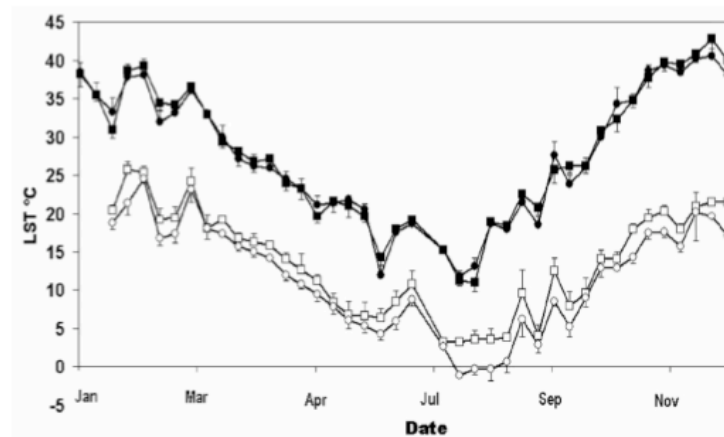


Fig. 4. Annual variation (2003) of diurnal land surface temperature (LST day) produced by the MODIS sensor (closed symbols) and LST night (open symbols) in locality groups of high (squares) and low (circles) house infestation rate. LST values are 8-d composites.

Reference: X. Porcasi, , S. S. Catala, H. Hrellac, M. C. Scavuzzo, D. E. Gorla, 2006: Infestation of Rural Houses by *Triatoma Infestans* (Hemiptera: Reduviidae) in Southern Area of Gran Chaco in Argentina, J. Med. Entomol. 43(5): 1060-1067.