



Retrieval of optical and microphysical properties of ocean constituents using polarimetric remote sensing

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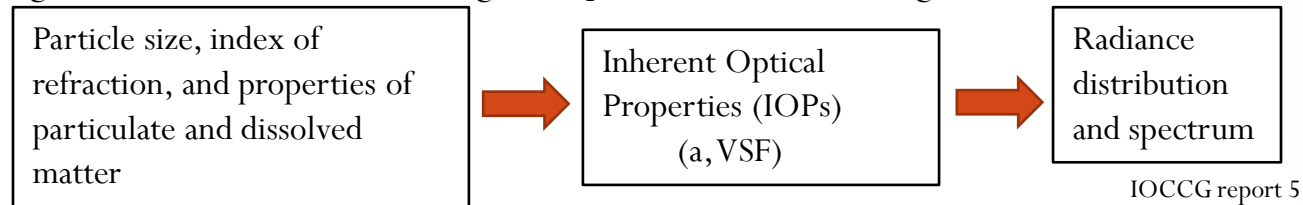
Background

Ocean Color remote sensing

Classical definitions

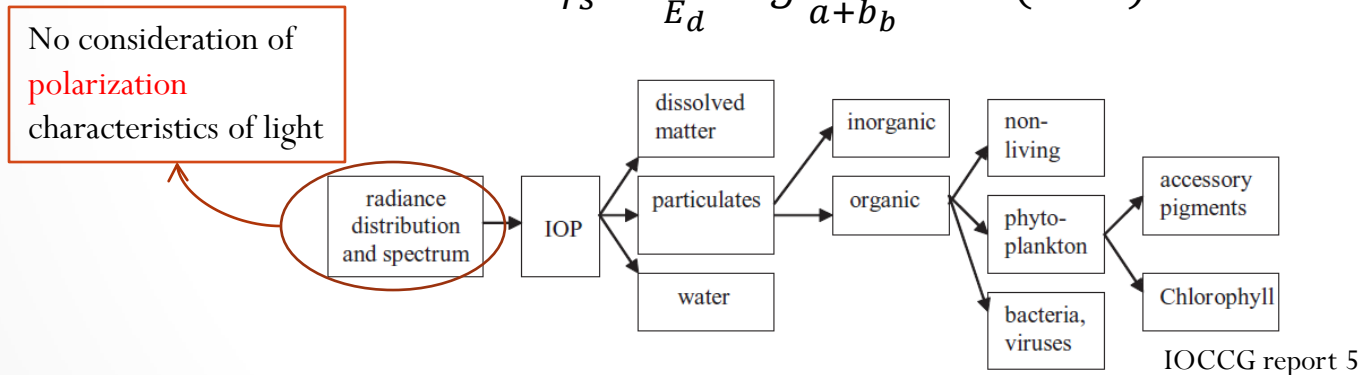
- **Forward model:**

- Oceanic waters embody a large variability of constituents including living and non-living.
- Changes in bulk IOPs lead to changes in spectral radiometric signature of the ocean.



- **Inverse Model:**

$$R_{rs} = \frac{L_w}{E_d} = g \frac{b_b}{a+b_b} \quad (sr^{-1})$$



- Remote sensing reflectance (R_{rs}) is proportional to the the backscattering b_b and inversely to absorption (a) coefficient.
- R_{rs} does not contain information on forwardly scattered light.

g: is related to bidirectionality and geometric structure of the ambient light field

Why polarimetric remote sensing of ocean can be important?

Retrieve additional optical and microphysical properties

- Polarization characteristics of light carry extra information that can be utilized in the retrieval process of IOPs.
- New optical properties of the constituents can be retrieved from polarization in conjunction with standard reflectance retrieval methods.
- Polarization of light is very sensitive to the scattering process.
- Scatterers (hydrosols) and dissolved matter in the ocean have a microphysical and optical properties.
- Polarized light is sensitive to these properties.

Relationship between DOP and IOPs

- Inverse model using polarized light:

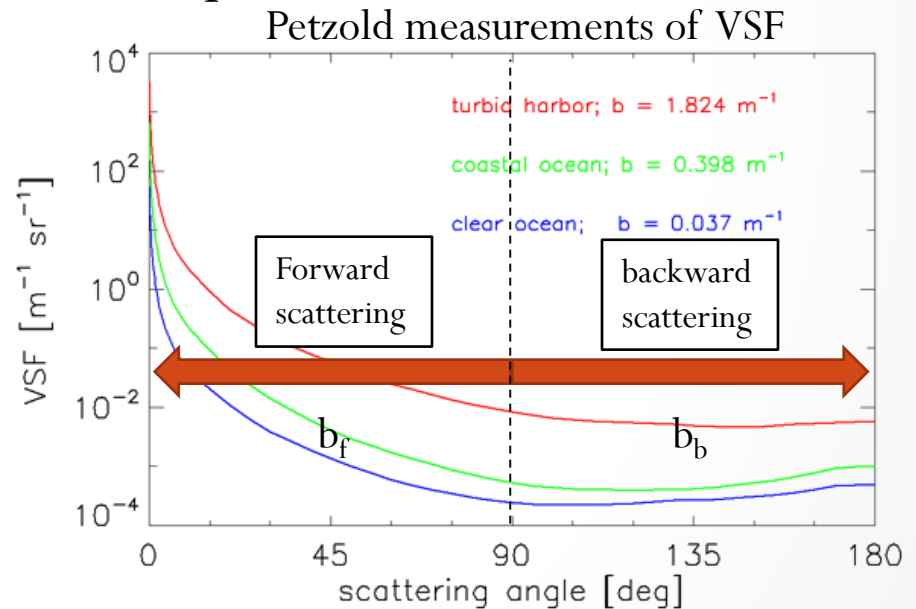
$$R_{rs}(POL) = f\left(\frac{a+b}{a}\right)$$

- $R_{rs}(POL)$ is a function of the absorption (a) and total scattering coefficients (b).

- $b = b_b + b_f$

$$b_f = 2\pi \int_0^{\pi/2} \beta(\theta) \sin(\theta) d\theta \quad \text{and} \quad b_b = 2\pi \int_{\pi/2}^{\pi} \beta(\theta) \sin(\theta) d\theta$$

- Instead of using $R_{rs}(POL)$ we use Degree of Polarization (DOP).



Timofeyeva found a relationship between the DoLP and “the parameter T which is equal to the ratio of the attenuation coefficient of the scattered light flux to the direct light flux” in milky solutions.

V. A. Timofeyeva, “Degree of polarization of light in turbid media,” *Izvestiya Akademii Nauk Sssr Fizika Atmosfery i Okeana*, 6, 513 (1970).

Nature of polarized light

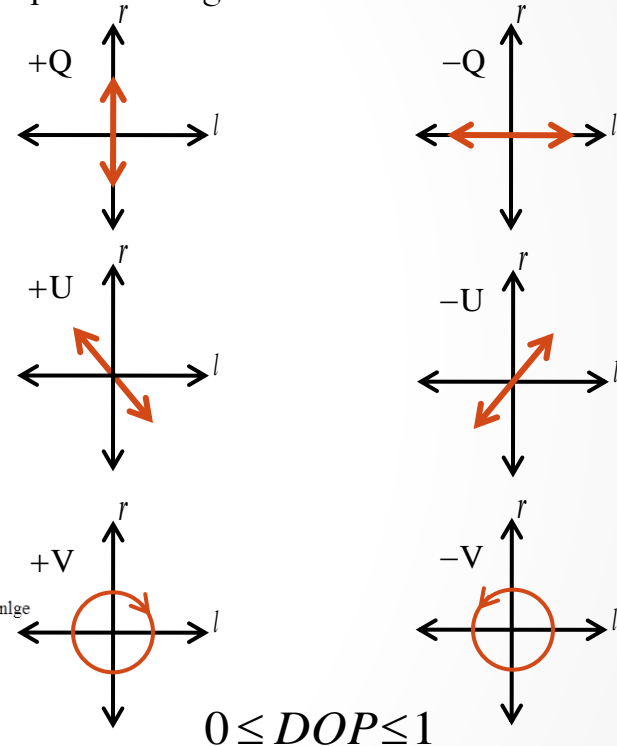
- Light is an electromagnetic wave that propagates through a medium along propagation vector \mathbf{k}

$$\mathbf{E} = (E_l \mathbf{l} + E_r \mathbf{r}) e^{i(\mathbf{kz} - \omega t)}$$

- Stokes components (vector) is a mathematical representation of the polarized light field

$$\begin{bmatrix} I \\ Q \\ U \\ V \end{bmatrix} = \begin{bmatrix} E_l E_l^* + E_r E_r^* \\ E_l E_l^* - E_r E_r^* \\ E_l E_r^* + E_r E_l^* \\ i(E_l E_r^* - E_r E_l^*) \end{bmatrix}$$

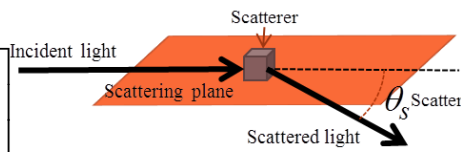
I: the total intensity
 Q: light polarized along 0/90° direction
 U: light polarized along +/ - 45° direction
 V: right or left circularly polarized



- Polarized light is mainly produced by scattering process.
- Mueller Matrix of scatterer:

$$\begin{bmatrix} I' \\ Q' \\ U' \\ V' \end{bmatrix} = \begin{bmatrix} M_{11} & M_{12} & 0 & 0 \\ M_{21} & M_{22} & 0 & 0 \\ 0 & 0 & M_{33} & M_{34} \\ 0 & 0 & M_{43} & M_{44} \end{bmatrix} \begin{bmatrix} I \\ Q \\ U \\ V \end{bmatrix}$$

Scattered light Incident light



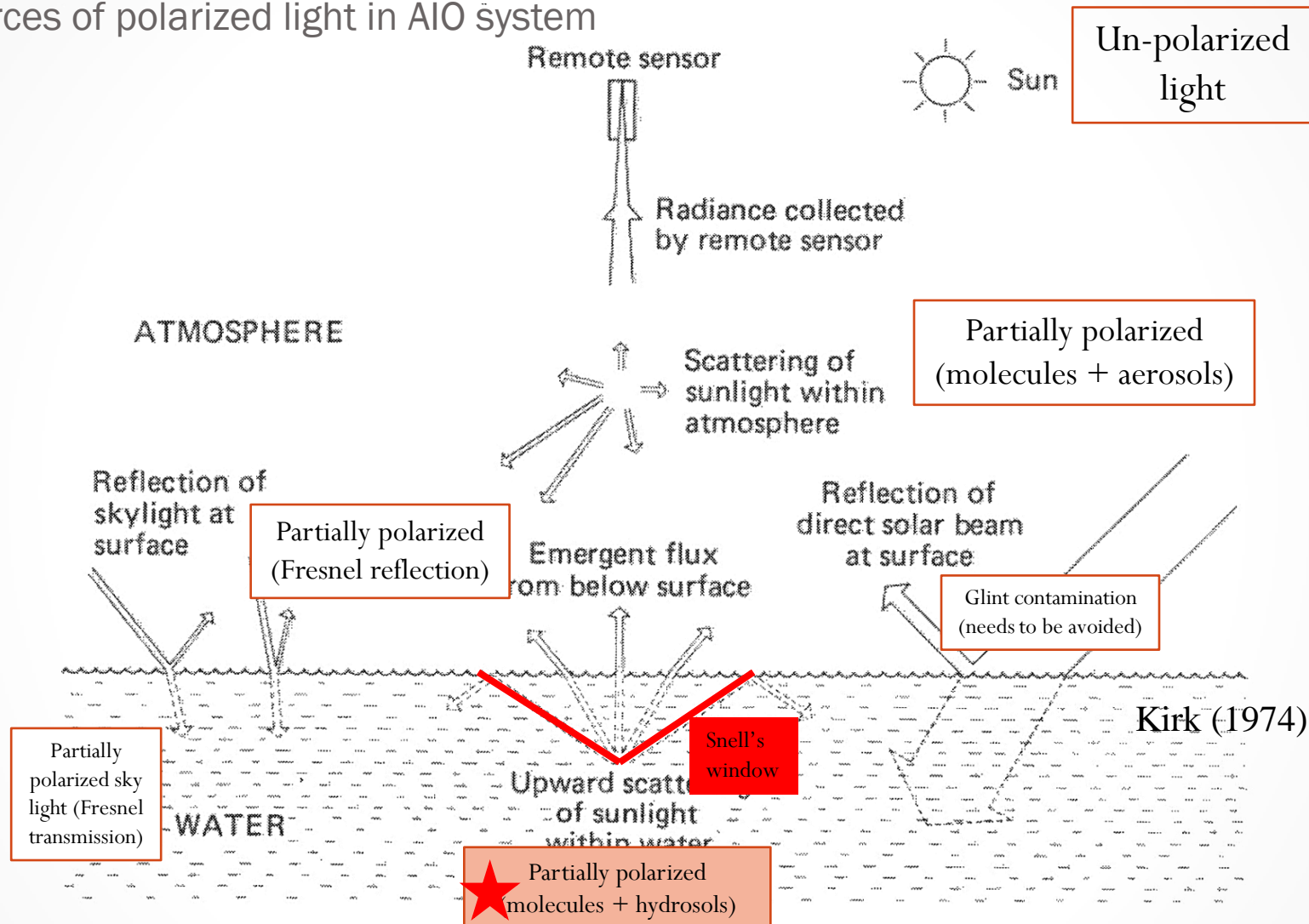
- Degree of Polarization (DOP) is the ratio of the amount of polarized radiation to total radiation.

$$DOP = \frac{\sqrt{Q^2 + U^2 + V^2}}{I}$$

DOP = 0 ⇒ Light is completely unpolarized
 DOP = 1 ⇒ Light is completely polarized

Polarized light field in the ocean

Sources of polarized light in AIO system



- Microphysical properties (particle shape, size, refractive index) affect polarization
- Absorption of water constituents affects polarization

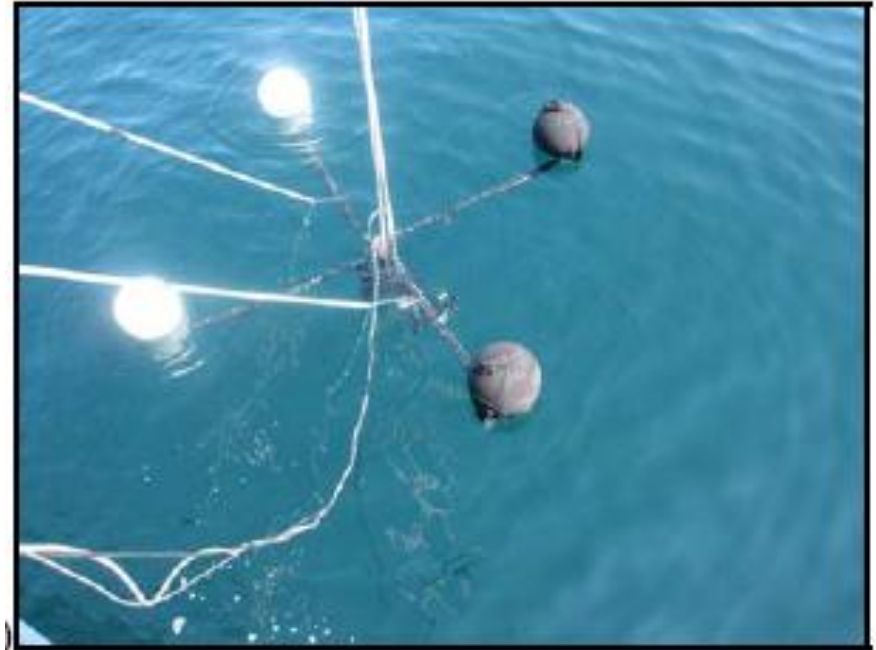
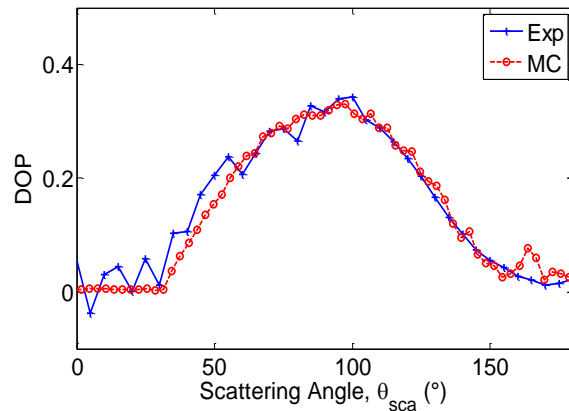
Goal of this study is to relate DOP to the bio-optical properties of the hydrosols IOPs and especially to the ratio of attenuation to absorption (c/a)

Underwater Degree of Polarization (DOP)

Underwater measurements with the hyperspectral multi-angular polarization probe



Station 1, $\lambda=510\text{nm}$

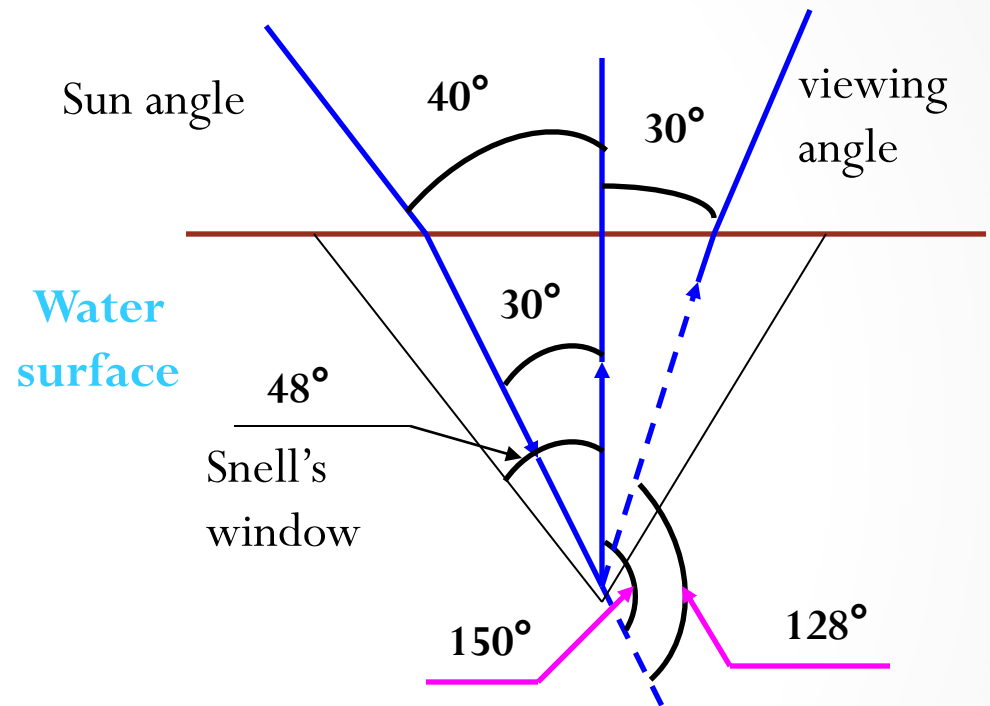
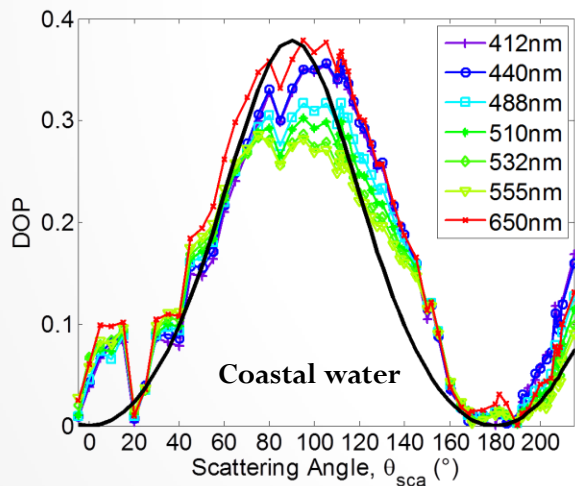
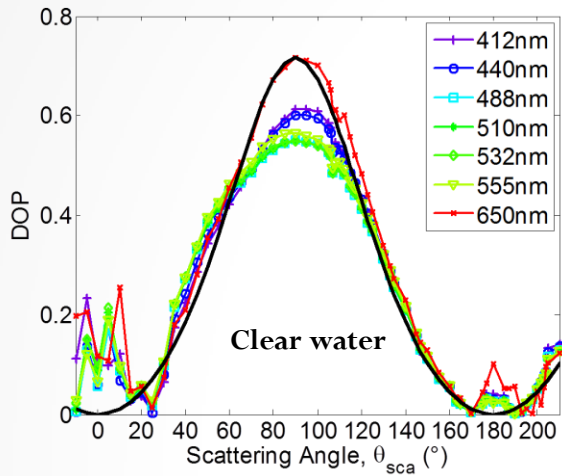


Comparison of measurements and Monte Carlo simulations

Optics Express, 2009, Applied Optics, 2011

Underwater Degree of Polarization (DOP)

Degree of polarization dependence on the geometry (**scattering angle**)



With vertical viewing degree of polarization is substantially lower than $DoLP_{max}$ and is expected to be 0.3-0.4 in the open ocean and 0.15-0.25 in the coastal waters

For above water measurements polarized sky component should be properly eliminated, sunglint should be avoided (measurements outside of the main plane)

Bio-optical model and RT simulations

- Four main constituents assumed in the bio-optical model:



- Each mentioned particle has inherent optical properties (IOPs) which are related to its concentration
- IOPs are measured and modeled based on well established formalism
- Particles also have microphysical characteristics such as, refractive index, size, effective radius

Two main properties of the particles needed in order to simulate for natural water environment:

- IOPs (absorption and scattering coefficients)
- Microphysical parameters (Scattering Matrix for all stokes components)



Inputs
$n_{\text{nap}} = 1.18$
$n_{\text{ph}} = 1.06$
$\xi_{\text{nap}} = 3.5, 4.0, 4.5$
$\xi_{\text{ph}} = 3.5, 4.0, 4.5$
$[\text{Chl}] = 1:20 \text{ mg m}^{-3}$
$[\text{NAP}] = 0.5:10 \text{ g m}^{-3}$
$a_g(400) = 0.3, 0.6, 1, 2, 3 \text{ m}^{-1}$

Bio-optical Model	
Phytoplankton	NAP
$a_{\text{ph}}^* = S_f a_{\text{pico}}^* + (1 - S_f) a_{\text{micro}}^*$	$a_{\text{nap}}(412) = a_{\text{nap}}^*(412) \times [\text{NAP}]$
$S_f = 0.3$	$a_{\text{nap}}^*(412) = 0.05$
$a_{\text{ph}}(\lambda) = a_{\text{ph}}^*(\lambda) \times [\text{Chl}]$	$a_{\text{nap}}(\lambda) = a_{\text{nap}}(412) \times e^{S_{\text{nap}}(412-\lambda)}$
$c_{\text{ph}}(550) = \rho \times [\text{Chl}]^P$	$S_{\text{nap}} = 0.01$
$P = 0.57$	$b_{\text{nap}}(550) = b_{\text{nap}}^*(550) \times [\text{NAP}]$
$\rho = 0.25$	$b_{\text{nap}}^*(550) = 0.6$
$c_{\text{ph}}(\lambda) = c_{\text{ph}}(550) \times \left(\frac{550}{\lambda}\right)^{Y_{\text{ph}}}$	$b_{\text{nap}}(\lambda) = b_{\text{nap}}(550) \times \left(\frac{550}{\lambda}\right)^{Y_{\text{nap}}}$
$Y_{\text{ph}} = \xi_{\text{ph}} - 3$	$Y_{\text{nap}} = \xi_{\text{nap}} - 3 (\text{Approximation})$
$b_{\text{ph}}(\lambda) = c_{\text{ph}}(\lambda) - a_{\text{ph}}(\lambda)$	
CDOM $a_g(\lambda) = a_g(400) e^{-(S_y \times (\lambda - 400))}$	
$S_y = 0.014$	

Mie Calculations	
NAP SM	Phytoplankton SM
$F_{\text{nap}}(\lambda)$	$F_{\text{ph}}(\lambda)$
Mixing scattering matrices	
$F_{\text{Bulk}}(\lambda)$	
$= \frac{b_{\text{nap}}(\lambda) \times F_{\text{nap}}(\lambda) + b_{\text{ph}}(\lambda) \times F_{\text{ph}}(\lambda)}{b_{\text{nap}}(\lambda) + b_{\text{ph}}(\lambda)}$	
Water scattering matrix from standard data bank is mixed through the RT	

$\omega_{\text{sol}}(\lambda) = \frac{b_{\text{nap}}(\lambda) + b_{\text{ph}}(\lambda)}{c_{\text{nap}}(\lambda) + c_{\text{ph}}(\lambda)}$
$c_{\text{nap}}(\lambda) = a_{\text{nap}}(\lambda) + b_{\text{nap}}(\lambda)$
$c_{\text{ph}}(\lambda) = a_{\text{ph}}(\lambda) + b_{\text{ph}}(\lambda)$
$c_{\text{sol}}(\lambda) = c_{\text{nap}}(\lambda) + c_{\text{ph}}(\lambda)$
$\tau_{\text{sol}}(\lambda) = c_{\text{sol}}(\lambda) \times \text{depth}$
$a_g(\lambda), a_w(\lambda), b_w(\lambda)$

Permuted 10125 different cases of IOPs

Radiative Transfer Simulations (RayXP)

Output
$\begin{bmatrix} I \\ Q \\ U \\ V \end{bmatrix} \left\{ \begin{array}{l} \theta_{\text{view}} = 0^\circ:80^\circ \\ \varphi_{\text{sun}} = 0^\circ:360^\circ \\ \theta_{\text{sun}} = 30^\circ \end{array} \right.$

$$DoLP = \frac{\sqrt{Q^2 + U^2}}{I}$$

Notation:
 Subscripts: 'w' for water, 'ph' for phytoplankton, 'nap' for Non-algal particles, and 'g' for CDOM
 'a' is absorption coefficient
 'b' is scattering coefficient
 'c' is attenuation coefficient
 ξ_{ph} & ξ_{nap} are slopes of hyperbolic Particle Size Distribution (PSD)

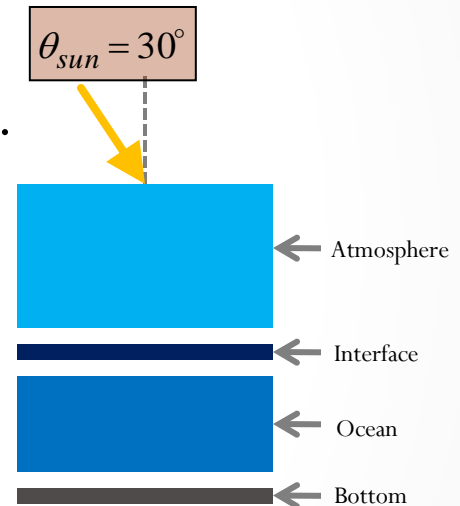
Modeling

Radiative Transfer computations

- Simulations using RayXP program by Zege.
- Optimize computational time by incorporating various techniques of solving the RT equation (**very fast**).
- Plane-parallel homogenous layers for AIO system.

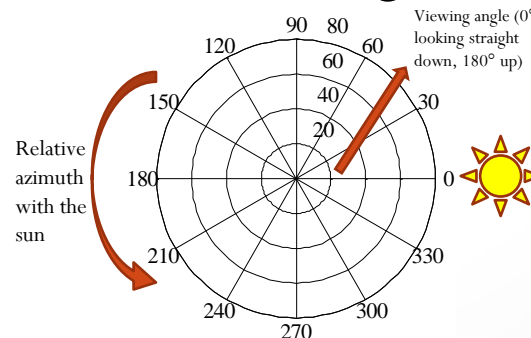
- Assumptions:

- Rayleigh, non-absorbing atmosphere
- Wind ruffled surface (speed of 3 m/s)
- Optically deep waters (no bottom boundary effects)
- Sensor position is just below interface



- Stokes components and DOP calculated for geometries:

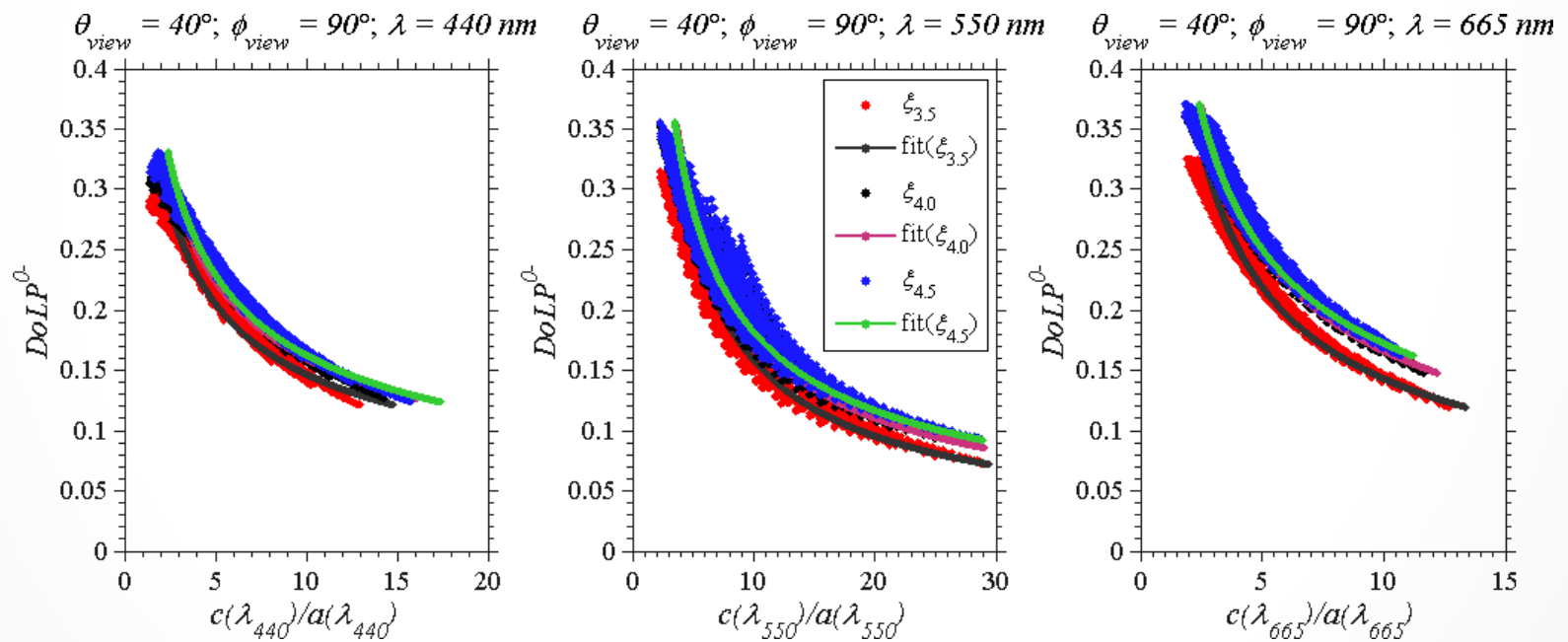
- $\theta_{sun} = 30^\circ$
- $\Phi_{sun} = 0^\circ : 10^\circ : 360^\circ$
- $\theta_{viewing} = 0^\circ : 5^\circ : 180^\circ$



Fitting the relationship

Power law fit

- Parameterizing the relationship can be useful for remote sensing applications.
- Power law fit is a simple representation of the relationship.

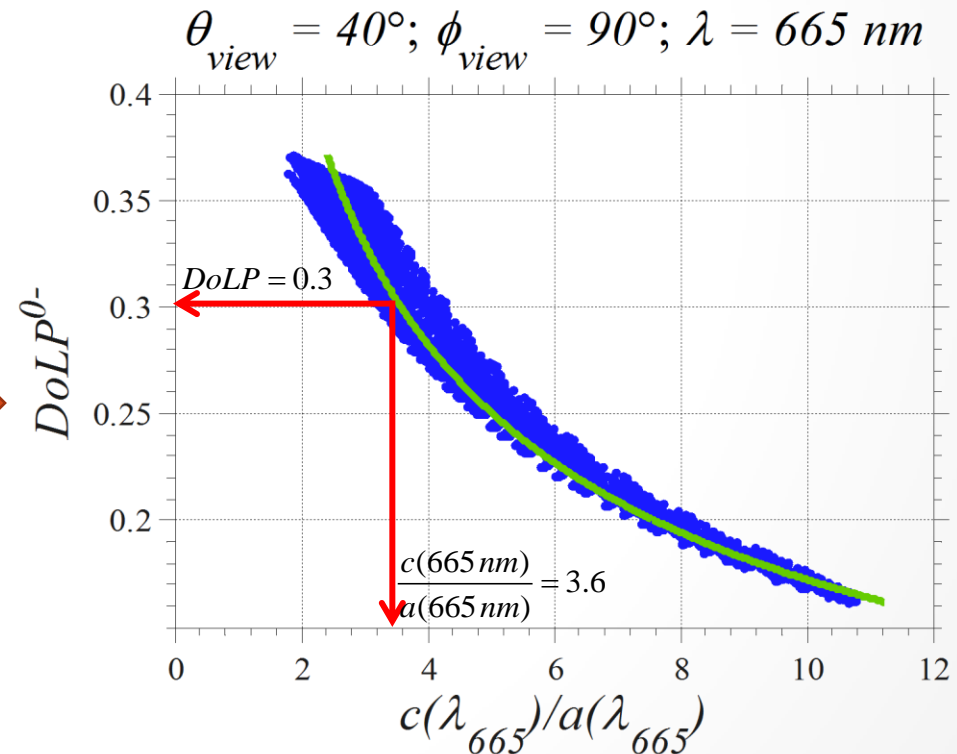
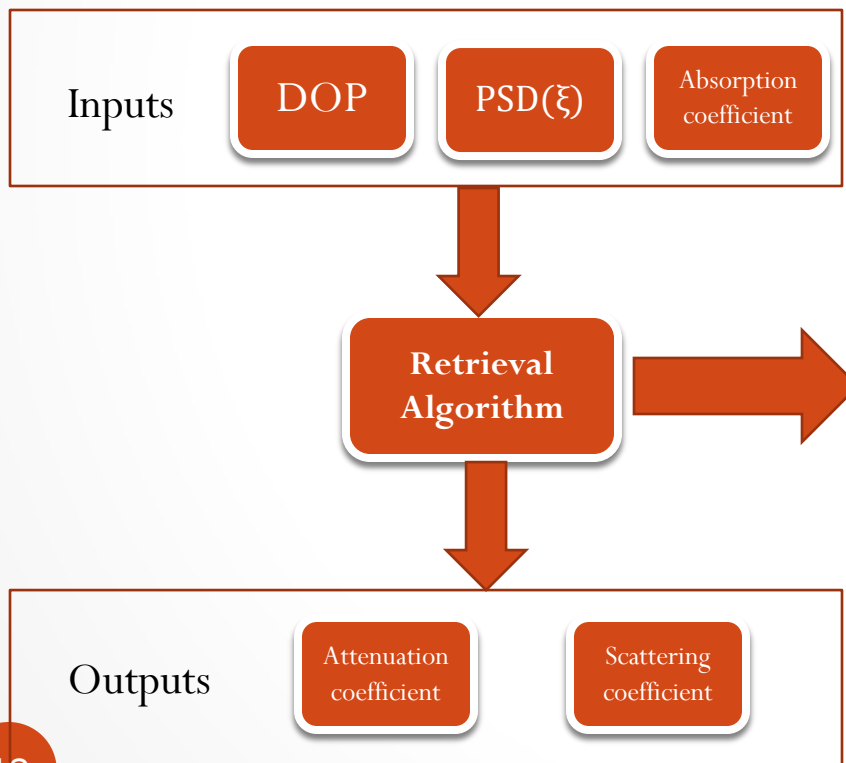


$$\left\{ \left(\frac{c}{a} \right)_{fit} \right\}_{\xi_{nap}=3.5,4.0,4.5} = \left\{ \chi (DoLP)^\gamma \right\}_{\xi_{nap}=3.5,4.0,4.5}$$

Results

Example of Retrieval process

- Attenuation and scattering coefficients can be retrieved from DOP measurements $\frac{c}{a} = \frac{a+b}{a} = 1 + \frac{b}{a}$
- A prior knowledge of PSD and absorption coefficient is needed for retrieval.
- Establishing an algorithm for the retrieval process.

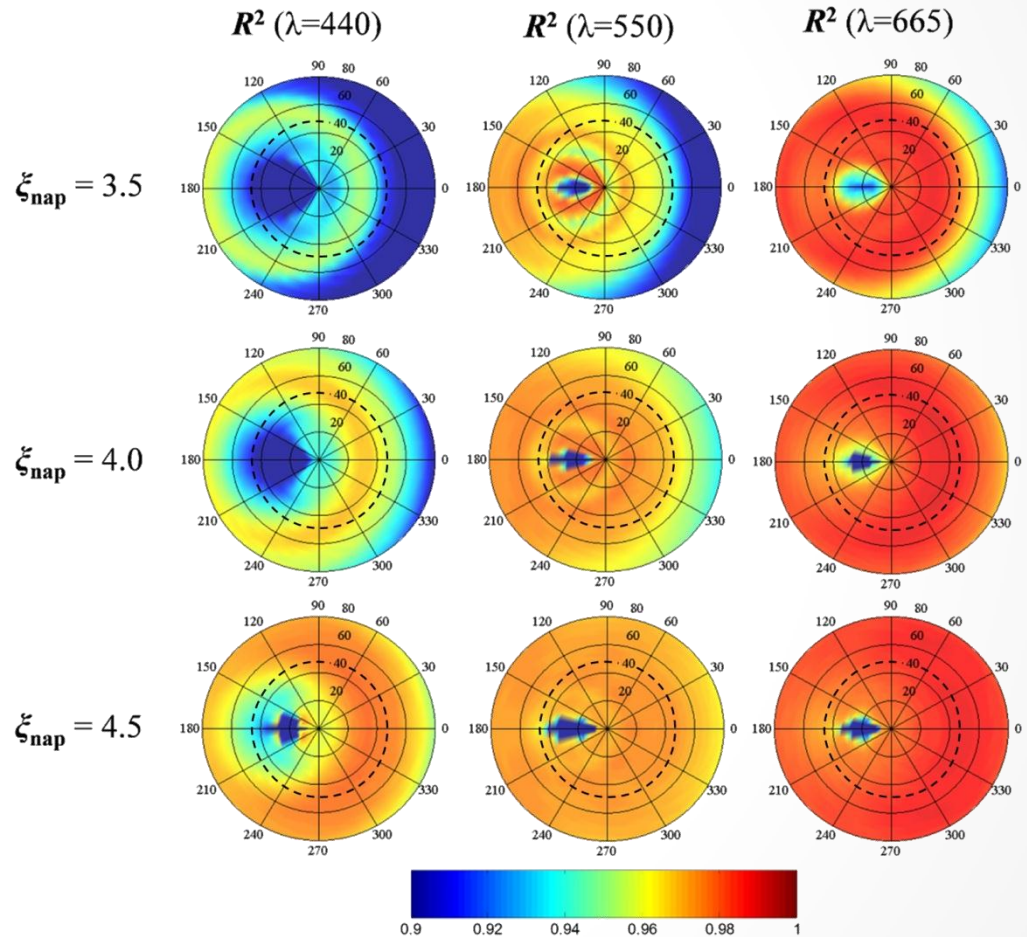


Quality of fitting

Geometrical dependence

$$R^2 = \frac{SSR}{SST} = \frac{\sum_{i=1}^{10125} \left[\left(\frac{c}{a} (DoLP_i) \right)_{fit} - \overline{\left(\frac{c}{a} \right)}_{fit} \right]^2}{\sum_{i=1}^{10125} \left[\left(\left(\frac{c}{a} \right)_i \right) - \overline{\left(\frac{c}{a} \right)}_{fit} \right]^2}$$

- High R^2 values for most of geometries.
- Degradation of sensitivity in the anti-solar plane (backscattering region).
- R^2 values dependence on slope of PSD and wavelength.
- 665 nm shows highest R^2 values.
- ξ_{nap} of 4.5 shows highest R^2 values.

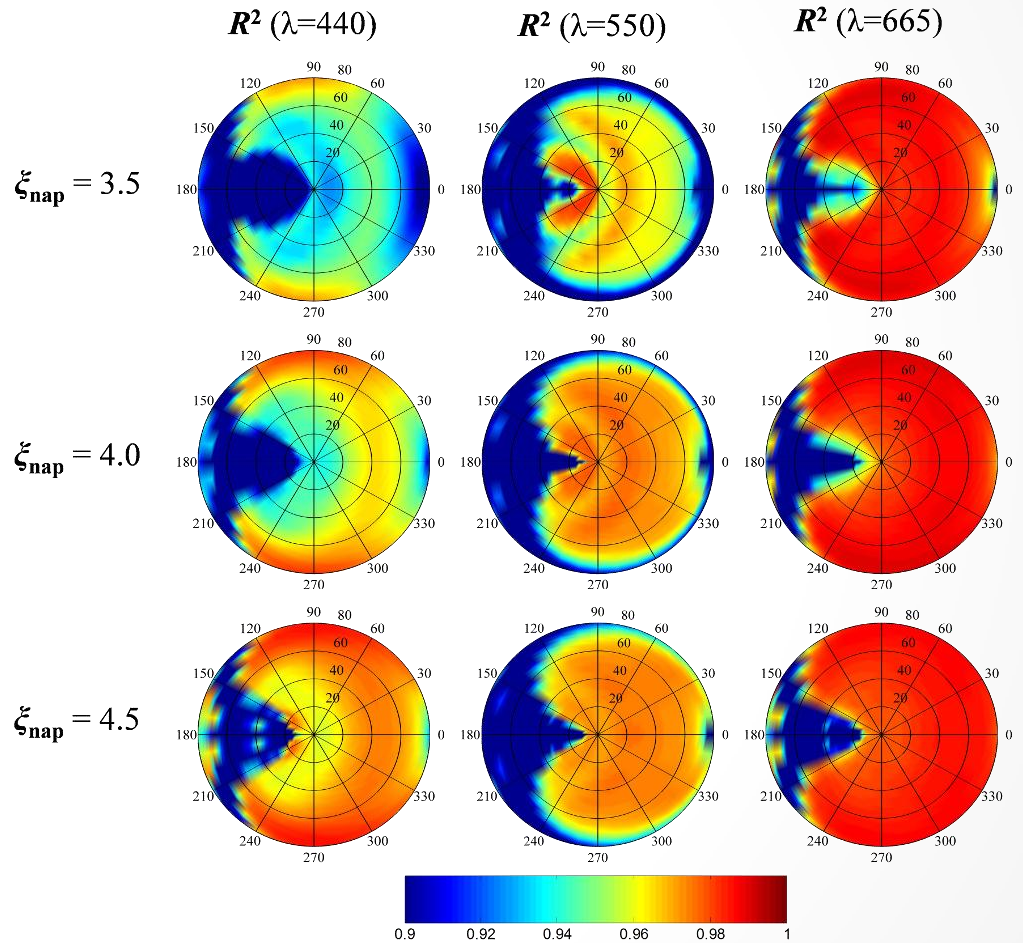


Above water detection

DoLP sensitivity just above water

$$T_{AM} = \begin{bmatrix} \alpha' + \eta' & \alpha' - \eta' & 0 & 0 \\ \alpha' - \eta' & \alpha' + \eta' & 0 & 0 \\ 0 & 0 & \gamma'_{Re} & 0 \\ 0 & 0 & 0 & \gamma'_{Re} \end{bmatrix}$$

- Stokes vector is transmitted to above surface using transmission matrix (kattawar et al.1989).
- Light within Snell's window is transmitted above.
- Low correlation region expands to larger range of viewing geometries (anti-solar).
- High R^2 values for viewing geometries appropriate for remote sensing applications.
- Possible to avoid sun glint contamination.



$$\alpha' = \frac{1}{2} \left[\frac{2 \sin \theta_i \cos \theta_i}{\sin(\theta_i + \theta_r) \cos(\theta_i - \theta_r)} \right]^2,$$

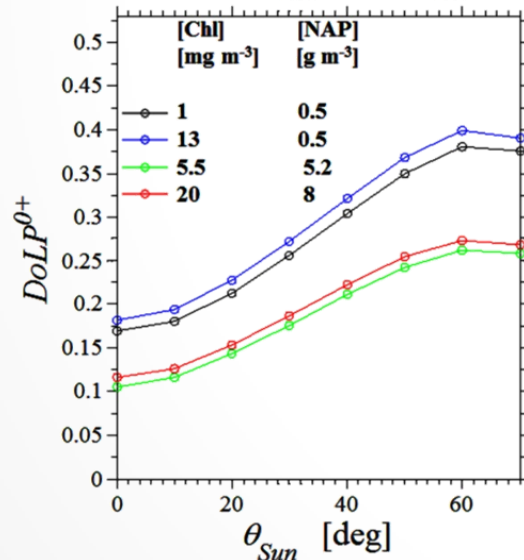
$$\eta' = \frac{1}{2} \left[\frac{2 \sin \theta_i \cos \theta_i}{\sin(\theta_i + \theta_r)} \right]^2,$$

$$\gamma'_{Re} = \frac{4 \sin^2 \theta_i \cos^2 \theta_i}{\sin^2(\theta_i + \theta_r) \cos(\theta_i - \theta_r)} \quad \text{Kattawar (1989)}$$

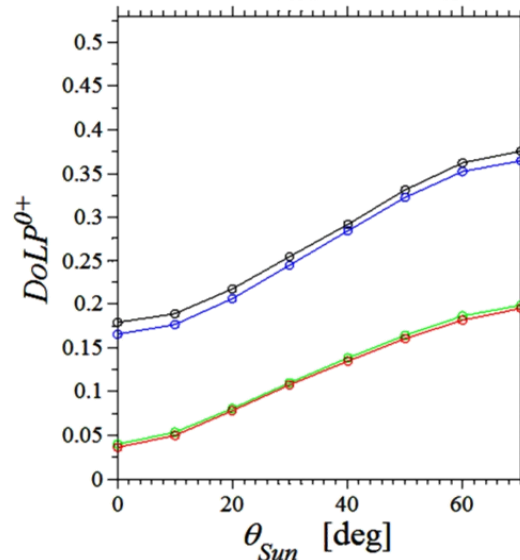
Sun's zenith angle effect

- High dependence of DoLP on sun angle.
- Lower sun angle tends to increase detected DoLP of water leaving radiance.
- Maximal range of DoLP is transmitted to above water surface.
- Dependence of DoLP on the constituents of the water.

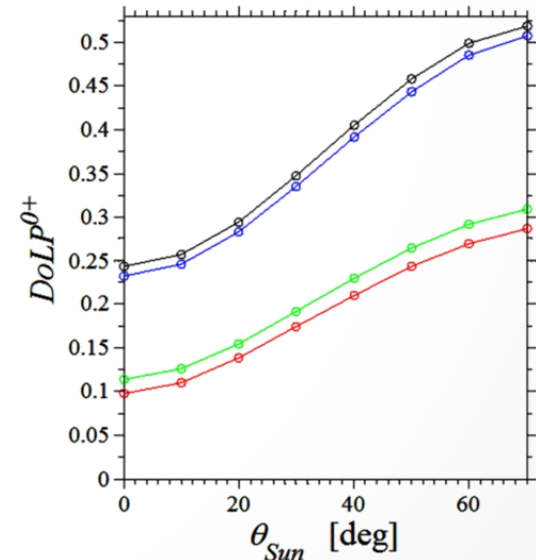
$\lambda = 440 \text{ nm}; \theta_{\text{view}} = 60^\circ; \phi_{\text{view}} = 90^\circ$



$\lambda = 550 \text{ nm}; \theta_{\text{view}} = 60^\circ; \phi_{\text{view}} = 90^\circ$



$\lambda = 665 \text{ nm}; \theta_{\text{view}} = 60^\circ; \phi_{\text{view}} = 90^\circ$



Conclusions

- We simulated the stokes components for 10125 different cases of bio-optical and microphysical properties of the ocean's constituents.
- We explored the relationship between DoLP and IOPs (c/a).
- The relationship was parameterized/fitted easily with a power law fit.
- We showed the possibility to retrieve hydrosol's attenuation to absorption (c/a) ratio from DoLP measurements leading to the retrieval of attenuation and total scattering coefficients.
- Synoptic view of the R^2 between fitted and simulated relationship for below and above water have been presented.
- We showed the best range of angles (lowest R^2) for retrieval purposes.